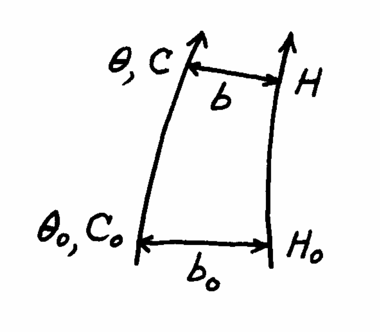
Chap 3. Transformation and Deformation of Random Sea Waves

3.1 Waves Refraction (+Shoaling)

3.1.1 Introduction

Ray theory for regular waves



Conservation of energy:

$$\frac{1}{8}\rho g H^2 C_g b = \frac{1}{8}\rho g H_0^2 C_{g0} b_0$$

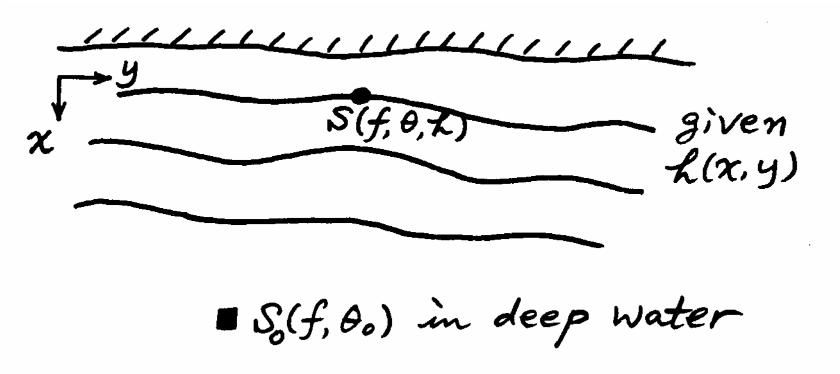
which gives

$$H = H_0 K_s K_r$$

shoaling coefficient, $K_s = \sqrt{\frac{C_{g0}}{C_g}} = \left[\tanh kh\left(1 + \frac{2kh}{\sinh 2kh}\right)\right]^{-1/2} = K_s(f,h)$

refraction coefficient, $K_r = \sqrt{\frac{b_0}{b}} = K_r(f,\theta,h)$; $\theta = \theta(f,h,\theta_0)$

3.1.2 Refraction Coefficient of Random Sea Waves



$$S(f,\theta) = \left[K_s(f,h)K_r(f,h,\theta)\right]^2 S_0(f,\theta_0)\frac{\partial\theta_0}{\partial\theta}$$

$$S(f) = \int_{-\pi}^{\pi} S(f,\theta)d\theta = \left[K_s(f,h)\right]^2 \int_{-\pi}^{\pi} S_0(f,\theta_0)\left[K_r(f,h,\theta_0)\right]^2 d\theta_0$$

$$m_0 = \int_0^\infty S(f)df = \int_0^\infty \int_{-\pi}^{\pi} S_0(f,\theta_0)\left[K_s(f,h)K_r(f,h,\theta_0)\right]^2 d\theta_0 df$$

$$= \int_0^\infty \int_{\theta_{\min}}^{\theta_{\max}} S_0(f,\theta_0)\left[K_s(f,h)K_r(f,h,\theta_0)\right]^2 d\theta_0 df$$

Goda's book uses

$$m_{s0} = \int_0^\infty \int_{\theta_{\min}}^{\theta_{\max}} S_0(f,\theta_0)\left[K_s(f,h)\right]^2 d\theta_0 df$$

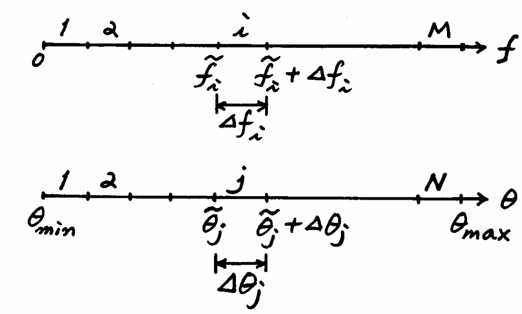
Define $(K_r)_{eff} = \left(\dfrac{m_0}{m_{s0}}\right)^{1/2}$

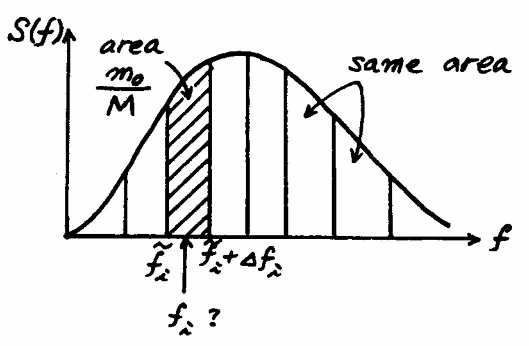
For example, $H_{m0} = 4\sqrt{m_0}$ = significant wave height after shoaling and refraction

$H_{ms0} = 4\sqrt{m_{s0}}$ = significant wave height due to shoaling only

then, $H_{m0} = (K_r)_{eff} H_{ms0}$

Goda's book explains how to discretize f and θ.





$$m_0 = \int_0^\infty S(f)df = \text{total area}$$

$$\int_{\tilde{f}_i}^{\tilde{f}_i+\Delta f_i} S(f)df = \frac{m_0}{M} \quad (i = 1,2,\cdots,M)$$

$S(f)$ can be integrated analytically (e.g. B-M or P-M spectra), say

$$S(f) = af^{-5}\exp(-bf^{-4})$$

$$m_0 = \int_0^\infty S(f)df = \left[\frac{a}{4b}\exp(-bf^{-4})\right]_0^\infty = \frac{a}{4b}$$

Similarly,

$$\int_{\tilde{f}_i}^{\tilde{f}_i+\Delta f_i} S(f)df = \left[\frac{a}{4b}\exp(-bf^{-4})\right]_{\tilde{f}_i}^{\tilde{f}_i+\Delta f_i} = \frac{a}{4b}\left\{\exp[-b(\tilde{f}_i+\Delta f_i)^{-4}] - \exp(-b\tilde{f}_i^{-4})\right\} = \frac{m_0}{M}$$

Hence,

$$\exp[-b(\tilde{f}_i+\Delta f_i)^{-4}] - \exp(-b\tilde{f}_i^{-4}) = \frac{1}{M} \quad (i = 1,2,\cdots,M)$$

Now we can find $\Delta f_i \ (i = 1,2,\cdots,M)$ starting from $\tilde{f}_1 = 0$.

Representative frequency f_i for the band ($\tilde{f}_i$ to $\tilde{f}_i + \Delta f_i$)?

Goda suggests on the basis of $\overline{T} = \sqrt{m_0 / m_2}$ and $m_2 = \int_0^\infty f^2 S(f) df$

$$f_i \cong \sqrt{\frac{(m_2)_i}{(m_0)_i}}$$

$$(m_0)_i = \frac{m_0}{M} = \frac{a}{4b}\frac{1}{M}$$

$$(m_2)_i = \int_{\tilde{f}_i}^{\tilde{f}_i + \Delta f_i} f^2 S(f) df = \int_{\tilde{f}_i}^{\tilde{f}_i + \Delta f_i} a f^{-3} \exp(-b f^{-4}) df$$

Putting $\sqrt{b} f^{-2} = \frac{\xi}{\sqrt{2}} \to f^{-3} df = \frac{d\xi}{-2\sqrt{2b}}$,

$$(m_2)_i = -\frac{a}{2\sqrt{2b}} \int_{\sqrt{2b}(\tilde{f}_i)^{-2}}^{\sqrt{2b}(\tilde{f}_i + \Delta f_i)^{-2}} \exp(-\xi^2/2) d\xi = \frac{a}{2\sqrt{2b}} \int_{\sqrt{2b}(\tilde{f}_i + \Delta f_i)^{-2}}^{\sqrt{2b}(\tilde{f}_i)^{-2}} \exp(-\xi^2/2) d\xi$$

