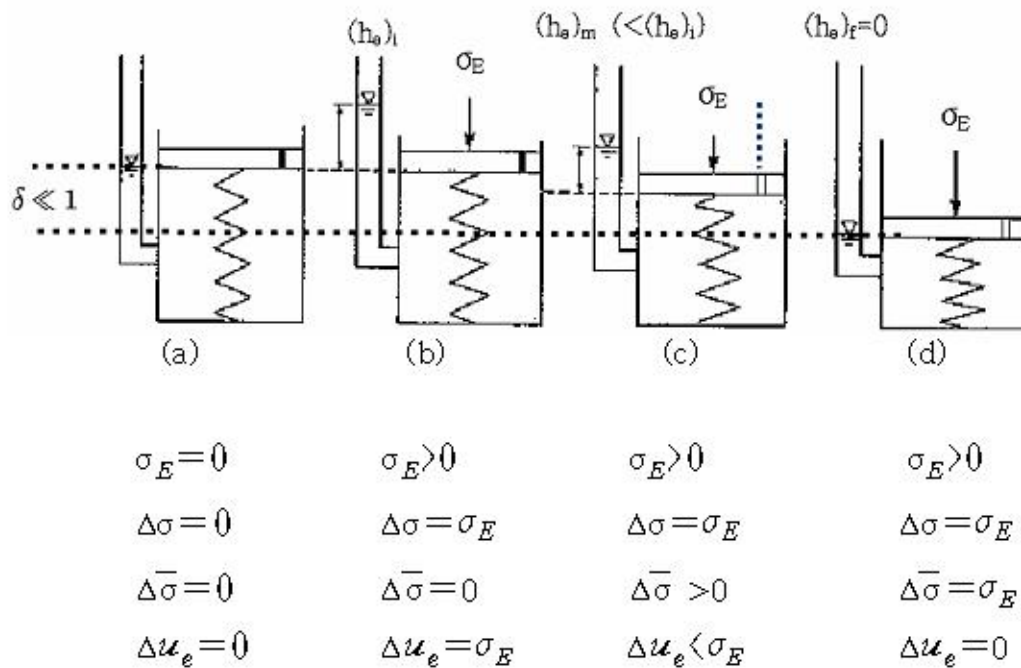


## Chapter 5. Consolidation (outlines)

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## Chapter 5. Consolidation

### 1. Spring analogy (Fig. 2.17, p.57)

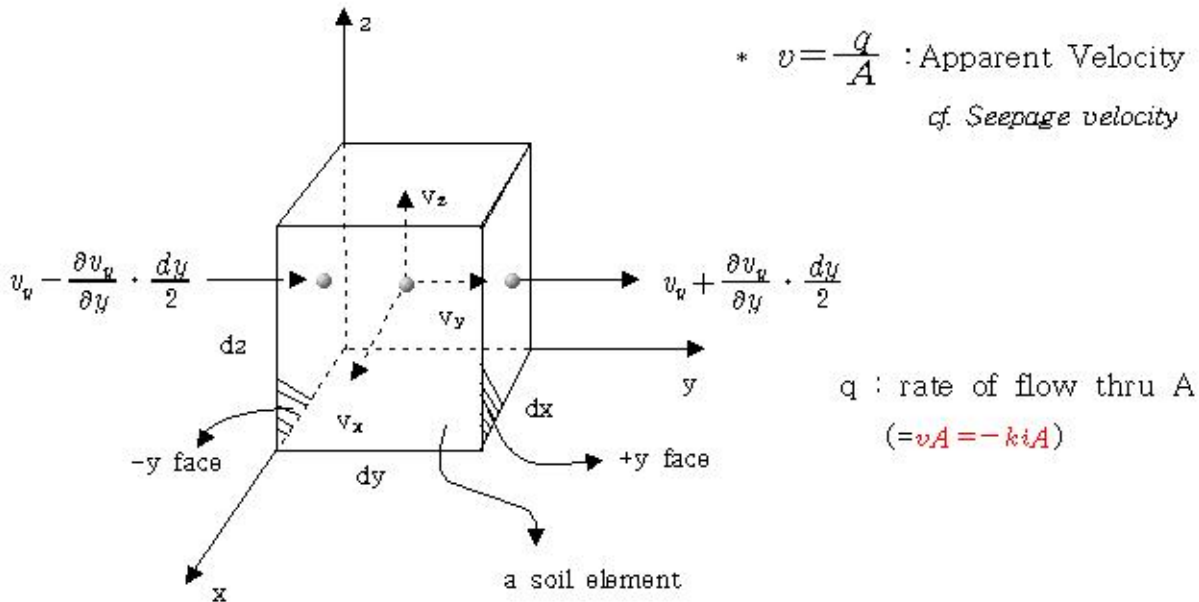


\* Consolidation ?

