

## Lecture 3

# Tectonic stress field and its determination

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# Introduction

## Objectives of the course

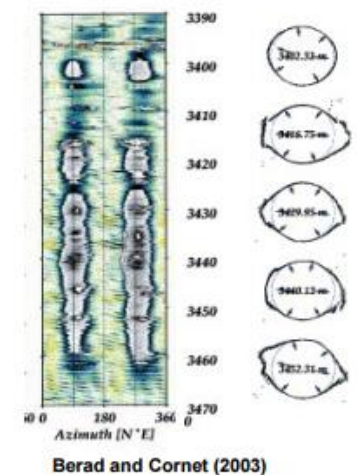
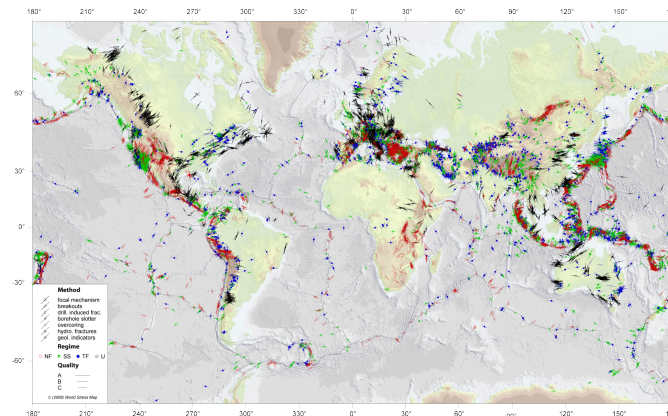
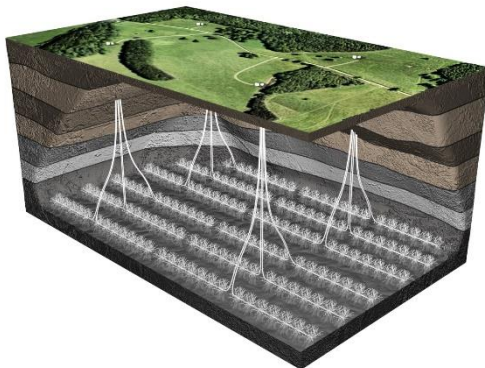
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- Objective;
  - Understand the importance of in situ stress for geomechanics
  - In situ stress measurement;
  - Estimation
    - ↗ Integrated method
    - ↗ World Stress

- Boundary condition for a engineering problem
  - In situ stress orientation and magnitude is a critical factor for various geomechanics applications
    - ✎ Hydraulic fracturing
    - ✎ Borehole stability
    - ✎ Earthquake analysis
    - ✎ Tunnel/mine/opening design/stability



# Introduction

## Definition of stress



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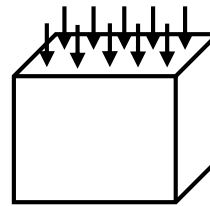
### – Stress

↯: a force acting over a given area,  $F/A \leftarrow$  simple definition

↯ the internal distribution of force per unit area that balances and reacts to external loads applied to a body  $\leftarrow$  exact definition

### – Normal stress: Normal force/Area

$$\sigma = \frac{F_n}{A}$$



### – Shear stress: Shear force/Area

$$\tau = \frac{F_s}{A}$$

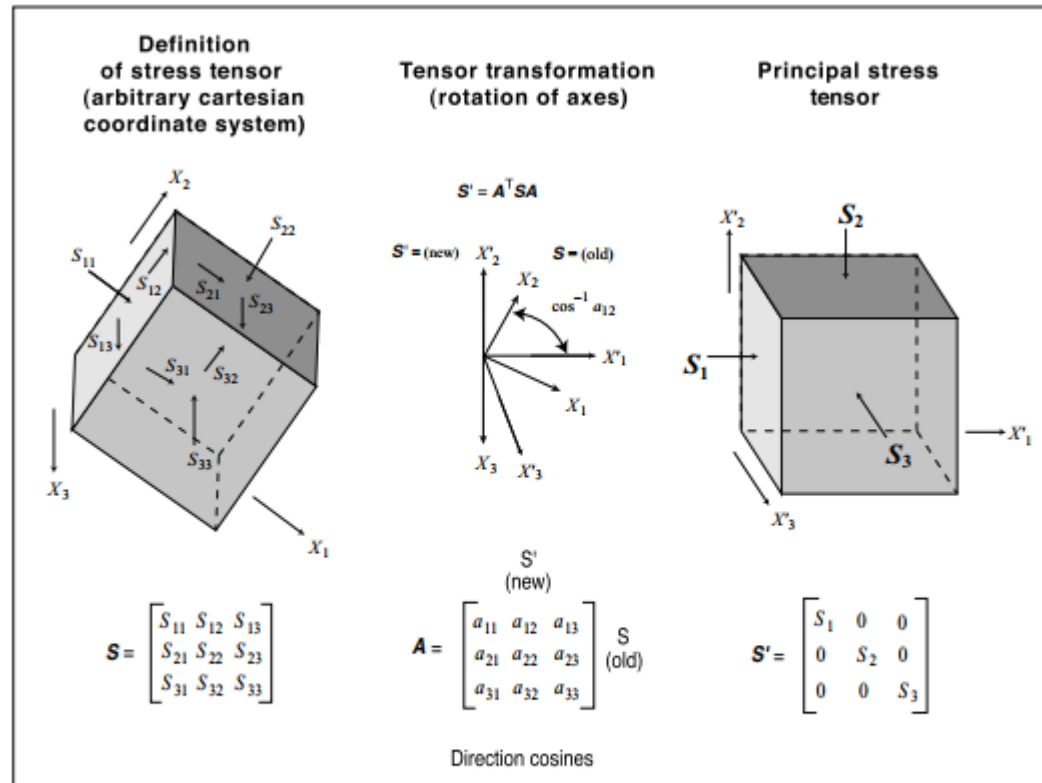
– Unit:  $\text{N/m}^2 = \text{Pa}$ ,  $10^6 \text{Pa} = \text{MPa}$ ,  $10^9 \text{Pa} = \text{GPa}$

$145 \text{ psi} = 1 \text{ MPa} = 10 \text{ bar} = 10 \text{ kgf/cm}^2$

# Introduction

## Definition of stress

- Six components  $\rightarrow$  three components + three angles



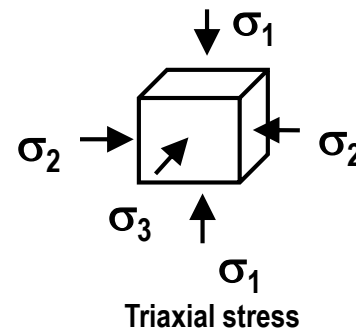
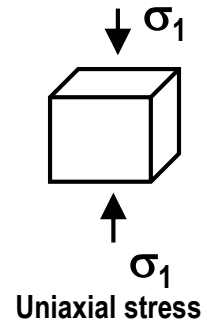
**Figure 1.1.** Definition of stress tensor in an arbitrary cartesian coordinate system (Engelder and Leftwich 1997), rotation of stress coordinate systems through tensor transformation (center) and principal stresses as defined in a coordinate system in which shear stresses vanish (right).