Goda defines error function, $\Phi(t) = 1/\sqrt{2\pi}\int_0^t \exp(-x^2/2)dx$, though usual definition is $erf(t) = 2/\sqrt{\pi}\int_0^t \exp(-x^2)dx$ so that $erf(\infty) = 1$. Thus,

$$(m_2)_i = \frac{a}{2}\sqrt{\frac{\pi}{b}}\left\{\Phi\left[\sqrt{2b}\left(\tilde{f}_i\right)^{-2}\right] - \Phi\left[\sqrt{2b}\left(\tilde{f}_i + \Delta f_i\right)^{-2}\right]\right\}$$

$$f_i = \sqrt{\frac{(m_2)_i}{(m_0)_i}} = \left(2\sqrt{b\pi}M\right)^{1/2}\left\{\Phi\left[\sqrt{2b}\left(\tilde{f}_i\right)^{-2}\right] - \Phi\left[\sqrt{2b}\left(\tilde{f}_i + \Delta f_i\right)^{-2}\right]\right\}^{1/2}$$

which should correspond to Eq. (3.7) in Goda's book if $b = 1.03 T_s^{-4}$ (B-M spectrum):

$$f_i = \frac{1}{0.9 T_s}(2.912 M)^{1/2}\left\{\Phi\left(\sqrt{2\ln\frac{M}{i-1}}\right) - \Phi\left(\sqrt{2\ln\frac{M}{i}}\right)\right\}^{1/2}$$

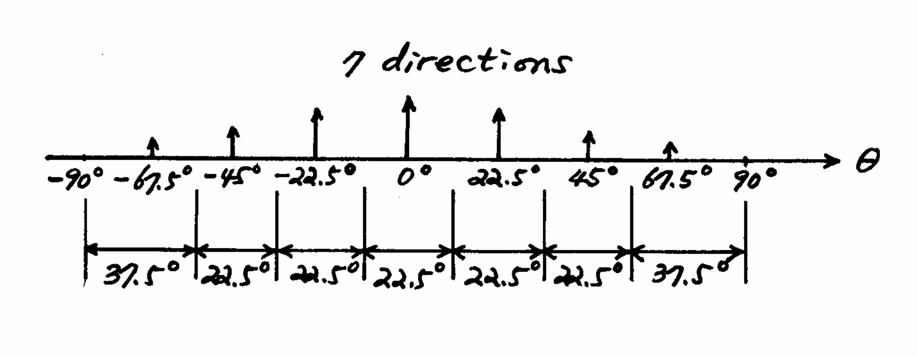
It is required

$$\sqrt{2\ln\frac{M}{i-1}} = \sqrt{2b}\left(\tilde{f}_i\right)^{-2} \quad (i = 1,2,\cdots,M)$$

On the other hand, $\exp\left[-b\left(\tilde{f}_i + \Delta f_i\right)^{-4}\right] - \exp\left[-b\left(\tilde{f}_i\right)^{-4}\right] = \dfrac{1}{M}$

Take $b\tilde{f}_i^{-4} = \ln\dfrac{M}{i-1}$. Then $\left(\dfrac{i}{M}\right) - \left(\dfrac{i-1}{M}\right) = \dfrac{1}{M}$ satisfied.

As for the discretization of wave angle θ (16-point bearing, see Table 3.2),



3.1.3 Computation of Random Wave Refraction by Means of the Energy Flux Equation

General transport equation for S (any scalar quantity):

$$\frac{\partial S}{\partial t} + \nabla \cdot (S\vec{V}) = Q \quad \text{(sink or source of } S\text{)}$$

where $\vec{V}$ = transport velocity of S. For $S(t, x, y, f, \theta)$ = directional random waves,

$\vec{V}$ = velocity following waves

$$\vec{V} = \left(\frac{dx}{dt}, \frac{dy}{dt}, \frac{d\theta}{dt}, \frac{df}{dt}\right) = \left(v_x, v_y, v_\theta, v_f\right)$$

with

$$v_x = C_g \cos\theta, \quad v_y = C_g \sin\theta$$

$$v_\theta = \frac{C_g}{C}\left(\frac{\partial C}{\partial x}\sin\theta - \frac{\partial C}{\partial y}\cos\theta\right) \text{ to account for refraction}$$

$v_f = 0$ assuming f does not change following the wave.

Then

$$\frac{\partial S(f,\theta)}{\partial t} + \frac{\partial}{\partial x}[S(f,\theta)v_x] + \frac{\partial}{\partial y}\left[S(f,\theta)v_y\right] + \frac{\partial}{\partial \theta}[S(f,\theta)v_\theta] = Q$$

For steady state ($\partial S/\partial t = 0$) with no sink or source ($Q = 0$),

$$\frac{\partial}{\partial x}(Sv_x) + \frac{\partial}{\partial y}\left(Sv_y\right) + \frac{\partial}{\partial \theta}(Sv_\theta) = 0 \quad \text{for} \quad S(x, y, f, \theta)$$

Assuming $\theta \neq \theta(x, y)$, or x, y, θ are independent variables,

$$\cos\theta \frac{\partial\left(SCC_g\right)}{\partial x} + \sin\theta \frac{\partial\left(SCC_g\right)}{\partial y} + C_g \frac{\partial S}{\partial \theta}\left(\sin\theta \frac{\partial C}{\partial x} - \cos\theta \frac{\partial C}{\partial y}\right) = 0$$

where C and C_g using linear wave theory depend on $h(x, y)$ and frequency f. θ is computed by ray theory. We need boundary conditions for S.

Example in Goda's book Fig. 3.3: Waves over a circular shoal

- T_s as well as H_s changes depending on locations (Fig. 3.4).
- Fig. 3.5 for regular waves shows larger spatial variations of wave heights.

Ref. Vincent and Briggs (1989). Refraction-diffraction of irregular waves over a mound, JWPCOE, 115(2), 269-284: Performed laboratory experiments on transformation of monochromatic and random directional waves over an elliptic shoal. They concluded that monochromatic waves using representative wave height and period (e.g. H_s and T_s) provide a poor approximation of irregular wave conditions if there is directional spread or high wave steepness.

Ref. 권혁민 (1998). 방향 스펙트럼 파랑에 대한 3 차원 쇄파변형 모델, 대한토목학회논문집, 18(II-6), 591-599: Include sink term due to wave breaking.

Ref. Mase, H. (2001). Multi-directional random wave transformation model based on energy balance equation, Coastal Engineering Journal, 43(4), 317-337: Include wave diffraction.

3.1.4 Wave Refraction on a Coast with Straight, Parallel Depth Contours

Snell's law: $\dfrac{\sin\theta}{C} = \dfrac{\sin\theta_0}{C_0}$ $\rightarrow$ can find $\theta(f,h,\theta_0)$ $\rightarrow$ $\dfrac{\partial\theta}{\partial\theta_0}$

Refraction coefficient $K_r(f,h,\theta_0) = \sqrt{\dfrac{\cos\theta_0}{\cos\theta}}$

Shoaling coefficient $K_s = \sqrt{\dfrac{C_{g0}}{C_g}} = K_s(f,h)$

Directional spectrum $S(f,\theta)$ in water depth h:

$$S(f,\theta) = \left[K_s(f,h)K_r(f,h,\theta_0)\right]^2 \left(\frac{\partial\theta}{\partial\theta_0}\right)^{-1} S_0(f,\theta_0)$$

Need to specify $S_0(f,\theta_0)$ in deep water. For example, $S_0(f,\theta_0) = S_0(f)G(f,\theta_0)$ with $S_0(f)$ = B-M spectrum with given $H_s = H_{m0}$ and $T_s = T_p/1.05$,

$G(f,\theta_0)$ = Mitsuyasu-type with given $s_{\max}$ and $(\alpha_p)_0$,

$(\alpha_p)_0$ = predominant wave direction in deep water,

$$G(f,\theta_0) = G_0 \cos^{2s}\left[\frac{\theta_0 - (\alpha_p)_0}{2}\right]; \quad -\pi \le \left[\theta_0 - (\alpha_p)_0\right] \le \pi\,,$$

$G = G_0$ is maximum at $\theta_0 = (\alpha_p)_0$.