The gradual reduction in volume of a fully saturated soil of low permeability due to drainage of some of the ( *pore pressure* ), ( *the process* ) continuing until the ( *excess pore water pressure* ) set up by an increase in total stress has completely dissipated : The simplest case is that of one-dimensional consolidation in which a condition of ( *no lateral strain* ) is implicit.

**2. Theory of Consolidation**

( 5.2 , Terzaghi's one-dimensional consolidation equation )



- Rate of flow out of the element thru the +y face

$$q_{out} = \left( v_y + \frac{\partial v_y}{\partial y} \frac{dy}{2} \right) dx \cdot dz \quad \dots \dots \dots (1)$$

- Rate of flow into the element thru the -y face

$$q_{in} = \left( v_y - \frac{\partial v_y}{\partial y} \frac{dy}{2} \right) dx \cdot dz \quad \dots \dots \dots (2)$$

- net rate of flow into the element in the y direction

$$q_{in} - q_{out} = (2) - (1) = \left( -\frac{\partial v_y}{\partial y} \right) dx \cdot dy \cdot dz \quad \text{[check (-) sign]}$$

flow quantity per time

- Total weight change of flux due to 3-D flow into the element

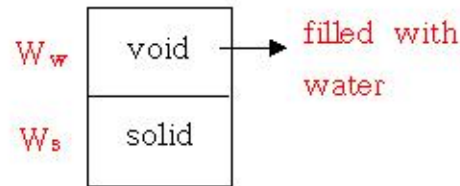
$$= -\gamma_w \left( \frac{\partial v_x}{\partial x} + \frac{\partial v_y}{\partial y} + \frac{\partial v_z}{\partial z} \right) dx \cdot dy \cdot dz$$

- For a soil element,

$$W_T = W_s + W_w$$

per unit volume of soil

$$W_T/V = W_s/V + W_w/V$$



- Let  $W_w/V = W_w^*$  : wt. of water per unit vol. of soil (흙의 단위 부피당 물의 무게)

- Then,

$$-v_w \left( \frac{\partial v_x}{\partial x} + \frac{\partial v_y}{\partial y} + \frac{\partial v_z}{\partial z} \right) \overbrace{dxdydz}^{\text{one}} / V = \frac{\partial W_w^*}{\partial t}$$

↓  
[ Eq. of continuity ]

- If ( *steady-state* ) condition exists,

$$\frac{\partial W_w^*}{\partial t} = 0 \quad : \text{i.e., ( no accumulation or no dissipation of water )}$$

$$\rightarrow \frac{\partial v_x}{\partial x} + \frac{\partial v_y}{\partial y} + \frac{\partial v_z}{\partial z} = 0 \quad : \text{steady-state seepage equation}$$

- If ( *transient* ) case,

$$-\left( \frac{\partial v_x}{\partial x} + \frac{\partial v_y}{\partial y} + \frac{\partial v_z}{\partial z} \right) = \frac{1}{v_w} \frac{\partial W_w^*}{\partial t} \quad \dots \dots \text{Eq. A}$$

*Assumption #1*

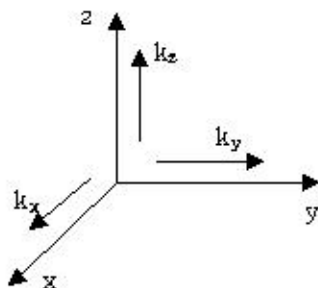
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- Darcy's law - holds for water flow (laminar) in soil

$$v \propto i \rightarrow v = -ki$$

$$i = \frac{dh}{dl} : \text{hydraulic gradient}$$

- In three-dimensional & anisotropic (general case)



$$k : k_x, k_y, k_z$$

$$v_x = -k_x \frac{\partial h}{\partial x}, \quad v_y = -k_y \frac{\partial h}{\partial y}, \quad v_z = -k_z \frac{\partial h}{\partial z}$$

- Substitute these into Eq. (A)