# Introduction

## State of stress underground



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Close to the reality in  
deep underground

- Structural Engineering
  - Stress (force) can be **directly** measurable or,
  - Stress (force) can be **calculated**
  - Less uncertainties
- Underground Engineering
  - Stress (force) has to be **indirectly measured**
  - Stress (force) has to be estimated
  - Large uncertainties



What is my weight? Hence force and stress at my ankles?



What is in situ stress here?

# Introduction

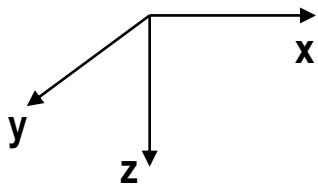
## Prediction of in situ stress



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- Heim's rule
  - Assumption: no lateral deformation

$$\begin{pmatrix} \varepsilon_x \\ \varepsilon_y \\ \varepsilon_z \\ \gamma_{yz} \\ \gamma_{xz} \\ \gamma_{xy} \end{pmatrix} = \begin{pmatrix} \frac{1}{E} & -\frac{\nu}{E} & -\frac{\nu}{E} & 0 & 0 & 0 \\ -\frac{\nu}{E} & \frac{1}{E} & -\frac{\nu}{E} & 0 & 0 & 0 \\ -\frac{\nu}{E} & -\frac{\nu}{E} & \frac{1}{E} & 0 & 0 & 0 \\ 0 & 0 & 0 & \frac{1}{G} & 0 & 0 \\ 0 & 0 & 0 & 0 & \frac{1}{G} & 0 \\ 0 & 0 & 0 & 0 & 0 & \frac{1}{G} \end{pmatrix} \begin{pmatrix} \sigma_x \\ \sigma_y \\ \sigma_z \\ \tau_{yz} \\ \tau_{xz} \\ \tau_{xy} \end{pmatrix}$$



NO. \_\_\_\_\_  
DATE \_\_\_\_\_

$$\begin{aligned} 0 &= \frac{1}{E} \sigma_x - \frac{\nu}{E} \sigma_y - \frac{\nu}{E} \sigma_z \\ 0 &= -\frac{\nu}{E} \sigma_x + \frac{1}{E} \sigma_y - \frac{\nu}{E} \sigma_z \\ \varepsilon_z &= -\frac{\nu}{E} \sigma_x - \frac{\nu}{E} \sigma_y + \frac{1}{E} \sigma_z \end{aligned}$$

$$\begin{aligned} \sigma_x &= \nu (\sigma_y + \sigma_z) \\ \sigma_y &= \nu (\sigma_x + \sigma_z) \\ \sigma_x &= \nu (\nu (\sigma_x + \sigma_z) + \sigma_z) \\ &= \nu^2 \sigma_x + \nu (\nu + 1) \sigma_z \\ (1 - \nu^2) \sigma_x &= \nu (\nu + 1) \sigma_z \\ (1 - \nu)(1 + \nu) \sigma_x &= \nu (1 + \nu) \sigma_z \\ \sigma_x &= \frac{\nu}{1 - \nu} \sigma_z \end{aligned}$$

Similarly  $\sigma_y = \frac{\nu}{1 - \nu} \sigma_z$

with  $\nu = \sim 0.25$

$$\sigma_x = \sigma_y = \frac{1}{3} \sigma_z$$

This rule usually works for soils. It does not generally work in rock/reservoir due to tectonic activities, erosion, ...



- Erosion
- topography
- Rock anisotropy
- Discontinuity
- Tectonic activity:
  - Prime interest for Reservoir Engineering ~ km depths

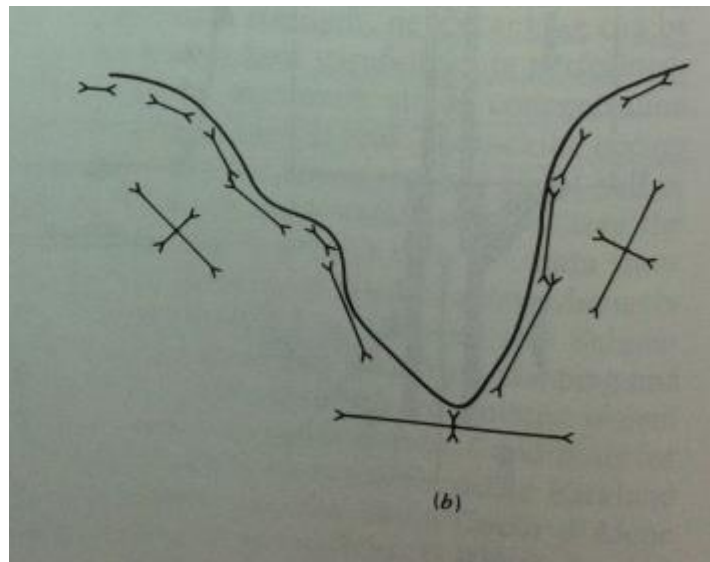
# Factors affecting in situ stress measurement

## Topography



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- Topography



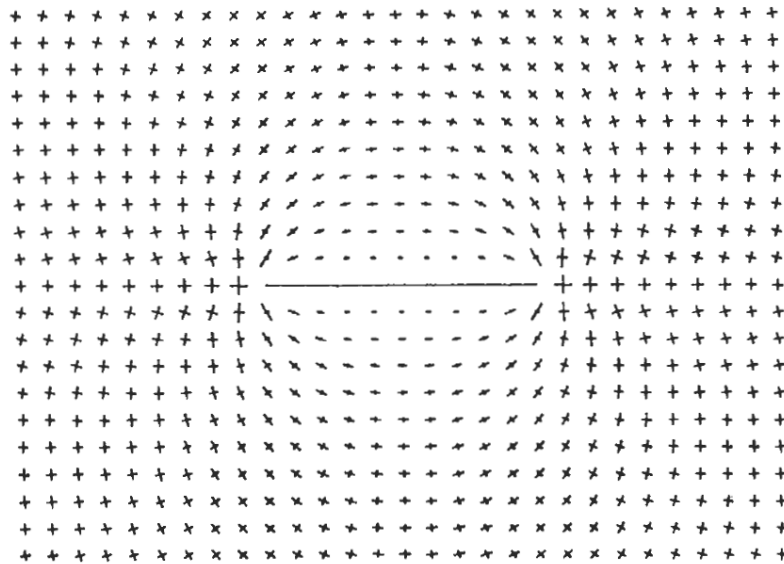
# Factors affecting in situ stress measurement