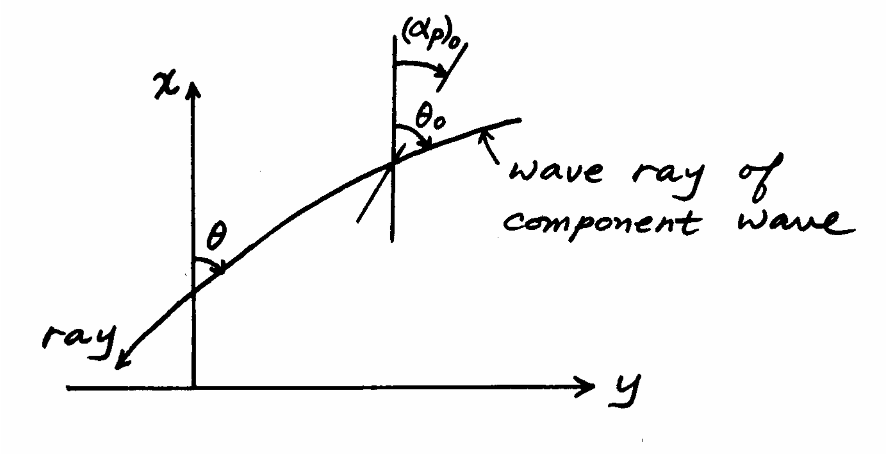


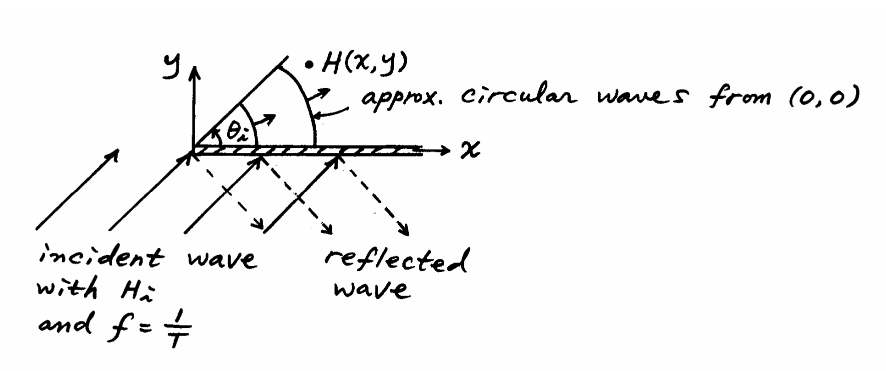
Fig. 3.6 shows $(K_r)_{eff}$ given by Eq. (3.2) as a function of h/L_0 with $L_0 = gT_s^2/2\pi$, $(\alpha_p)_0$ and $s_{\max}$.

Note: $(K_r)_{eff} \neq 1$ even for $(\alpha_p)_0 = 0$ ($\because$ directional spreading).

3.2 Wave Diffraction

3.2.1 Principle of Random Wave Diffraction Analysis

For linear monochromatic waves in constant water depth, Sommerfeld solution for a semi-infinite thin breakwater:



Diffraction coefficient $K_d = H(x,y)/H_i$ depends on f, θ_i, and h=constant.

$$
\begin{array}{lll}
H(x,y) & = & K_d(f,\theta_i;x,y,h)H_i \\
\downarrow & & \qquad\qquad\qquad\quad\downarrow \\
S(f) & & \qquad\qquad\qquad\quad S_i(f,\theta_i)
\end{array}
$$

Frequency spectrum

$$S(f) \text{ at given } (x,y) = \int_{-\pi}^{\pi} [K_d(f,\theta_i)]^2 S_i(f,\theta_i)d\theta_i$$

Since

$$\int_0^\infty S(f)df = \int_0^\infty \int_{-\pi}^{\pi} S(f,\theta)d\theta df = \int_0^\infty \int_{-\pi}^{\pi} [K_d(f,\theta_i)]^2 S_i(f,\theta_i)d\theta_i df$$

therefore

$$S(f,\theta) = [K_d(f,\theta_i)]^2 S_i(f,\theta_i)\frac{\partial\theta_i}{\partial\theta}$$

Then

$$S(f) = \int_{-\pi}^{\pi} S(f,\theta)d\theta = \int_{-\pi}^{\pi} [K_d(f,\theta_i)]^2 S_i(f,\theta_i)d\theta_i$$

In terms of zeroth moment,

$$(m_0)_i = \int_0^\infty \int_{-\pi}^{\pi} S_i(f,\theta_i)d\theta_i df \rightarrow (H_{m0})_i = 4\sqrt{(m_0)_i}$$: incident significant wave height

$$m_0 = \int_0^\infty \int_{-\pi}^{\pi} S(f,\theta)d\theta df \rightarrow H_{m0} = 4\sqrt{m_0}$$: significant wave height at (x,y)

$$m_0 = \int_0^\infty \int_{-\pi}^{\pi} [K_d(f,\theta_i)]^2 S_i(f,\theta_i)d\theta_i df$$

Define effective diffraction coefficient:

$$(K_d)_{eff} = \frac{H_{m0}}{(H_{m0})_i} = \left[\frac{m_0}{(m_0)_i}\right]^{1/2} = (3.14)\text{ in Goda's book where } i \text{ added}$$

$$= \left[\frac{1}{(m_0)_i}\int_0^\infty \int_{-\pi}^{\pi} S_i(f,\theta_i)[K_d(f,\theta_i)]^2 d\theta_i df\right]^{1/2}$$

Read Goda's book for field measurement (Figs. 3.9 and 3.10). $(K_d)_{eff} > K_d$ based on regular waves with $H = H_{1/3}$ and $T = T_{1/3}$ = 0.07, which is significantly underestimated in this case.

3.2.2 Diffraction Diagrams of Random Sea Waves

$$S(f;x,y,h) = \int_{-\pi}^{\pi} [K_d(f,\theta_i;x,y,h)]^2 S_i(f,\theta_i)d\theta_i$$

$$m_0(x,y,h) = \int_0^\infty S(f)df\ ; \qquad (m_0)_i = \int_0^\infty S_i(f)df$$

$$m_2(x,y,h) = \int_0^\infty f^2 S(f)df\ ; \qquad (m_2)_i = \int_0^\infty f^2 S_i(f)df$$

peak $T_p(x,y,h)$ from $S(f)$; peak $(T_p)_i$ from $S_i(f)$

$$H_{m0} = 4\sqrt{m_0},\quad \overline{T} = \sqrt{m_0/m_2}\ ; \qquad (H_{m0})_i = 4\sqrt{(m_0)_i},\quad (\overline{T})_i = \sqrt{(m_0)_i/(m_2)_i}$$

$$H_s \cong H_{m0},\quad T_s \cong T_p/1.05\ ; \qquad (H_s)_i \cong (H_{m0})_i,\quad (T_s)_i \cong (T_p)_i/1.05$$

Wave height ratio = $(K_d)_{eff} = \dfrac{H_{m0}}{(H_{m0})_i} \cong \dfrac{H_s}{(H_s)_i}$ only for $H_{m0} = H_s$

Period ratio $\cong \dfrac{\overline{T}}{(\overline{T})_i}$ or $\dfrac{T_p}{(T_p)_i} \cong \dfrac{T_s}{(T_s)_i}$

It is not specified in Goda's book which relation is used for period ratio. It is likely to use $\overline{T}/(\overline{T})_i$ since T_p may be difficult to find. But Goda uses $(T_s)_i$ to find L (p.59).

Goda assumed $S_i(f,\theta_i) = S_i(f)G(f,\theta_i)$ with B-M frequency spectrum and Mitsuyasu-type directional spreading.

Need to specify $(H_s)_i = (H_{m0})_i$, $(T_s)_i = (T_p)_i / 1.05$, $s_{\max}$, $(\alpha_p)_i$, constant depth h, and breakwater geometry.