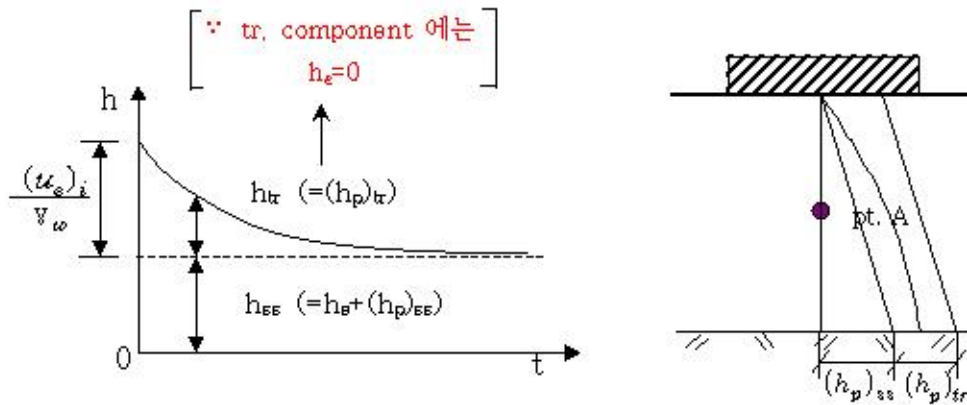
$$\frac{\partial}{\partial x} \left( k_x \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left( k_y \frac{\partial h}{\partial y} \right) + \frac{\partial}{\partial z} \left( k_z \frac{\partial h}{\partial z} \right) = \frac{1}{\gamma_w} \frac{\partial W_w^*}{\partial t}$$

- Assuming  $k_x, k_y, k_z$  are constant w.r.t.  $x, y, z$  coordinates

[ i.e., homogeneous ].

*Assumption #2*

$$k_x \frac{\partial^2 h}{\partial x^2} + k_y \frac{\partial^2 h}{\partial y^2} + k_z \frac{\partial^2 h}{\partial z^2} = \frac{1}{\gamma_w} \frac{\partial W_w^*}{\partial t} \quad \dots \dots \text{Eq. B}$$



<at pt. A>

$$\begin{aligned}
 - \text{ Since } h_t &= h_e + h_p = \overbrace{h_e}^{h_{ss}} + ((h_p)_{ss}) + ((h_p)_{tr}) \\
 &= (h_{ss}) + ((h_p)_{tr}) = h_{ss} + \left( \frac{u_e}{\gamma_w} \right)
 \end{aligned}$$

where  $u_e$  : excess pore water pressure

$\frac{\partial^2}{\partial x^2}$   
 $\downarrow$   
 linear operator

- Sub. this into Eq. B, then L.H.S. becomes

$$\begin{aligned}
 k_x \frac{\partial^2 h_{ss}}{\partial x^2} + k_y \frac{\partial^2 h_{ss}}{\partial y^2} + k_z \frac{\partial^2 h_{ss}}{\partial z^2} \\
 + \frac{k_x}{\gamma_w} \frac{\partial^2 u_e}{\partial x^2} + \frac{k_y}{\gamma_w} \frac{\partial^2 u_e}{\partial y^2} + \frac{k_z}{\gamma_w} \frac{\partial^2 u_e}{\partial z^2} = \frac{k}{\gamma_w} \nabla^2 u_e, \quad k_x = k_y = k_z
 \end{aligned}$$

isotropic Assumption #3

where  $\nabla^2$  : Laplacian operator  $\left( = \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2} + \frac{\partial^2}{\partial z^2} \right)$

- R.H.S  $\left( = \frac{1}{\gamma_w} \frac{\partial W_w^*}{\partial t} \right)$

$$W_w^* = \frac{\gamma_w V_w}{V} = \gamma_w \cdot S \cdot V_v / V \quad (\because S = \frac{V_w}{V_v})$$

$$W_s^* = \gamma_w \cdot G_s \cdot V_s / V$$

$$\frac{W_w^*}{W_s^*} = \frac{S \cdot V_v}{G_s \cdot V_s} = \frac{S \cdot e}{G_s} \quad (\because e = \frac{V_v}{V_s})$$

- Thus,

$$\begin{aligned}
 W_w^* &= \left( \frac{S \cdot e}{G_s} \right) W_s^* \\
 &= \left( \frac{e}{G_s} \right) W_s^* \leftarrow \left( \text{Assume } S=1, \text{ fully saturated} \right) \\
 &\hspace{10em} \text{Assumption \# 4}
 \end{aligned}$$

and, we get

$$\frac{1}{V_w} \frac{\partial W_w^*}{\partial t} = \frac{1}{V_w} \cdot \frac{W_s^*}{G_s} \frac{\partial e}{\partial t}$$

here,

$$\begin{aligned}
 \frac{W_s^*}{V_w \cdot G_s} &= \frac{\text{wt. of solid / vol. of soil}}{\text{unit wt. of solid}} \\
 &= \frac{\text{wt. of solid}}{(\text{unit wt. of solid})(\text{vol. of soil})} \\
 &= \frac{\text{vol. of solid}}{\text{vol. of soil}} = \frac{1}{1 + e_0}
 \end{aligned}$$