## Effect of discontinuities



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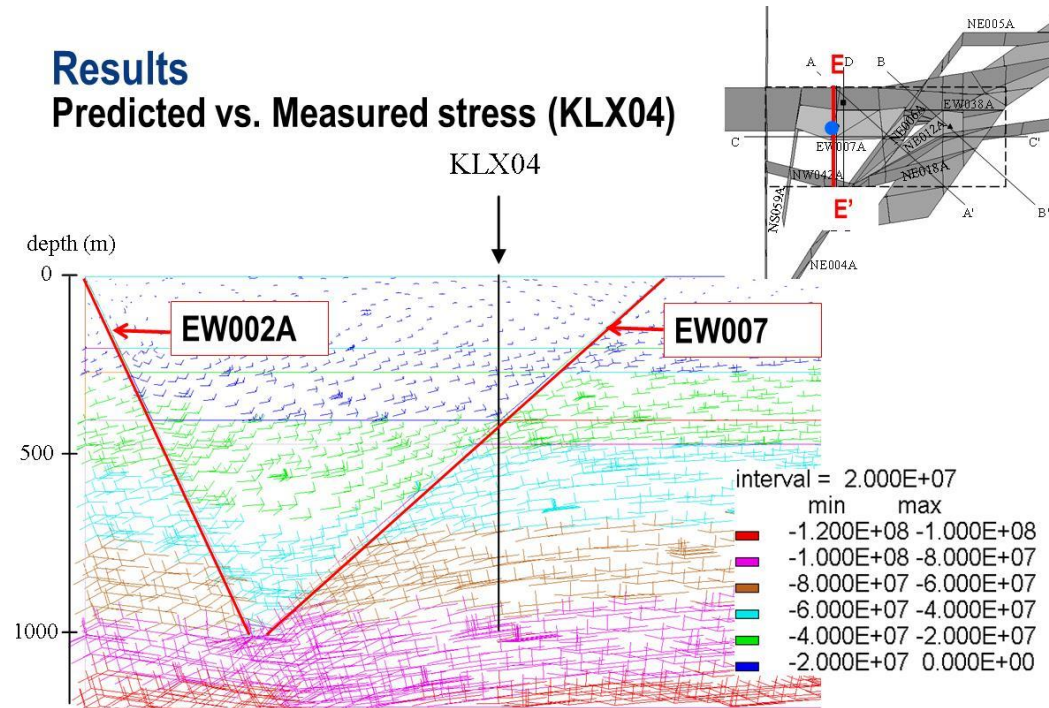
- Discontinuity



**Figure 4.19** Example of the effect of a discontinuity on the near-field stress state, for an applied hydrostatic two-dimensional stress with the discontinuity having a modulus of 10% of the host rock (from Hyett, 1990). The crosses represent the magnitudes and directions of the principal stresses. Note how the stress field close to the discontinuity is quite different from the far-field stress.

### Results

#### Predicted vs. Measured stress (KLX04)



Borehole passes deformation zone EW007 at around 300 ~ 400 m depth

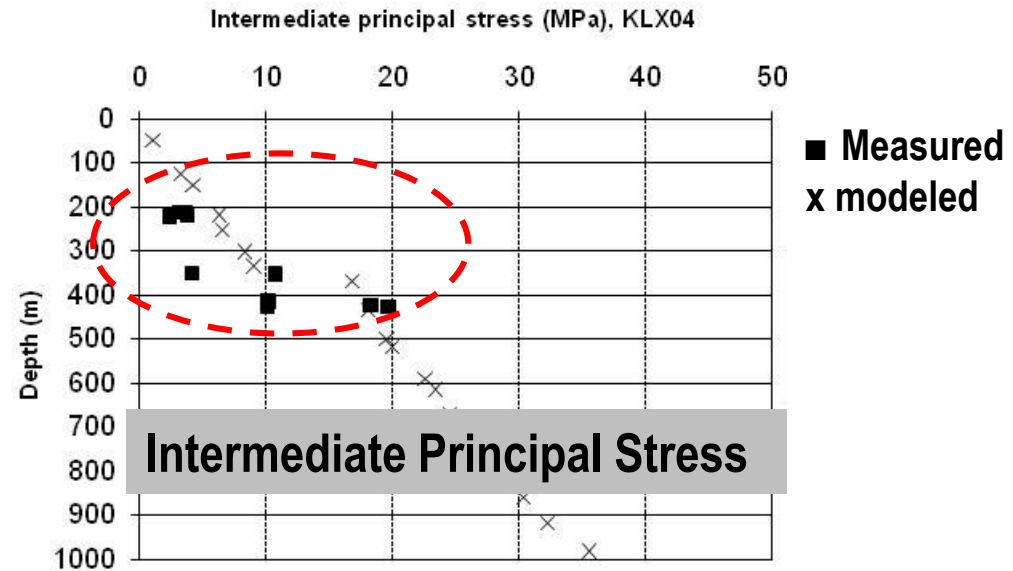
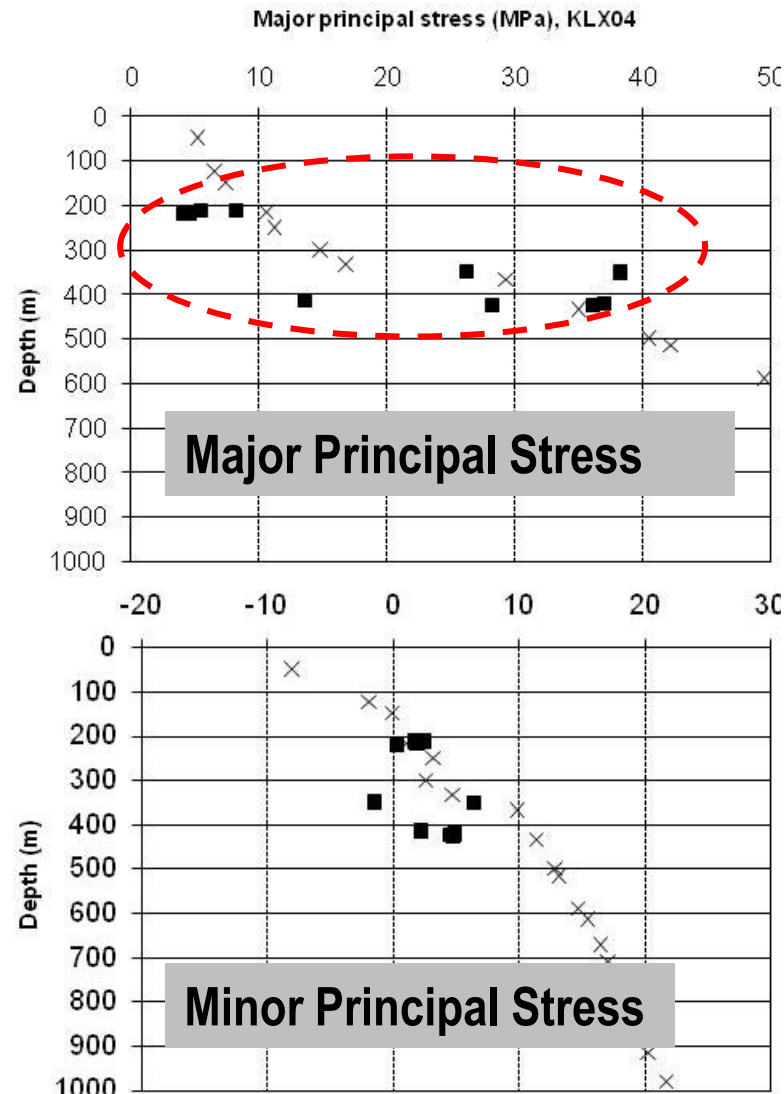
Dramatic change of stress occur due to the deformation zone – smaller stress in the wedge formed by two deformation zones (EW002A and EW007A).