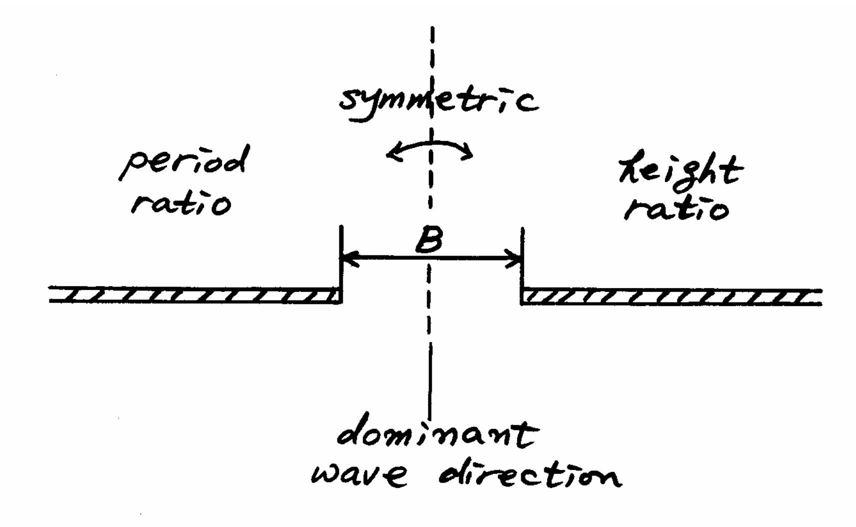
$$\text{Plotted} \begin{cases} \text{height ratio} = (K_d)_{eff} \\ \text{period ratio} = \dfrac{T_s}{(T_s)_i} \ \text{(probably)} \end{cases} \quad \text{for normal incidence only, } (\alpha_p)_i = 0^\circ .$$

Fig. 3.11 for a semi-infinite breakwater, for $s_{\max} = 10$ (wind waves) and $s_{\max} = 75$ (swell, more unidirectional).

Monochromatic versus directional random waves:

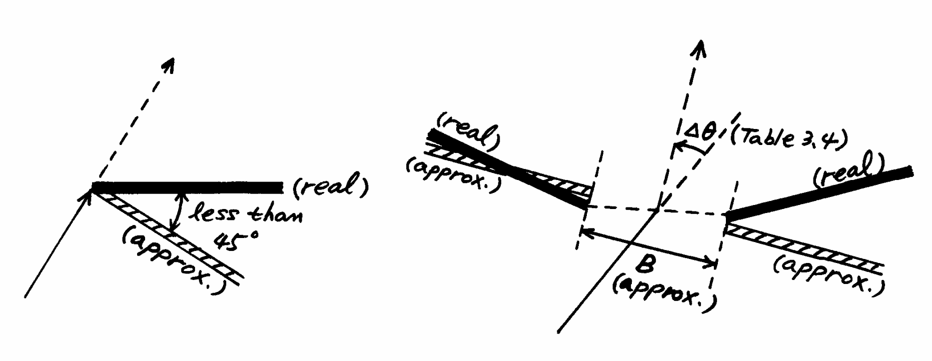
1) In general, monochromatic wave underestimates wave heights in sheltered area, and overestimates in open area.
2) The wave height ratio along the boundary of the geometric shadow (or the straight line from the tip of the breakwater to the wave direction) is 0.7 for directional random waves, while it is 0.5 for monochromatic waves.

Figs. 3.12~3.15 for breakwater gap ($B/L = 1,2,4,8$)



3.2.3 Random Wave Diffraction of Oblique Incidence

Construct your own computer program if exact solution is needed. Otherwise, use an approximate method suggested in the book.

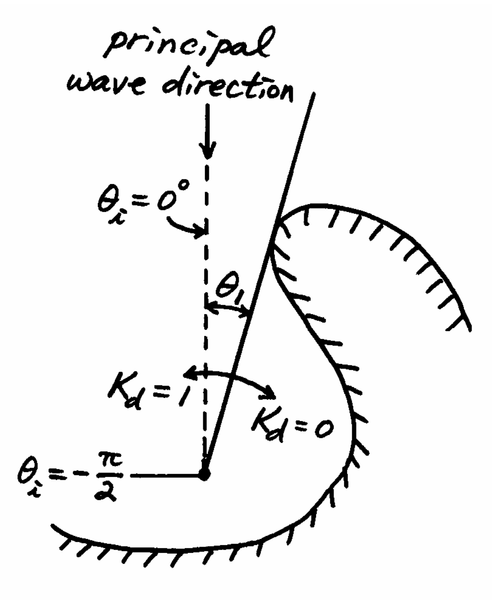


3.2.4 Approximate Estimation of Diffracted Height by the Angular Spreading Method

For large barriers (e.g. headlands and islands),

$$\text{roughly}\begin{cases} K_d \cong 0 \text{ in geometric shadow} \\ K_d \cong 1 \text{ in illuminated region} \end{cases}$$

Neglect wave refraction.



$$(m_0)_i = \int_0^\infty \int_{-\pi}^{\pi} S_i(f,\theta_i)\,d\theta_i df$$

Assume $S_i = 0$ for $|\theta_i| > \pi/2$.

$$(m_0)_i = \int_0^\infty \int_{-\pi/2}^{\pi/2} S_i(f,\theta_i)\,d\theta_i df$$

$$m_0 = \int_0^\infty \int_{-\pi/2}^{\pi/2} [K_d(f,\theta_i)]^2 S_i(f,\theta_i)\,d\theta_i df$$

$$\text{Assume}\begin{cases} K_d = 1 \quad \text{for } -\pi/2 \le \theta_i \le \theta_1 \\ K_d = 0 \quad \text{for } \theta_1 < \theta_i \le \pi/2 \end{cases}$$

Then

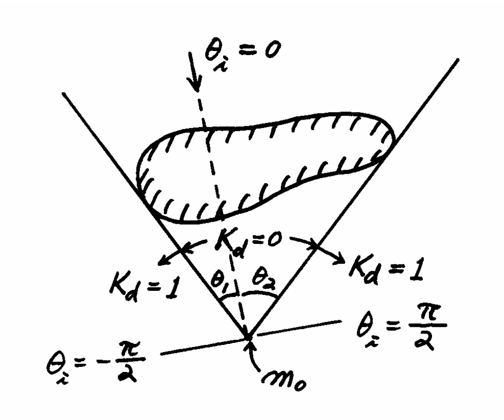
$$m_0 = \int_0^\infty \int_{-\pi/2}^{\theta_1} S_i(f,\theta_i)\,d\theta_i\,df$$

$$\frac{m_0}{(m_0)_i} = P_E(\theta_1) = \text{cumulative relative energy from } -\pi/2 \text{ to } \theta_1$$

$$= \text{Eq. (2.27) or Fig. 2.15 (B - M spectrum + Mitsuyasu spreading)}$$

$$(K_d)_{eff} = \left[\frac{m_0}{(m_0)_i}\right]^{1/2} = \left[P_E(\theta_1)\right]^{1/2}$$

$\theta_1 < 0$ and $\theta_2 > 0$ for this problem



$$m_0 = \int_0^\infty \int_{-\pi/2}^{\theta_1} S_i(f,\theta_i)\,d\theta_i\,df + \int_0^\infty \int_{\theta_2}^{\pi/2} S_i(f,\theta_i)\,d\theta_i\,df$$

$$\frac{m_0}{(m_0)_i} = P_E(\theta_1) + \left[P_E(\pi/2) - P_E(\theta_2)\right]$$

$$(K_d)_{eff}^2 = \frac{m_0}{(m_0)_i} = P_E(\theta_1) + \left[1 - P_E(\theta_2)\right]$$

$$= (K_d)_1^2 + (K_d)_2^2 \text{ in the text}$$

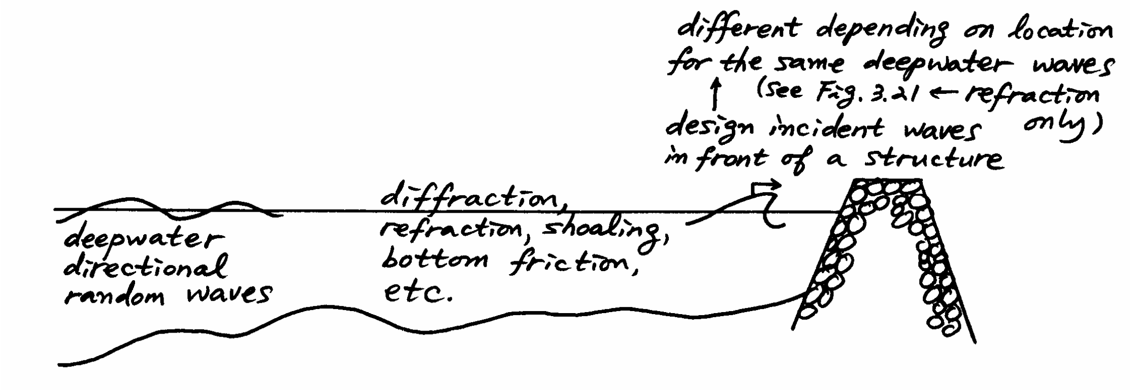
3.2.5 Applicability of Regular Wave Diffraction Diagrams

