Part 2. The Hydrosphere

Chapter 11. Gases in Water

Chapter 11. Gases in Water 11.1 Simple gases

Henry's Law:

 $P_g=KX_1$, $P_g=$ partial pressure of the gas in the bulk atmosphere/Pa, K is the Henry's law constant/Pa, X_1 is the equil. mole fraction of the solute in the bulk liquid (aqueous) phase

 Practical form: (at low conc.) reciprocal of Henry's law

 $[G]_{f=}K_{H}P_{g}$, $[G]_{I}$ is the equil. conc. of the solute in the bulk liquid (aqueous) phase/mol L⁻¹, K_{H} is the Henry's law constant/mol L⁻¹ P_{a}^{-1} , P_{g} =partial pressure of the gas in the bulk atmosphere/Pa.

Table 11.1 Henry's law constants forselected gases dissolved in water at 25 ℃	
Gas	K _H /mol L⁻¹ Pa⁻¹
0 ₂	1.3 x 10 ⁻⁸
N ₂	6.4 x 10 ⁻⁹
CH ₄	1.3 x 10 ⁻⁸
CO ₂	3.3 x 10 ⁻⁷
SO ₂	1.2 x 10 ⁻⁵
NH ₃	5.7 x 10 ⁻⁴
Hg	8.6 x 10 ⁻⁷
	3.7 x 10 ⁻⁷
	3.9 x 10 ⁻³

Chapter 11. Gases in Water 11.1 Simple gases

• E.g. Equil. conc. of oxygen in water at 25 °C

$$P_{O_{2}} = (P^{\circ} - P_{H_{2}O})X_{O_{2}}$$
where P° = atmospheri c pressure at 25°C is 3.2×10³ Pa.

$$P_{O_{2}} = (1.01 \times 10^{5} - 3.2 \times 10^{3})Pa \times 0.209$$
Mixing ratio of O₂ in air
= 2.04×10⁴ Pa.
Using eqn 11.2:

$$[O_{2}]_{aq} = 1.3 \times 10^{-8} molL^{-1}Pa^{-1} \times 2.04 \times 10^{4} Pa$$
= 2.7×10⁻⁴ molL⁻¹
= 8.5mgL⁻¹

 Solubility of oxygen: 5 ⁰C 12.4 mg/L 30 ⁰C 7.5 mg/L

- E.g. CO₂, SO₂ and NH₃
- Dissolution of CO₂ in water

 $CO_2(aq) + H_2O \rightarrow H_2CO_3(aq) K_r = [H_2CO_3]/[CO_2(aq)] = 2 \times 10^{-3}$ $K_{a1}' = 2 \times 10^{-4}$ $H_2CO_3(aq) \rightarrow HCO_3(aq) + H^+$ $K_{a2} = 4.7 \times 10^{-11}$ $HCO_3^-(aq) \rightarrow CO_3^{2-}(aq) + H^+$ $HCO_3^{-}(aq) + H^+ \rightarrow CO_2(aq) + H_2O$ $K_{a1} = K_{a1}$ ' x $K_r = 4.5 \times 10^{-7}$ $K_{H} = 3.3 \times 10^{-7} \qquad K_{r} = 2 \times 10^{-3} \qquad K_{a1} = 2 \times 10^{-4} \qquad K_{a2} = 4.7 \times 10^{-11}$ $CO_{2} (g) = CO_{2} (aq) = H_{2}CO_{3} (aq) = HCO_{3}^{-} (aq) = CO_{3}^{2-} (aq)$ $K_{\rm e1} = K_{\rm e1} \times K_{\rm e} = 4.5 \times 10^{-7}$ (11.5)

In the atmosphere where its mixing ratio is 365 ppmv, the partial pressure of carbon dioxide is

 $P_{CO_2} = (1.01 \times 10^5 - 3.2 \times 10^3) \times 365 \times 10^{-6} Pa = 35.8 Pa$

Dissolution of CO₂ in water (cont'd) Applying eqn 11.2, the concentration of CO₂ (aq) is then calculated to be

$$G_{l} = [CO_{2} (aq)] = 3.3 \times 10^{-7} \text{ mol } L^{-1} \text{ Pa}^{-1} \times 35.8 \text{ Pa} = 1.2 \times 10^{-5} \text{ mol } L^{-1}$$

The other aqueous carbonate species are determined using the standard acid-base relations. In these and other calculations, the square brackets refer to aqueous molar concentrations, an approximation for activities as described in Chapter 10.

$$H_2CO_3(aq)] = K_r \times [CO_2(aq)] = 2.4 \times 10^{-8} \text{ mol } L^{-1}$$
 (11.6)

$$K_{a1} = \frac{[\text{HCO}_3^-][\text{H}_3\text{O}^+]}{[\text{CO}_2(\text{aq})]} = 4.5 \times 10^{-7}$$
(11.7)

If, besides carbon dioxide, there are no additional sources of acid or carbonate species:

$$[HCO_3^-] = [H_3O^+] = ([CO_2(aq)] \times K_{a1})^{\frac{1}{2}} = \underline{2.2 \times 10^{-6} \text{ mol } L^{-1}}$$
(11.8)

$$[CO_3^{2-}] = \frac{[CO_2(aq)] \times K_{a1} \times K_{a2}}{[H_3O^+]^2} = 5.0 \times 10^{-11} \text{ mol } L^{-1}$$
(11.9)