# Factors affecting in situ stress measurement

## Effect of discontinuities - Predicted vs. Measured stress (KLX04)

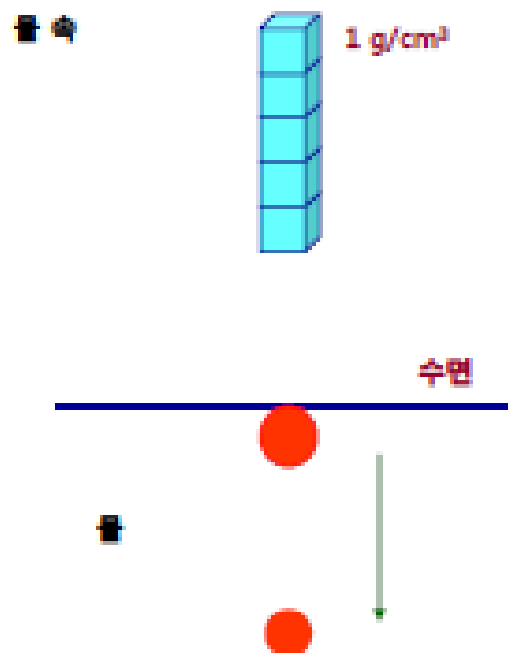


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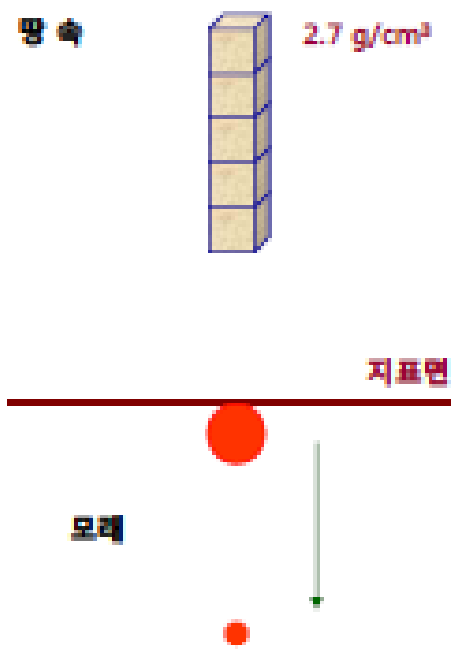


**Stress anomaly due to the presence of large scale fault**

$$p = \rho g z = \text{밀도} \times \text{중력가속도} \times \text{심도}$$



$$p = \rho g z = 0.01 \times z (\text{MPa})$$



$$p = \rho g z = 0.027 \times z (\text{MPa})$$

# Vertical Stress Calculation



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- Calculation of vertical stress
  - onshore

The magnitude of  $S_v$  is equivalent to integration of rock densities from the surface to the depth of interest,  $z$ . In other words,

$$S_v = \int_0^z \rho(z)g dz \approx \bar{\rho}gz \quad (1.5)$$

where  $\rho(z)$  is the density as a function of depth,  $g$  is gravitational acceleration and  $\bar{\rho}$  is the mean overburden density (Jaeger and Cook 1971).

- Offshore

$$S_v = \rho_w g z_w + \int_{z_w}^z \rho(z)g dz \approx \rho_w g z_w + \bar{\rho}g(z - z_w) \quad (1.6)$$

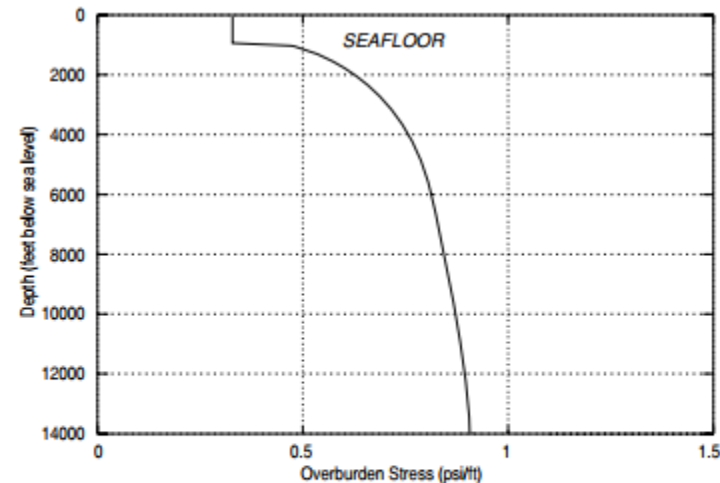
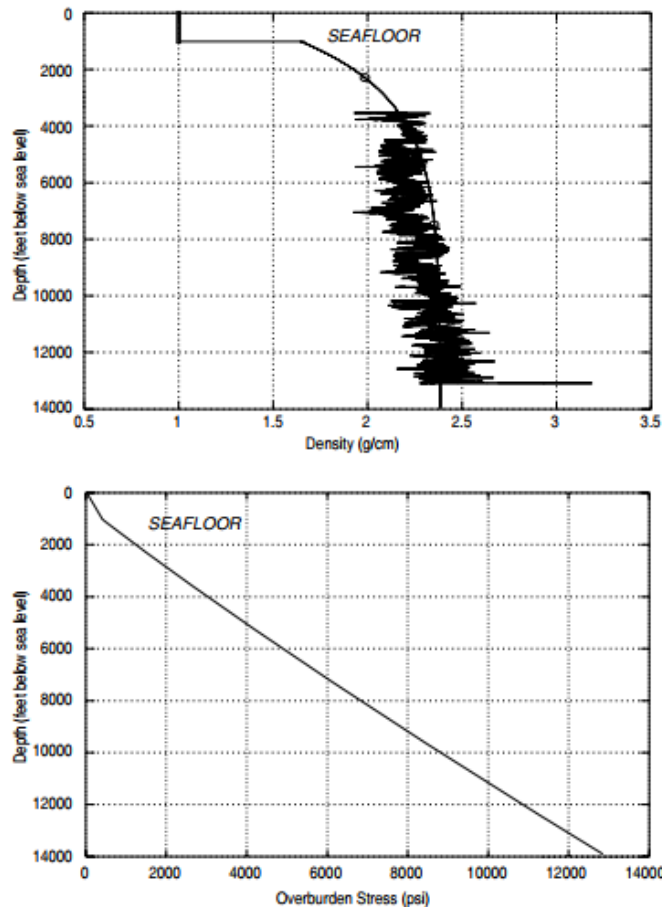
where  $\rho_w$  is the density of water and  $z_w$  is the water depth. As  $\rho_w \sim 1 \text{ g/cm}^3$  (1.0 SG), water pressure (hydrostatic pressure) increases at a rate of 10 MPa/km (0.44 psi/ft).