↑

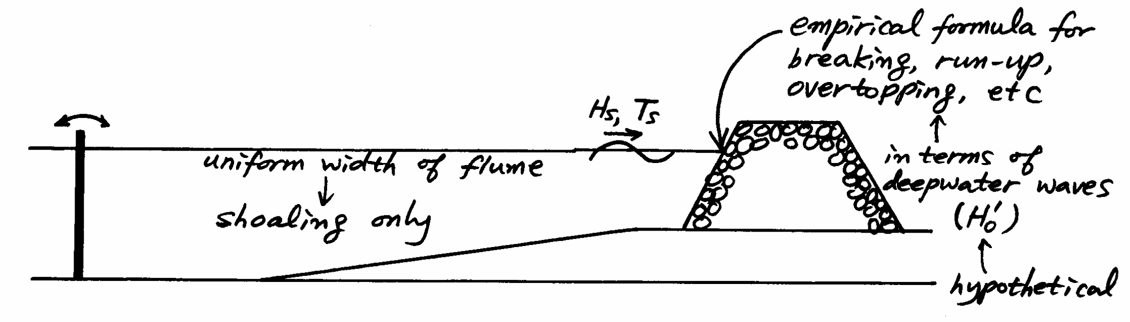
Only for very narrow directional spreading

3.3 Equivalent Deepwater Wave

In real situation,



Hydraulic model test in 2D wave flume,



In real situation, $H_s = K_d K_r K_s K_f (H_s)_0$

In 2D wave flume, $H_s = K_s H_0'$

Thus, $H_0' = K_d K_r K_f (H_s)_0$

↑

(unrefracted) equivalent deepwater wave height

For wave period, usually assumes $T_s = (T_s)_0$ ← error if diffraction is dominant.

3.4 Wave Shoaling

Linear wave shoaling coeff. $K_s = \frac{H}{H_0'} = \sqrt{\frac{C_{g0}}{C_g}} = \text{function of } \frac{h}{L}$ (3.16)

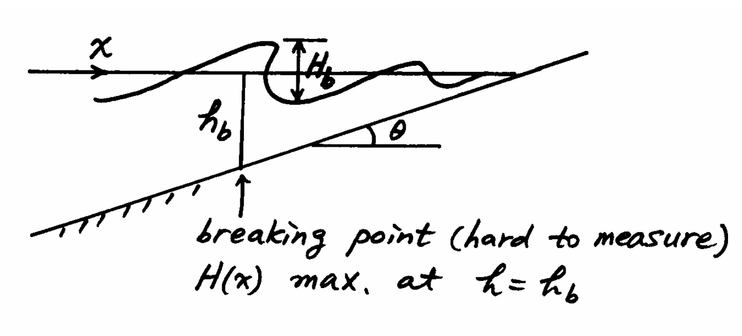
For shoaling of normally-incident linear random waves,

$$S(f;h) = [K_s(f;h)]^2 S_0(f) \leftarrow \text{can write a computer program easily.}$$

For shoaling of nonlinear monochromatic waves, use Shuto (1974) model (read text).

3.5 Wave Deformation Due to Random Breaking

3.5.1 Limiting Wave Height of Regular Waves by Breaking



$$\frac{H_b}{h_b} = f\left(\tan\theta, \frac{h_b}{L_0}\right) \rightarrow \text{Fig. 3.23}$$

Goda's empirical formula (1970):

$$\frac{H_b}{L_0} = A\left\{1 - \exp\left[-1.5\frac{\pi h_b}{L_0}\left(1 + 15\tan^{4/3}\theta\right)\right]\right\} \qquad (3.22)$$

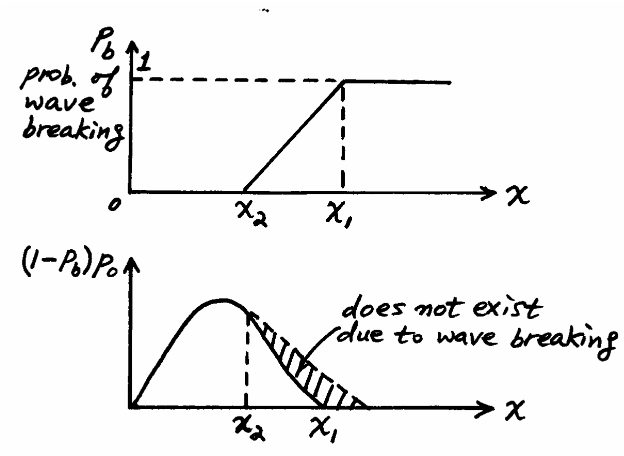
with $L_0 = gT^2/2\pi$; $A \cong 0.17$

3.5.2 Computational Model of Random Wave Breaking

Before wave breaking, Rayleigh distribution may be assumed

$$p_0(x) = \frac{\pi}{2} x \exp\left(-\frac{\pi}{4}x^2\right); \quad x = \frac{H}{\overline{\overline{H}}} \qquad (2.1)$$

After breaking, $p_0(x) \to p(x)$



$$1 - p_b = \text{probability of non-breaking}$$

$$1 - p_b = \begin{cases} 0 & \text{for } x \geq x_1 \\ \dfrac{x_1 - x}{x_1 - x_2} & \text{for } x_2 < x < x_1 \\ 1 & \text{for } x \leq x_2 \end{cases}$$

Let $A = \int_0^{x_1} (1 - p_b) p_0 dx < 1$ since $\int_0^{\infty} p_0 dx = 1$

Assume p.d.f. adjusted for wave breaking:

$$p(x) = \frac{1}{A}(1 - p_b)p_0(x) \quad \text{so that} \quad \int_0^{\infty} p(x)dx = 1$$

3.5.3 Computation of the Change in Wave Height Distribution Due to Random Wave Breaking

Need to estimate $\left\{\begin{matrix} x_1 = \text{upper limit} \\ x_2 = \text{lower limit} \end{matrix}\right\}$ of wave breaking.

Use $$\frac{H_b}{L_0} = A\left\{1 - \exp\left[-1.5\frac{\pi h}{L_0}\left(1 + 15\tan^{4/3}\theta\right)\right]\right\} \qquad (3.22)$$

with $L_0 = \dfrac{gT_s^2}{2\pi}$ and $\tan\theta$ = beach slope

$$A = \begin{cases} 0.18 \ \text{ for } x = x_1 \\ 0.12 \ \text{ for } x = x_2 \end{cases}$$

Eq. (3.22) was developed for breaking point ($h = h_b$) of regular waves. But it may be used inside the surf zone if H_b = broken wave height, h = local depth.

<u>Water level change</u>:

Wave setup ($\bar{\eta}$) was computed using the results of monochromatic waves with T_s and $\overline{H^2}$ = mean square of random waves, the latter of which is affected by $\bar{\eta}$. Therefore, we need iteration to solve $\bar{\eta}$ and $\overline{H^2}$ simultaneously. See Eq. (3.23).