Dissolution of CO₂ in water (cont'd)
 In many environmental situations, water is exposed to atmospheric CO₂ and also in contact with limestone (CaCO₃). So, if the system is at equil.,

$$K_{H}=$$

$$K_{a1} = \frac{[\text{HCO}_3^-][\text{H}_3\text{O}^+]}{[\text{CO}_2]} = 4.5 \times 10^{-7}$$
(11.11)

$$K_{a2} = \frac{[CO_3^{2-}][H_3O^+]}{[HCO_3^-]} = 4.7 \times 10^{-11}$$
(11.12)

$$K_{\rm sp} = [{\rm Ca}^{2+}][{\rm CO}_3^{2-}] = 5 \times 10^{-9}$$
 (11.13)

In any solution, the total ionic positive and negative charge must be equal. In the present case, this charge balance is described by eqn 11.14.

$$2[Ca^{2+}] + [H_3O^+] = [HCO_3^-] + 2[CO_3^{2-}] + [OH^-]$$
(11.14)

Dissolution of CO₂ in water (cont'd)

From Fig. 1.2, in the pH range 6-9:

 $[H_3O^+] \ll [Ca^{2+}]$ $[CO_3^{2-}]$ and $[OH^-] \ll [HCO_3^-].$

therefore

$$2 \times [Ca^{2+}] = [HCO_3^{-}] \tag{11.15}$$

From eqns. 11.10, 11.11, and 11.12

$$CO_{3}^{2-}] = \frac{K_{a1} \times K_{a2} \times K_{H} \times P_{CO_{2}}}{\left[H_{3}O^{+}\right]^{2}}$$
(11.16)

From eqns. 11.13 and 11.16

$$Ca^{2+}] = \frac{K_{sp}[H_3O^+]^2}{K_{a1} \times K_{a2} \times K_H \times P_{CO_2}}$$
(11.17)

From eqns. 11.10, 11.11, 11.15, and 11.17

$$[HCO_{3}^{-}] = \frac{P_{CO_{2}}K_{H}K_{a1}}{[H_{3}O^{+}]} = \frac{2K_{sp}[H_{3}O^{+}]^{2}}{K_{a1} \times K_{a2} \times K_{H} \times P_{CO_{2}}}$$
(11.18)

From eqn 11.18

 $\left[H_{3}O^{+}\right]^{3} = \tag{11.19}$

We showed above that the atmospheric partial pressure of carbon dioxide is 35.8 Pa and the equilibrium aqueous concentration is 1.2×10^{-5} mol L⁻¹. From eqn 11.19

$$[H_3O^+] = \left(\frac{(3.3 \times 10^{-7})^2 \times (4.5 \times 10^{-7})^2 \times 4.7 \times 10^{-11} \times (35.8)^2}{2 \times 5 \times 10^{-9}}\right)^{\frac{1}{3}}$$

= 5.1 × 10⁻⁹ mol L⁻¹
pH = -log[H₃O⁺] = 8.3





Dissolution of CO₂ in water (cont'd)

From eqn 11.18



$$\begin{aligned} \text{La}^{-1} &= \frac{1}{4.5 \times 10^{-7} \times 4.7 \times 10^{-11} \times 3.3 \times 10^{-7} \times 35.8} \\ &= 5.2 \times 10^{-4} \text{ mol } \text{L}^{-1} \end{aligned}$$

Fig. 11.1 Solubility of calcium carbonate as a function of the partial pressure of carbon dioxide in the atmosphere.

The calculations confirm that the two most important species contributing to mass balance are Ca^{2+} and HCO_3^{-} . The pH of the solution is 8.30. Undissociated CO_2 and CO_3^{2-} are present only in small concentrations.

- Alkalinity of water with dissolved carbonate species
- Alkalinity: a measure of the ability of a water body to neutralize acids, i.e. proton-accepting ability, and is very important in predicting the extent of acidification of lakes and rivers

alkalinity =
$$[OH^{-}] + [HCO_{3}^{-}] + 2[CO_{3}^{2-}] - [H_{3}O^{+}]$$
 (11.20)

A related but broader concept is the acid neutralizing capacity (ANC):

$$ANC = [OH^{-}] + [HCO_{3}^{-}] + 2[CO_{3}^{2-}] + [B(OH)_{4}^{-}] + [H_{3}SiO_{4}^{-}] + [HPO_{4}^{2-}] + [HS^{-}] + [NOM^{-}] + ... - [H_{3}O^{+}] - 3[AI^{3+}] - ...$$

(11.21)

In this definition, NOM refers to

Alkalinity of water with dissolved carbonate species (cont'd)

$$B(OH)_4^-(aq) + H_3O^+(aq) \rightarrow H_3BO_3 + 2H_2O$$
 pKa=9.14 (11.22)

$$H_3SiO_4^-(aq) + H_3O^+(aq) \rightarrow Si(OH)_4 + H_2O$$
 pKa=9.66 (11.23)

$$HPO_4^{2-}(aq) + H_3O^+(aq) \rightarrow H_2PO_4^-(aq) + H_2O$$
 pKa=7.20 (11.24)

$$HS^{-}(aq) + H_{3}O^{+}(aq) \rightarrow H_{2}S + H_{2}O$$
 pKa=7.00 (11.25)

-Above rxns have negligible effects on neutralization

-carbonate species and hydroxyl ions are the most significant species to alkalinity

-typical alkalinity value = $50 \sim 2000 \ \mu mol \ L^{-1}$

N.B. acid-base property

In terms of **pH**: intensity factor-measures the conc. of alkali or acid immediately available for reaction

In terms of **alkalinity**: capacity factor-a measure of the ability of a water sample to sustain reaction with added acids or base. Can have negative value, meaning that extensive acidification has already occurred so that there is an excess of proton-donating species

Alkalinity of water with dissolved carbonate species (cont'd)



Fig. 11.2 Lake status in terms of acid/base properties.

- N.B. Sensitivity classification can also be in terms of alkalinity using units proton-accepting capacity,
- -high value of alkalinity means that a large amount of acid would have to be added to bring the lake into the transitional and acidic ranges