# Vertical Stress Calculation - example



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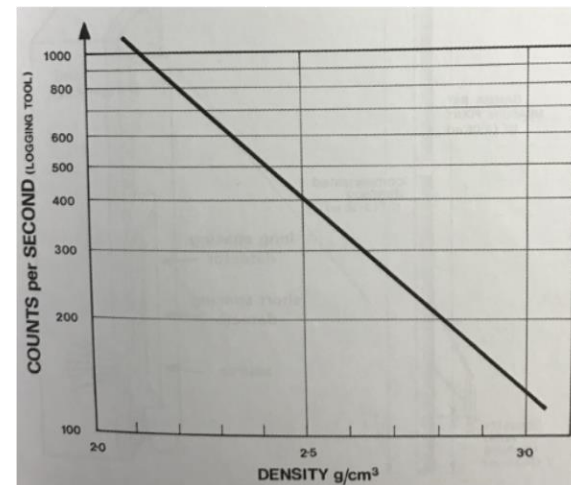
- Estimation of overburden at depth
  - Uncertainty: variability of density log + missing data



**Figure 1.3.** Illustration of how integration of density logs (upper figure) can be used to estimate overburden stress at depth (center figure) or overburden stress gradient (lower figure). Variability of the density logs, as well as the fact that they are often not obtained over the total depth of interest, leads to uncertainty in the calculated overburden (see text).



- Density log
  - Continuous record of a formation's bulk density including solid matrix and fluid in the pores.
- Principal uses
  - Density, or Porosity derived from density
  - Useful lithology indicator
- Principles of measurement
  - Exploit 'gamma-gamma' scattering to measure density
  - A radioactive source emits a continuous beam of high energy gamma rays into the formation → some of these are scattered back to a pair of gamma ray detectors.
  - The count rate at a detector depends on density



Density tool radiation count vs. bulk density

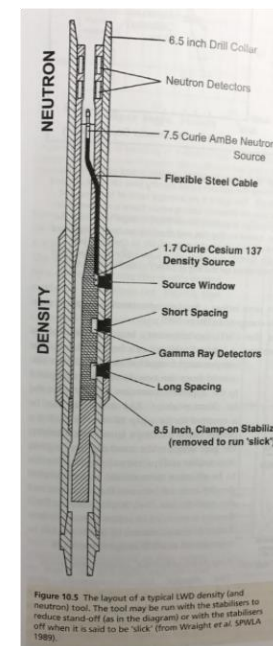


Figure 10.3 The layout of a typical LWD density (and neutron) tool. The tool may be run with the stabilizers to reduce stand-off (as in the diagram) or with the stabilizers off when it is said to be 'slick' (from Wright et al. SPWLA 1989).

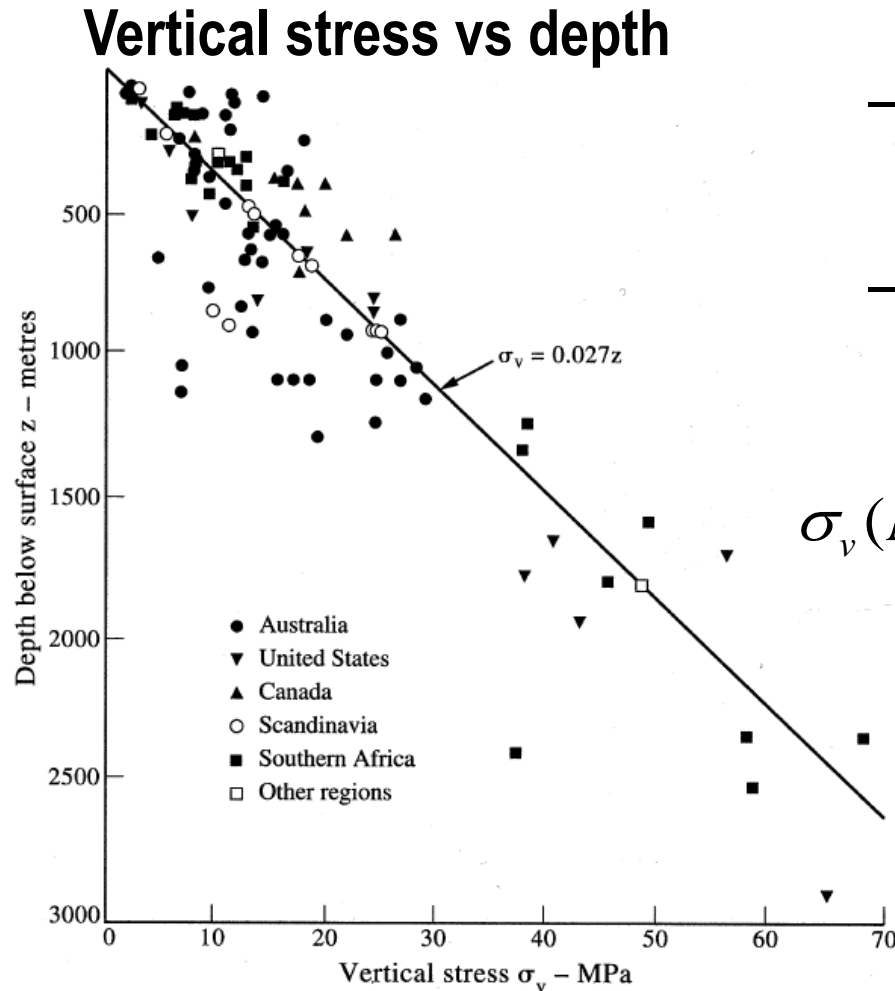


# Vertical Stress

## World wide in situ stress data



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- Vertical component of in situ stress
- More or less similar to predicted stress

