

### 457.562 Special Issue on River Mechanics (Sediment Transport) .05 Macroscopic view of sediment transport



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# 1. Submerged Angle of Repose

- If granular particles are allowed to pile up while submerg ed in a fluid, there is a specific slope angle beyond which spontaneous failure of the slope occurs.
- This angle is the angle of repose (or friction angle).
- The coefficient of Coulomb friction

 $\mu = \frac{\text{tangential resistive force}}{\text{downward normal force}}$ 







# 1. Submerged Angle of Repose

- The forces acting on the particle along the slope are:
  - The submerged force of gravity, which has a downslope component F<sub>gt</sub> and normal component F<sub>gn</sub>
  - A tangential resistive force  $F_r$  due to Coulomb friction.

$$F_{gt} = (\rho_s - \rho)gV_p \sin\phi$$
$$F_{gn} = (\rho_s - \rho)gV_p \cos\phi$$
$$F_r = \mu F_{gn}$$

The condition for incipient motion is

$$F_{gt} = F_r$$

 Down slope impelling force of gravity should just barely bal ance with the Coulomb resistive force.



- 1. Submerged Angle of Repose
- From the previous four relations, it is found  $\mu = \tan \phi$
- The angle of repose is an empirical quantity.
- Test with well-sorted material indicate that this angle is n ear 30 degree for sand, gradually increasing to 40 for gra vel.
- Poorly-sorted angular material tends to interlock, giving g reater resistance to failure, and as a result, a higher fricti on angle of repose.



- When a granular bed is subjected to a turbulent flow, it is found that virtually no motion of the grains is observed at some flows, but that the bed is noticeably mobilized at ot her flows.
- Factor that affect the mobility of grains subjected to a flo w are
  - Randomness: including grain placement, and turbulent
  - Forces on grains: Gravity and fluid
    - Fluid : lift and drag (mean & turbulent)
- In the presence of turbulent flow,
  - random fluctuations typically prevent the clear definition of a critical or threshold condition for motion



- But, still we need some conditions for describing it.
- Ikeda-Coleman-Iwagaki model for threshold conditions.
- Assumptions:
  - The forces acting on a particle taken to "dangerously place d.
  - The flow is taken to follow the turbulent rough logarithmic I aw near the boundary.
  - Turbulent forces on the particle are neglected.
  - Drag and lift forces act through the particle center
  - On drag force, boundary effects can be neglected
  - Very low streamwise slopes
  - Roughness height is equated to diameter of sediment



- Particle is centered at z=D/2
- To solve this problem, it is ne cessary to use some informati on about turbulent boundary I ayers to to define effective flui d velocity u<sub>f</sub> acting on the part icle.



In the viscous sublayer, a linear velocity profile

 $D/\delta_v > 2$  then no viscous sublayer exist.



 If no viscous sublayer exists, apply the rough pipe log v elocity profile

$$\frac{u_f}{u_*} = 2.5 \ln \left( 30 \frac{z}{D} \right) \Big|_{z=\frac{1}{2}D} = 6.77$$

Then the general form for the both cases,

• Where 
$$\frac{u_f}{u_*} = F\left(\frac{u_*D}{v}\right)$$

$$F = \frac{1}{2} \frac{u_*D}{v} \quad for \frac{u_*D}{v} < 13.5$$
 Since rough flow  

$$F = 6.77 \quad for \frac{u_*D}{v} > 13.5$$



 However, it would be more convenient to have a contin uous function

$$F = \left\{ \left(\frac{2\nu}{u_*D}\right)^{10/3} + \left[\kappa^{-1}\ln\left(\frac{\frac{9}{2}\frac{u_*D}{\nu}}{1+0.3\frac{u_*D}{\nu}}\right)\right]^{-10/3} \right\}^{-0.3}$$