Surf beat, $\zeta(t)$: slow (30~300 s) fluctuation of free surface mainly inside surf zone

↑

from Gaussian distribution with ζ_{rms} given by Eq. (3.24)

Thus, $h = d + \bar{\eta} + \zeta(t)$

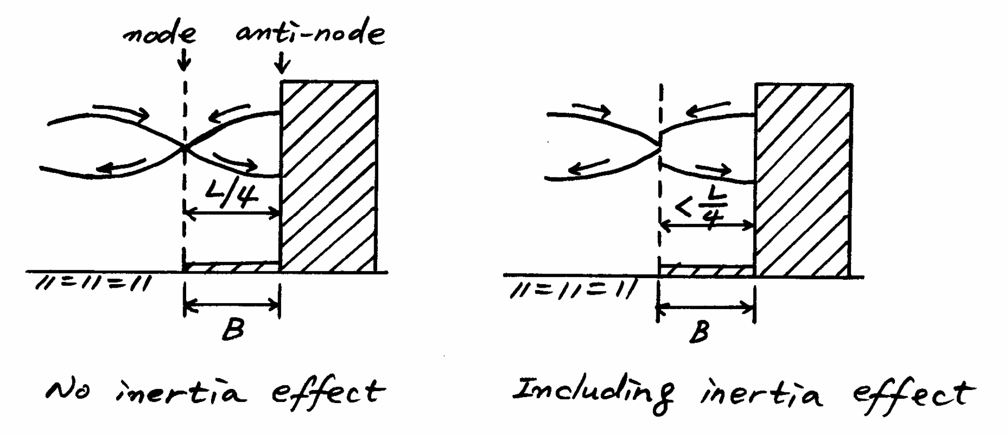
3.6 Wave Reflection and Dissipation

3.6.1 Coefficient of Wave Reflection

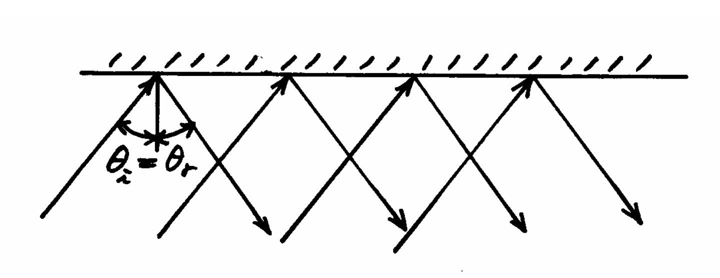
$$K_R = \frac{H_R}{H_i}$$

Typical reflection coefficients are given in Table 3.7.

For perforated wall caissons, K_R becomes minimum (0.3~0.4) at $B/L = 0.15 \sim 0.2$ (see Fig. 3.36). Under a standing wave system, maximum u at node $\rightarrow$ maximum energy dissipation & minimum reflection at $B = L/4$ $\rightarrow$ $B/L = 0.25$. However, in reality, minimum reflection occurs at $B/L = 0.15 \sim 0.2$, due to inertia effect.



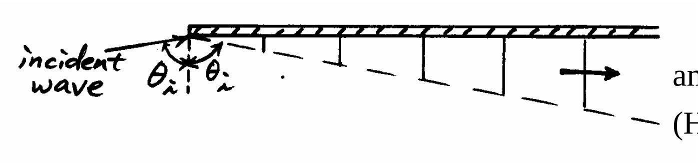
3.6.2 Propagation of Reflected Waves



$\theta_i = \theta_r$ (geometrical optics theory)

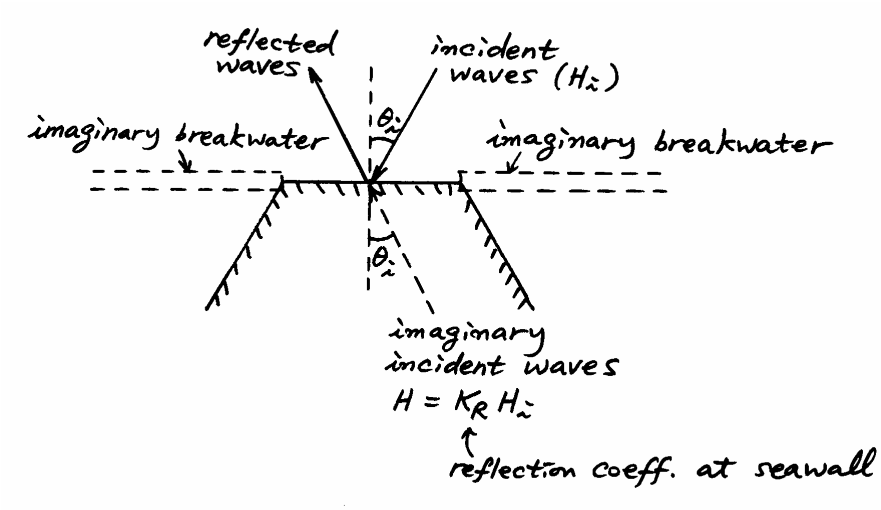
diamond pattern of surface profile

For long-period waves incident at large angle, Mach stem is formed.

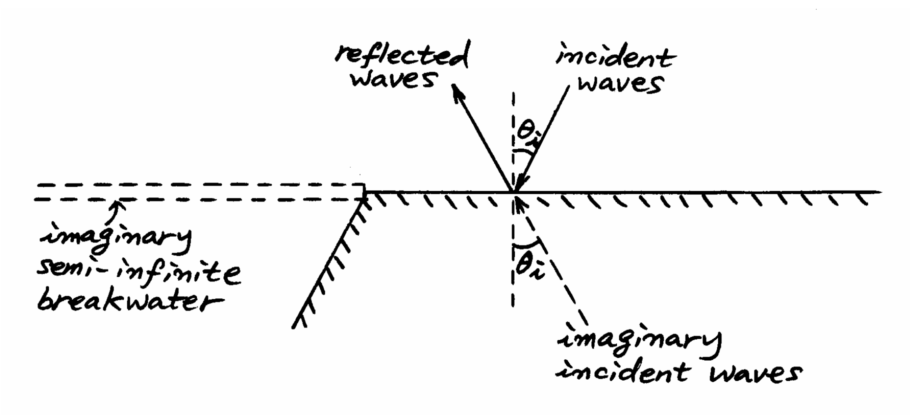


amplitude dispersion
(Higher waves go faster.)

Reflection from finite length of seawall ← diffraction by breakwater gap



Reflection from very long seawall ← diffraction by semi-infinite breakwater
(or angular spreading method for headland)



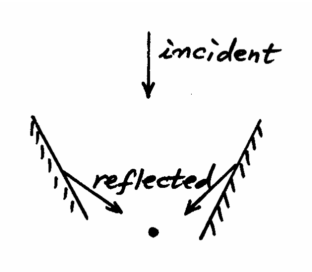
Effect of opposing wind (sea → land): attenuates waves of large steepness, but its effect is minor for swell of low steepness.

3.6.3 Superposition of Incident and Reflected Waves

For linear waves, we can superpose the free surface displacement:

$$\eta(t,x,y) = \eta_i(t,x,y) + \sum_{n=1}^{N} \eta_R^n(t,x,y)$$

total incident reflected waves



Time-averaged energy per unit surface area:

$$\rho g \overline{\eta^2} \text{ at given } (x,y) = \rho g m_0; \quad m_0 = \int_0^{\infty} S(f)df$$

If the distance from the reflective structure is more than one wavelength, we may assume

$$\overline{\eta_i \eta_R^n} = 0 \quad (n = 1,2,\cdots,N), \quad \overline{\eta_R^n \eta_R^m} = 0 \quad (n \neq m) \quad \text{uncorrelated.}$$

Then

$$\overline{\eta^2} = \overline{\eta_i^2} + \sum_{n=1}^{N} \overline{(\eta_R^n)^2}$$

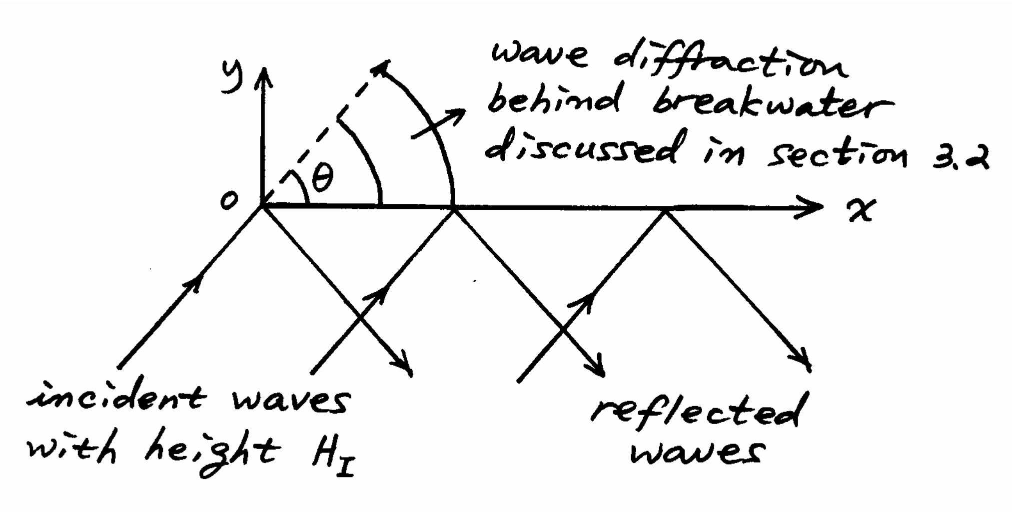
$$m_0 = (m_0)_i + \sum_{n=1}^{N} (m_0)_R^n \leftarrow \text{addition of area } m_0 \text{ under spectrum } S(f)$$

$$H_{m0}^2 = (H_{m0})_i^2 + \sum_{n=1}^{N} \left[(H_{m0})_R^n\right]^2 \qquad (3.28)$$