$$\sigma_v (MPa) = \rho gh = 0.027 \times h(m)$$

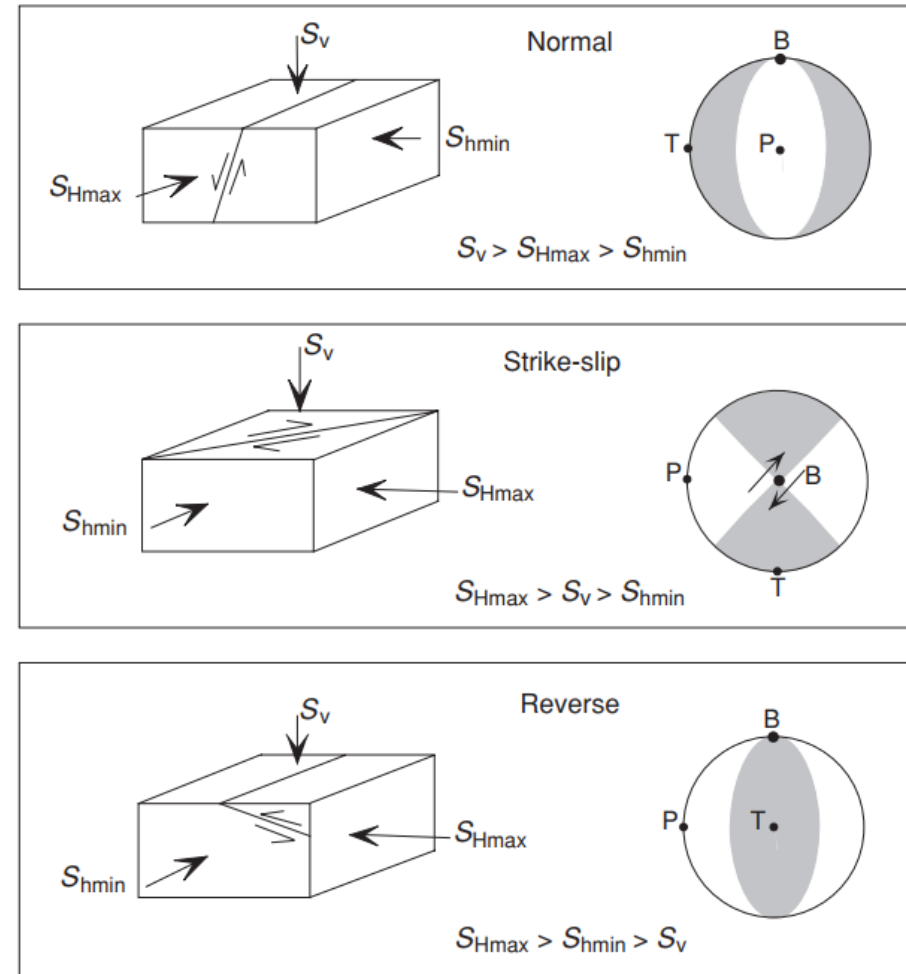
# Anderson's Classification of stress magnitude



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**Table 1.1.** *Relative stress magnitudes and faulting regimes*

Regime	Stress		
	$S_1$	$S_2$	$S_3$
Normal	$S_v$	$S_{Hmax}$	$S_{hmin}$
Strike-slip	$S_{Hmax}$	$S_v$	$S_{hmin}$
Reverse	$S_{Hmax}$	$S_{hmin}$	$S_v$



**Figure 1.2.** E. M. Anderson's classification scheme for relative stress magnitudes in normal, strike-slip and reverse faulting regions. Earthquake focal mechanisms, the *beach balls* on the right, are explained in [Chapter 5](#).

# In Situ Stress Measurement Method

## Direct/Indirect method



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- Direct method
  - Flatjack method
  - USBM overcoring method
  - CSIRO (type) overcoring method
  - Hydraulic fracturing method (mini-frac or micro-frac)
- Indicator method (indirect method)
  - Borehole breakout
  - Anelastic Strain Recovery (ASR)
  - Kaiser effect
  - Core diking
  - Focal Mechanism of earthquake

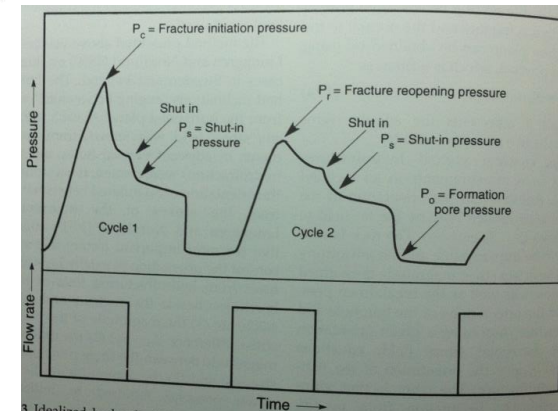
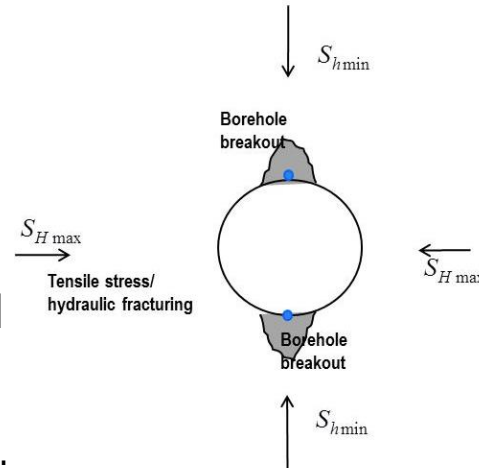
# In Situ Stress Measurement Method

## Direct method



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- Direct method
  - Flatjack method
  - USBM overcoring method
  - CSIRO (type) overcoring method
  - Hydraulic fracturing method (mini-frac or micro-frac) → **More relevant for reservoir engineering at deep depth**



Hydraulic fracturing method

In open circular hole!