- Forces acting on the particle  $D_f = \rho \frac{1}{2} \pi \left(\frac{D}{2}\right)^2 c_D u_f^2, \quad L_f = \rho \frac{1}{2} \pi \left(\frac{D}{2}\right)^2 c_L u_f^2, \quad F_g = \rho Rg \frac{4}{3} \pi \left(\frac{D}{2}\right)^3$
- Coulomb resistive force

$$F_r = \mu \left( F_g - L_f \right)$$



 For a channel with a downstream slope angle, the dow nslope effect of gravity has to be included in the force b alance presented earlier, resulting in

$$\tau_{c}^{*} = \frac{4}{3} \frac{\mu}{C_{D} + \mu C_{L}} \frac{1}{F^{2}(u_{*c}D/\nu)}$$

This is dimensionless critical shear stress and called as "the Shields particle  $\tau_{bc}^* = \tau_{bc}$ 

$$\tau_c^* = \frac{\tau_{bc}}{\rho g R D}$$



 The critical condition for incipient motion of the particle i s that the impelling drag force is just balanced by the re sisting Coulomb frictional force:

$$D_f = F_r$$

• With all forces above,

$$\frac{u_f^2}{RgD} = \frac{4}{3} \frac{\mu}{c_D + \mu c_L}$$

- Since  $\frac{u_f}{u_*} = F\left(\frac{u_*D}{v}\right)$  and  $\tau_b = \rho u_*^2$
- Ikeda-Coleman-Iwagaki

$$\tau_{c}^{*} = \frac{4}{3} \frac{\mu}{(c_{D} + \mu c_{L})} \frac{1}{F^{2}(u_{*c}D/v)}$$





 Comparison of Ikeda-Coleman-Iwagaki model for initiatio n of motion with Shields data





## 3. Shields Diagram

- Shields elucidate the condition for which sediment grins would be at the verge of moving.
- He used the fundamental concepts of similarity and dime nsional analysis.
- He said shield number should be function of shear Reyn olds number  $\int_{a_{*}}^{a_{*}} c_{vac}(u_{*c}D)$

$$\tau_c^* \sim function\left(\frac{u_{*c}D}{v}\right)$$

Key parameters for the system are

$$\tau_{bc}$$
,  $\gamma$  and  $\gamma_s$ ,  $D$ ,  $u_{*c} = \sqrt{\tau_{bc} / \rho}$ ,  $v$ 

• So simply  $\tau_c^* = \frac{\tau_{bc}}{\rho g R D} = F_* \left( \frac{u_{*c} D}{v} \right)$ 



3. Shields Diagram

$$\tau_c^* = \frac{\tau_{bc}}{\rho g R D} = F_* \left( \frac{u_{*c} D}{v} \right)$$

- But this form is not practical,
- Therefore, Vanoni (1964) proposed

$$\frac{u_*D}{v} = \frac{u_*}{\sqrt{gRD}} \frac{\sqrt{gRD}D}{v} = \left(\tau^*\right)^{1/2} \mathbf{R}_{ep}$$

Which was used in falling velocity relations





## 3. Shields Diagram





# 4. Granular Sediment on a Sloping Bed

- The work of Shields on initiation of motion applied only to the case of nearly horizontal slopes.
- Most streams, particularly in mountain areas, have steep gradients, creasing a need to account for the effect of th e downslope component of gravity on the initiation motio n.
- Wiberg and Smith (1987)

$$\tau_c^* = \frac{4}{3} \frac{\left(\tan\phi_0 \cos\alpha - \sin\alpha\right)}{\left(C_D + \tan\phi_0 C_L\right)} \frac{1}{F^2(z/z_0)}$$
  
$$\alpha = \text{Bed slope angle}$$
  
$$\phi_0 = \text{Angle of repose of the grains}$$



# 4. Granular Sediment on a Sloping Bed

- The effect of the streamwise bed slope on incipient sedi ment motion can be illustrated by considering the forces active on a particle lying in a bed consisting of similar par ticles over which water flows.
- Chiew and Parker (1994)

$$\frac{\tau_{c\alpha}^*}{\tau_{co}^*} = \cos\alpha \left(1 - \frac{\tan\alpha}{\tan\phi}\right)$$