$e_0$	$V_v$
1	$V_s$

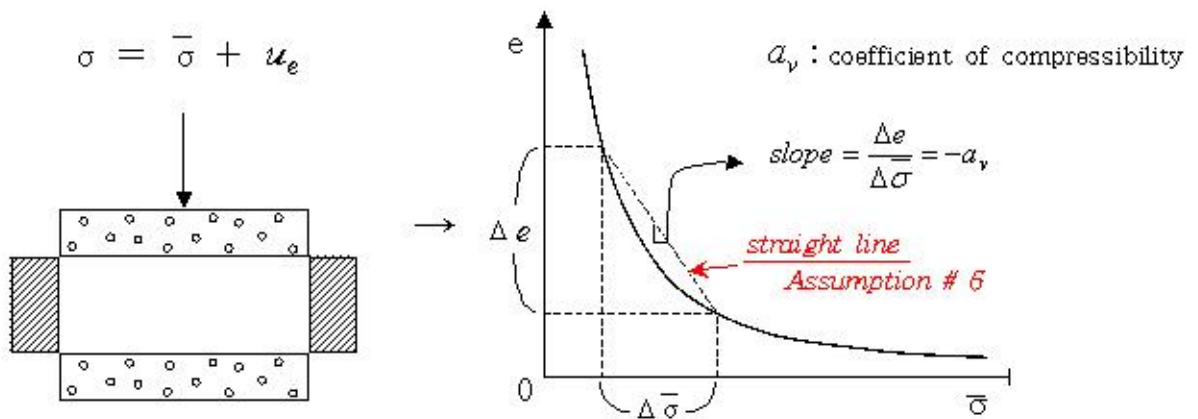
↑  
initial void ratio

↑  
small strain assumed

Assumption # 5

\*  $\frac{\partial e}{\partial t} = ?$

- from ( *Oedometer* ) test



-  $\Delta e = -a_v \Delta \bar{\sigma}$

$$\begin{aligned} \frac{\partial e}{\partial t} &= -a_v \frac{\partial \bar{\sigma}}{\partial t} = -a_v \frac{\partial (\sigma - u_e)}{\partial t} \\ &= +a_v \frac{\partial u_e}{\partial t} - a_v \frac{\partial \sigma}{\partial t} \\ &= a_v \left( \frac{\partial u_e}{\partial t} - \frac{\partial \sigma}{\partial t} \right) \end{aligned}$$

- Finally,

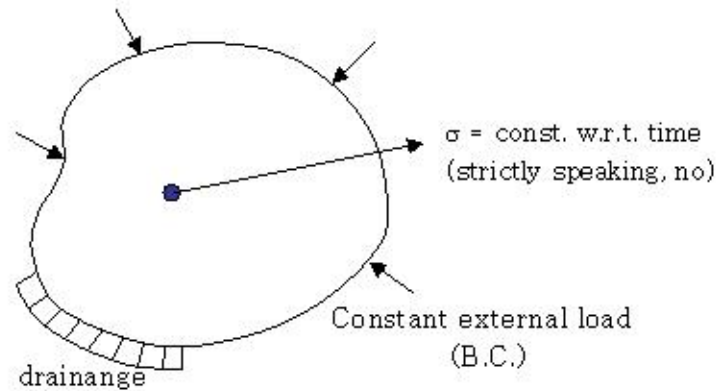
$$\frac{k}{\gamma_w} \cdot \nabla^2 u_e = \frac{a_v}{1+e_0} \left( \frac{\partial u_e}{\partial t} - \frac{\partial \sigma}{\partial t} \right)$$

or

$$\nabla^2 u_e = \frac{\gamma_w \cdot a_v}{k(1+e_0)} \left( \frac{\partial u_e}{\partial t} - \frac{\partial \sigma}{\partial t} \right) \dots\dots\dots \text{Eq.C}$$



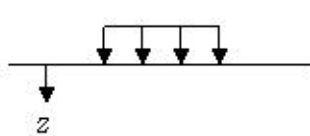
-  $\frac{\partial \sigma}{\partial t} = ?$



- Assume  $\frac{\partial \sigma}{\partial t} = 0$  , i.e., total stress  $\sigma = \text{constant}$  for constant boundary loading  
*Assumption # 7*

- Assume one dimensional flow - *Assumption # 8*

then, Eq. C becomes,



$$\frac{\partial^2 u_e}{\partial z^2} = \frac{1}{C_v} \frac{\partial u_e}{\partial t} \quad \dots \text{1-dim. consolidation Eq.}$$

$u_e$  : excessive pore water pressure

$$C_v = \frac{k(1+e_0)}{a_v \cdot \gamma_w} \quad ; \text{ coefficient of consolidation}$$

( $L^2 T^{-1}$ )

assumed to be ( *constant* ) for a certain soil  
*Assumption # 9*

$$m_v = \frac{a_v}{1+e_0} \quad ; \text{ coefficient of Volume compressibility}$$

$$\rightarrow C_v = \frac{k}{m_v \gamma_w}$$