$$P_b = 3S_{hmin} - S_{Hmax} + T_o$$

$$P_s = S_{hmin}$$

$$T_o = P_b - P_r$$

or from Lab Test

- ❧ Very different from hydraulic fracturing for enhanced production
- ❧ In Civil/Mining Industry: hydraulic fracturing (in open hole)
- ❧ Petroleum Industry: mini-frac, microfrac (in cased well with perforation). Leak-off test (open hole below cementing)

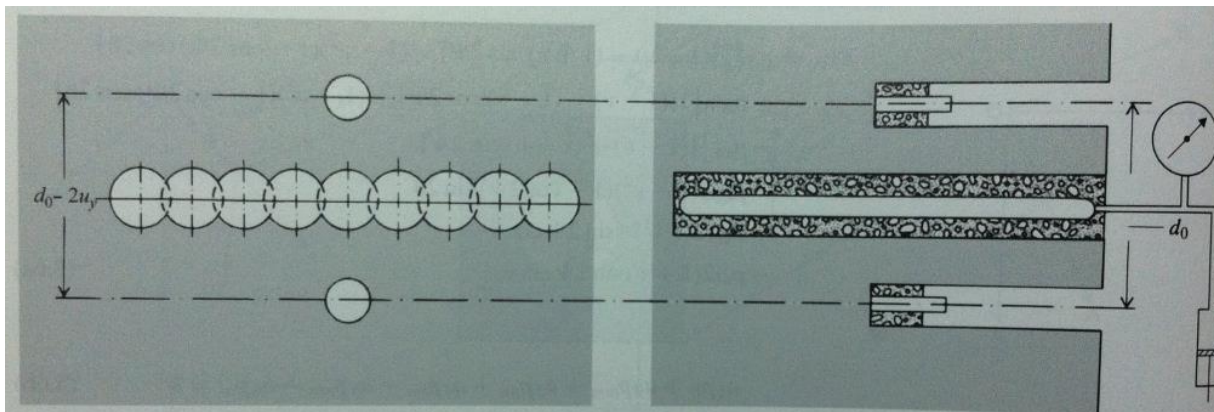
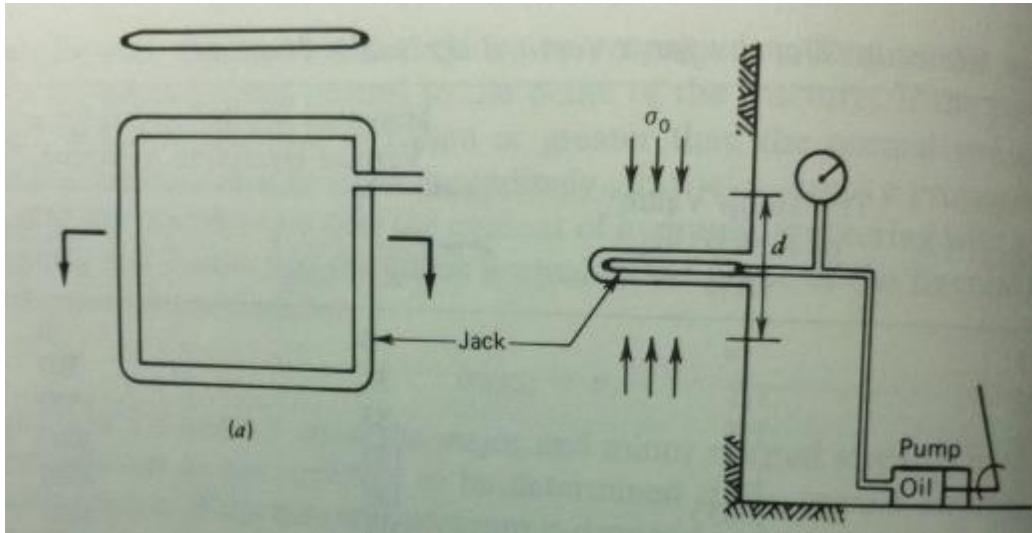
# In Situ Stress Measurement Method

## Direct - Flatjack method



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- Directly measure the tangential stress

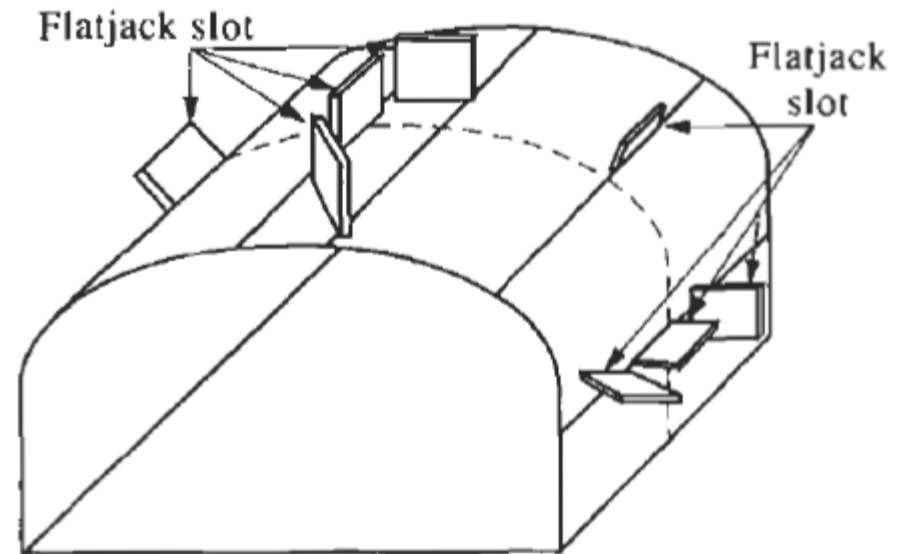
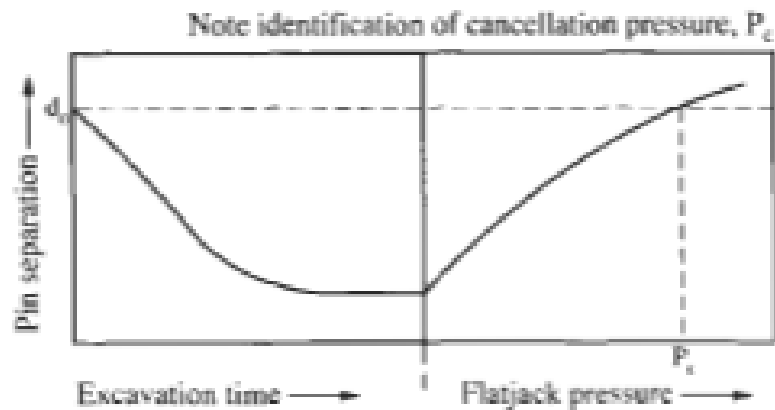