- $\phi$ : Repose angle
- $\tau_{c\alpha}^*$ : critical shear stress for sediment on a bed with a longitudinal slope angle  $\alpha$
- $au_{co}^*$ : critical shear stress for a bed with very small slope



## 4. Granular Sediment on a Sloping Bed

In terms of shear velocities,

$$\frac{u_{*_c}}{u_{*_o}} = \sqrt{\cos\alpha \left(1 - \frac{\tan\alpha}{\tan\phi}\right)}$$

 Right figure was done by Chiew and Parker in duct laboratory experiment





## 5. Threshold Condition on Side Slopes

- Until now, we just talk the horizontal in the transverse dir ection.
- In engineering application, side wall slope must be impor tant.
- Simplification of the analysis:
  - Logarithmic upward normal from the bed.

$$D_{f} = \rho \frac{1}{2} \pi \left(\frac{D}{2}\right)^{2} c_{D} u_{f}^{2} \vec{e}_{1} \quad \text{(Drag force)}$$

$$F_{g} = F_{g2} \vec{e}_{2} + F_{g3} \vec{e}_{3} \qquad \text{(Gravitational force)}$$

$$\left(F_{g2}, F_{g3}\right) = -\rho Rg \frac{4}{3} \pi \left(\frac{D}{2}\right)^{3} (\sin\theta, \cos\theta)$$



## 5. Threshold Condition on Side Slopes

The lift force is given as

$$L_f = \rho \frac{1}{2} \pi \left(\frac{D}{2}\right)^2 c_L u_f^2 \vec{e}_3$$

 Coulomb resistive force must balance the sum of the imp elling forces du to flow and transverse downslope pull of gravity at the critical conditions

$$\mu^{2} \left| F_{g3} \vec{e}_{3} + L_{f} \right|^{2} = \left| D_{f} \right|^{2} + \left| F_{g2} \vec{e}_{3} \right|^{2}$$

With the previous terms

$$\left[\left(\frac{u_f^2}{RgD}\right)^2 + \left(\frac{4}{3c_D}\sin\theta\right)^2\right]^{1/2} = \mu\left(\frac{4}{3c_D}\cos\theta - \frac{c_L}{c_D}\frac{u_f^2}{RgD}\right)$$



## 5. Threshold Condition on Side Slopes

With some manipulations

$$\left[\left(\tau_c^*\right)^2 + \left(\frac{4}{3c_D F^2}\sin\theta\right)^2\right]^{1/2} = \frac{4\mu}{3c_D F^2}\cos\theta - \mu\frac{c_L}{c_D}\tau_c^*$$

 The case of a transversely horizontal bed is recovering b y setting θ=0. The critical Shields stress if found as in th e previous

$$\pi_c^* = \frac{4}{3} \frac{\left(\mu \cos\alpha - \sin\alpha\right)}{\left(c_D + \mu c_L\right)} \frac{1}{F^2\left(u_{*c}D/\nu\right)}$$

And denotes with the subscript of o.



- 5. Threshold Condition on Side Slopes
  - Then

$$\left[\left(\frac{\tau_c^*}{\tau_{co}^*}\right)^2 + \left(\frac{\left(1 + \mu c_L / c_D\right)}{\mu} \sin\theta\right)^2\right]^{1/2} = \left(1 + \mu \frac{c_L}{c_D}\right) \cos\theta - \mu \frac{c_L}{c_D} \left(\frac{\tau_c^*}{\tau_{co}^*}\right)$$

When  $\mu = 0.84 \ (\phi = 40^{\circ}), c_L = 0.85 c_D$ .

 $(c_D \text{ can be found in diagram})$ 