Fig. 3.40 indicates $H_{m0} \cong \sqrt{(H_{m0})_i^2 + (H_{m0})_R^2}$ at $x/L \geq 1.0$

3.7 Spatial Variation of Wave Height along Reflective Structures

3.7.1 Wave Height Variation near the Tip of a Semi-Infinite Structure



Wave height (crest elevation – trough elevation) along vertical wall ($y = 0$):

$$\frac{H_S}{H_I} = \sqrt{(C + S + 1)^2 + (C - S)^2}$$

where

$$C = \int_0^u \cos\left(\frac{\pi}{2} t^2\right) dt\,, \quad S = \int_0^u \sin\left(\frac{\pi}{2} t^2\right) dt\,, \quad u = 2\sqrt{\frac{2x}{L}} \sin\frac{\theta}{2}$$

Note: at $x = 0$, $u = 0 \;\rightarrow\; C = S = 0 \;\rightarrow\; H_S / H_I = 1$

As $x \to \infty$, $u \to \infty$, then $C \to \frac{1}{2}$, $S \to \frac{1}{2}$ $\quad \therefore\; \frac{H_S}{H_I} = 2$

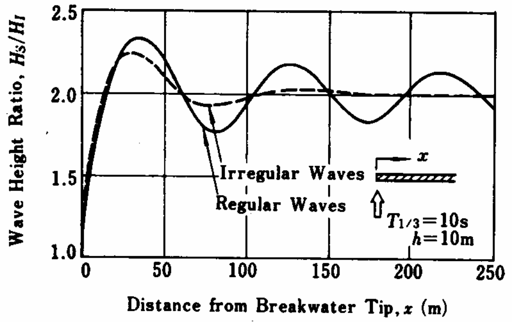


Fig. 3.42. Variation of wave height in front of a semi-infinite breakwater.

less undulation
for irregular waves

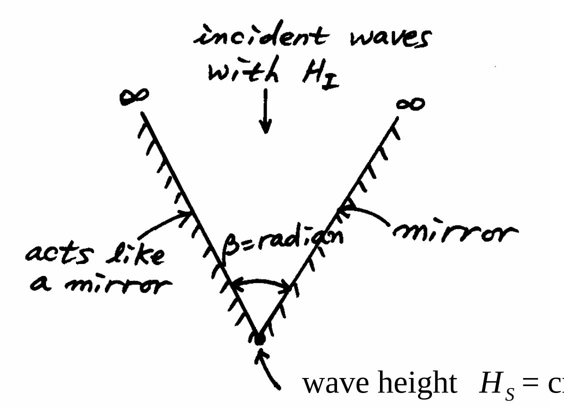
For irregular waves, $K_d = \dfrac{H_S}{H_I}$ was used in Eq. (3.14)

$\uparrow$

for component waves ($f \rightarrow L \rightarrow u$)

Explains meandering damage of concrete caissons.

3.7.2 Wave Height Variation at an Inward Corner of Reflective Structures

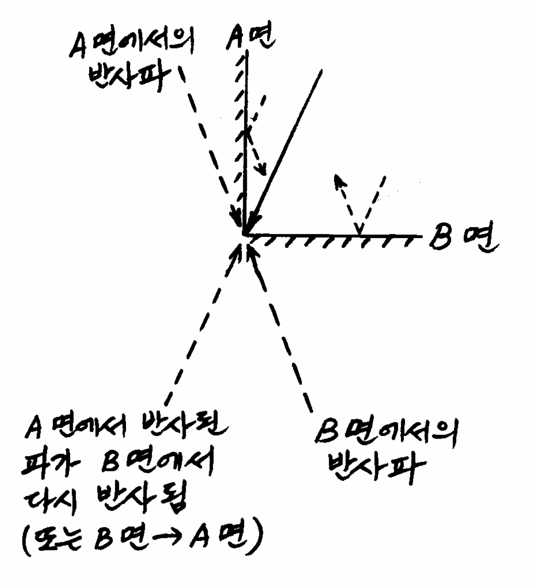


$$\frac{H_S}{H_I} = \frac{2\pi}{\beta} \qquad (3.31)$$

$$\text{for} \quad \beta = \pi, \quad \frac{H_S}{H_I} = 2$$

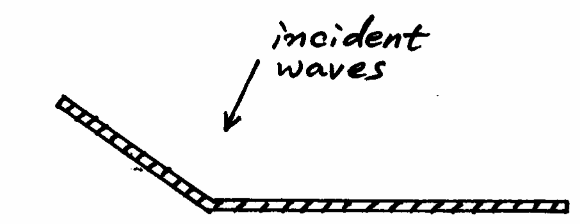
$$\beta = \frac{\pi}{2}, \quad \frac{H_S}{H_I} = 4$$

wave height H_S = crest elevation – trough elevation

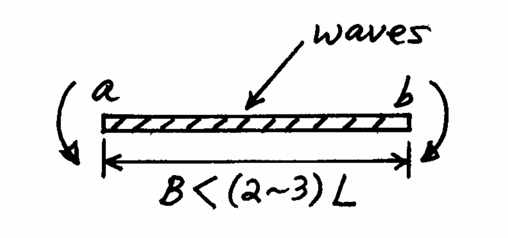


Same as sum of 4 waves propagating in 4 different directions

If the length is finite, use a computer program or an approximate method given in Goda’s book.

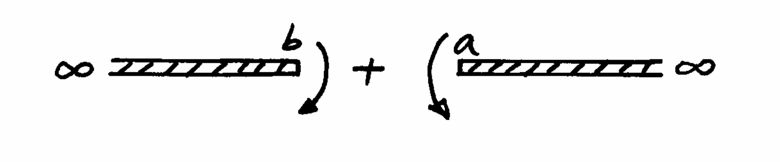


3.7.3 Wave Height Variation along an Island Breakwater



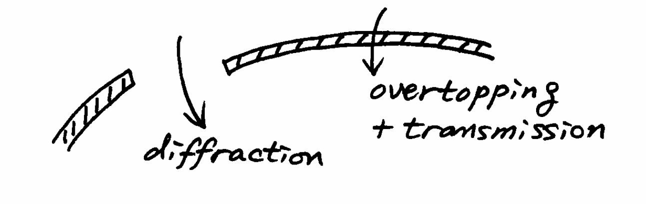
cause undulation along wall (Fig. 3.46 and 3.47)

If $B >> L$, may add two waves diffracted from each tip:

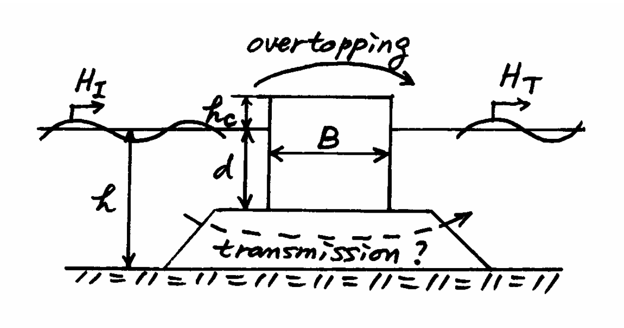


3.8 Wave Transmission over Breakwater

3.8.1 Wave Transmission Coefficient



(A) Vertical breakwater



transmission coefficient $K_T = \frac{H_T}{H_I}$

Wave transmission through rubble mound may be negligible.

Expect $K_T = \text{function}\left(\frac{h_c}{H_I}, \frac{d}{h}, B, T, \text{mound material}, \cdots\right)$

Fig. 3.48 for regular wave tests $\rightarrow$ may be applicable to irregular waves with $H_I = (H_{1/3})_I$ and $H_T = (H_{1/3})_T$ (see Fig. 3.49)

Eq. (3.33) $\rightarrow$ $K_T = \text{function}\left(\frac{h_c}{H_I} \text{ only}\right)$; Effect of $\frac{d}{h}$ is minor (see Fig. 3.48).

Eq. (3.34) $\rightarrow$ horizontally composite breakwaters

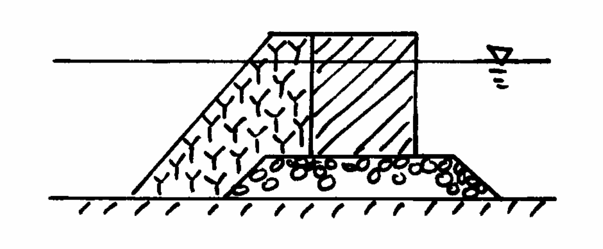
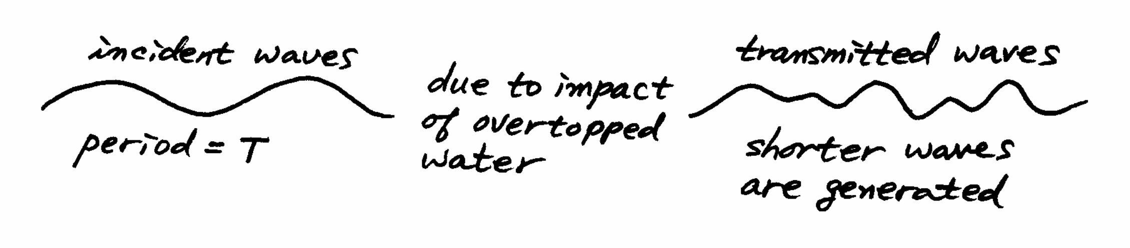


Fig. 3.50 suggests $\frac{(T_{1/3})_T}{(T_{1/3})_I}$ and $\frac{(\overline{T})_T}{(\overline{T})_I} \cong 0.5 \sim 0.8$ due to generation of higher harmonic waves.



(B) Breakwater consisting of energy-dissipating blocks (Tetrapod, Dolos,…)

↓

Fig. 3.51

(C) Rubble mound breakwater

Less void than concrete blocks → smaller K_T (0.1~0.3). But $K_T \uparrow$ as $T \uparrow$

3.8.2 Propagation of Transmitted Waves in a Harbor

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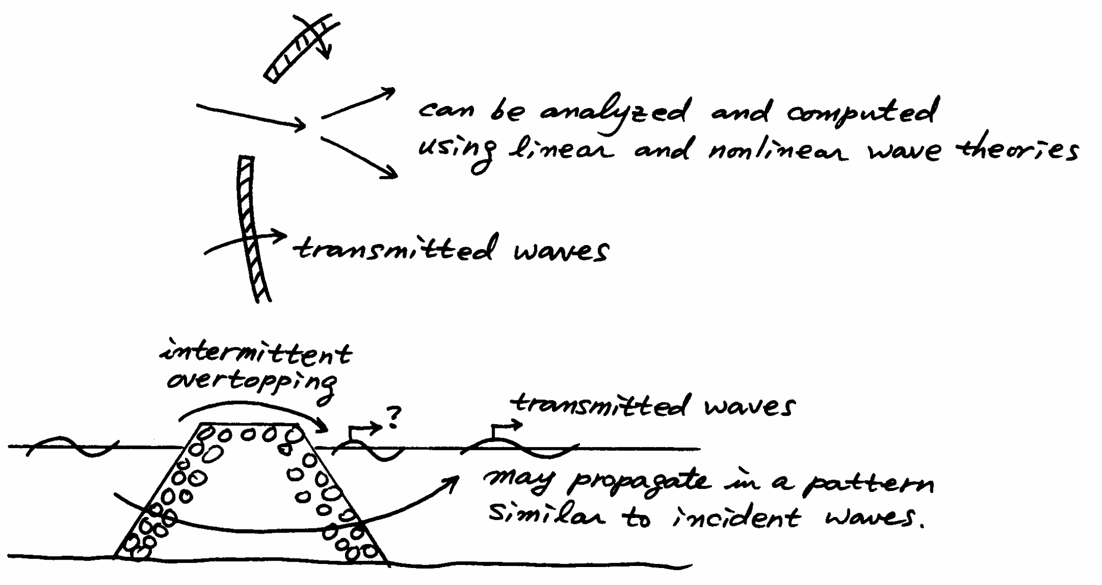

3.9 Longshore Currents by Random Waves on Planar Beach

3.9.1 Longshore Currents by Unidirectional Irregular Waves

Longshore current profile for regular waves

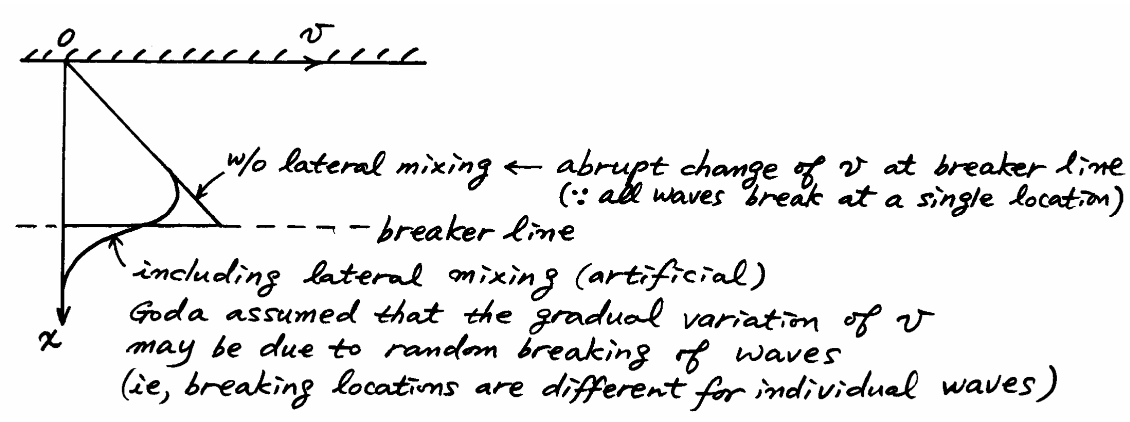

Fig. 3.52 gives cross-shore variation of longshore current velocity based on random breaking model discussed in Section 3.5. In terms of deepwater wave condition,

$$V = \text{dimensionless velocity} = \frac{C_f v}{\sqrt{gH_0}\tan\theta\sin\alpha_0}$$

$$z = \text{depth normalized w.r.t. } H_0 = \frac{h}{H_0}$$

$$v(z) = v_0\left(\frac{z-B}{A}\right)^{k-1}\exp\left[-\left(\frac{z-B}{A}\right)^{k}\right]$$

v_0, A, B, k are given in Eqs. (3.37) to (3.40)

To find $v_{\max}$,

$$\frac{dv}{dz} = v_0(k-1)\left(\frac{z-B}{A}\right)^{k-2}\exp\left[-\left(\frac{z-B}{A}\right)^{k}\right] - v_0\left(\frac{z-B}{A}\right)^{k-1}k\left(\frac{z-B}{A}\right)^{k-1}\exp\left[-\left(\frac{z-B}{A}\right)^{k}\right]$$

$$= v_0\left(\frac{z-B}{A}\right)^{k-2}\exp\left[-\left(\frac{z-B}{A}\right)^{k}\right]\left\{k-1-k\left(\frac{z-B}{A}\right)^{k}\right\} = 0$$

$$k\left(\frac{z-B}{A}\right)^k = k-1 \;\rightarrow\; \left(\frac{z-B}{A}\right)^k = 1-\frac{1}{k} \;\rightarrow\; \frac{z-B}{A} = \left(1-\frac{1}{k}\right)^{1/k}$$

$$\rightarrow\; z = B + A\left(1-\frac{1}{k}\right)^{1/k}$$

$$\therefore v_{\max} = v_0\left(1-\frac{1}{k}\right)^{1-\frac{1}{k}} \exp\left(\frac{1}{k}-1\right) \;\text{ at }\; z = \frac{h_{\text{mode}}}{H_0} = B + A\left(1-\frac{1}{k}\right)^{1/k} \qquad (3.41)$$

3.9.2 Longshore Currents by Directional Random Waves

$s_{\max}\downarrow$, more directional spreading, smaller longshore current velocity (see Fig. 3.53).