# In Situ Stress Measurement Method

## Direct - Flatjack method



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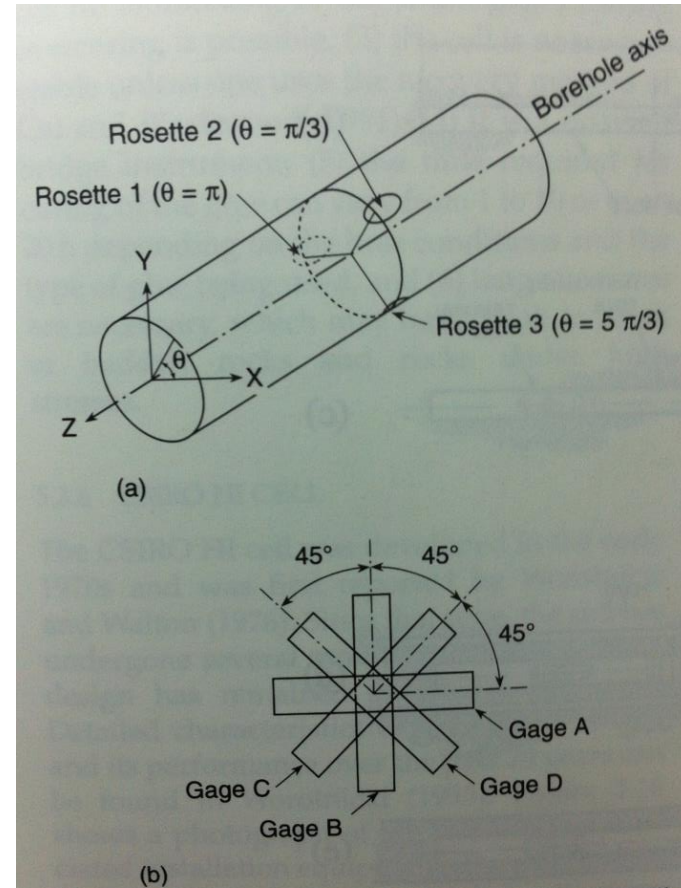
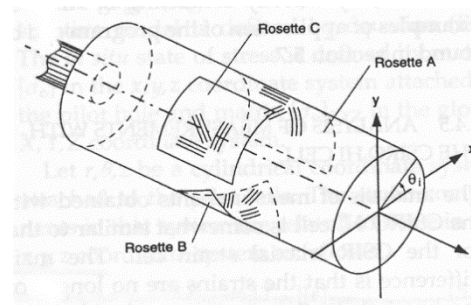
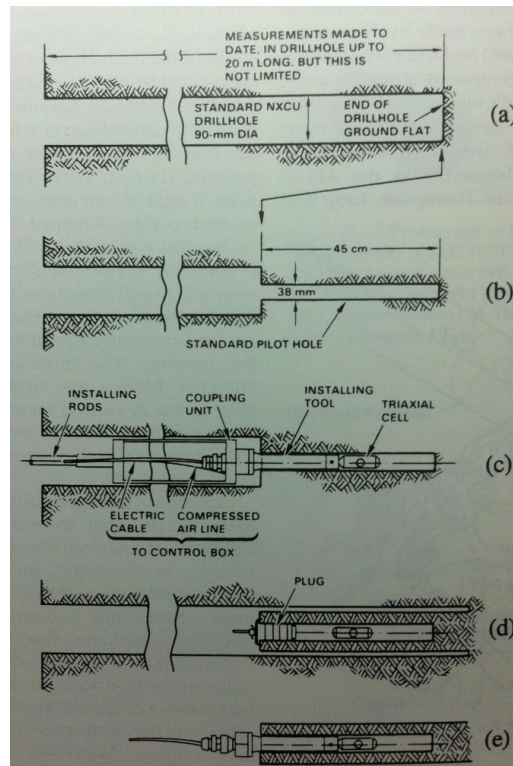




# In Situ Stress Measurement Method

## Direct - CSIRO type overcoring method

- Complete stress tensor can be determined from minimum of six strain gauges.



**Fig. 5.13** Position of the CSIR triaxial strain cell rosettes in (a) and strain gage configuration in (b). (After Van Heerden, 1976.)

# In Situ Stress Measurement Method

## Direct - CSIRO type overcoring method



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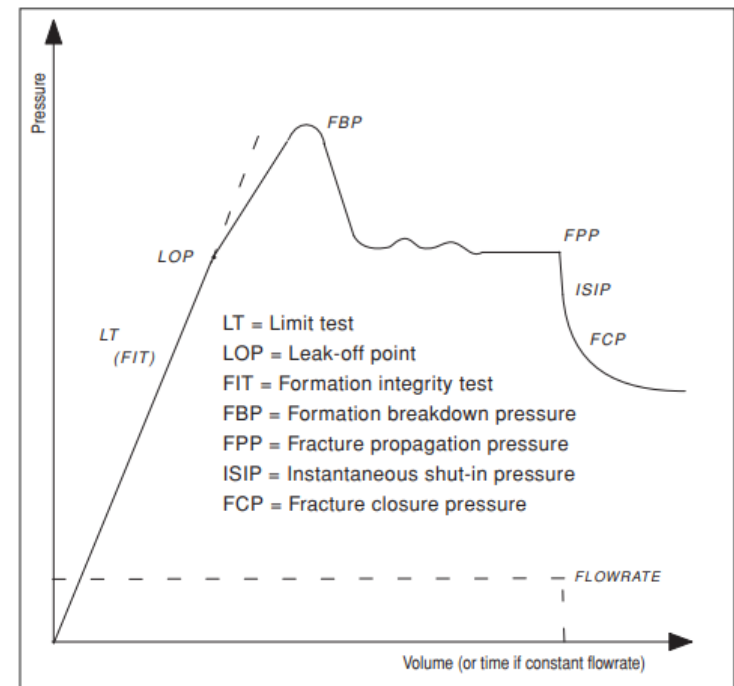
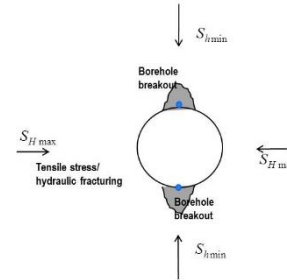
# In Situ Stress Measurement Method

## Direct – mini-frac



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- Mini-frac (or microfrac)
  - In reservoir engineering, only magnitude and direction of minimum principal stress is measured through mini-frac.
  - Usually conducted in cased hole.
  - Leak-off test
    - ✎ In open section after drilling a short distance



**Figure 7.2.** A schematic mini-frac or extended leak-off test showing pressure as a function of volume, or equivalently time (if the flow rate is constant). Modified after Gaarenstroom, Tromp *et al.* (1993). The significance of the various points indicated on the pressure record is discussed in the text.

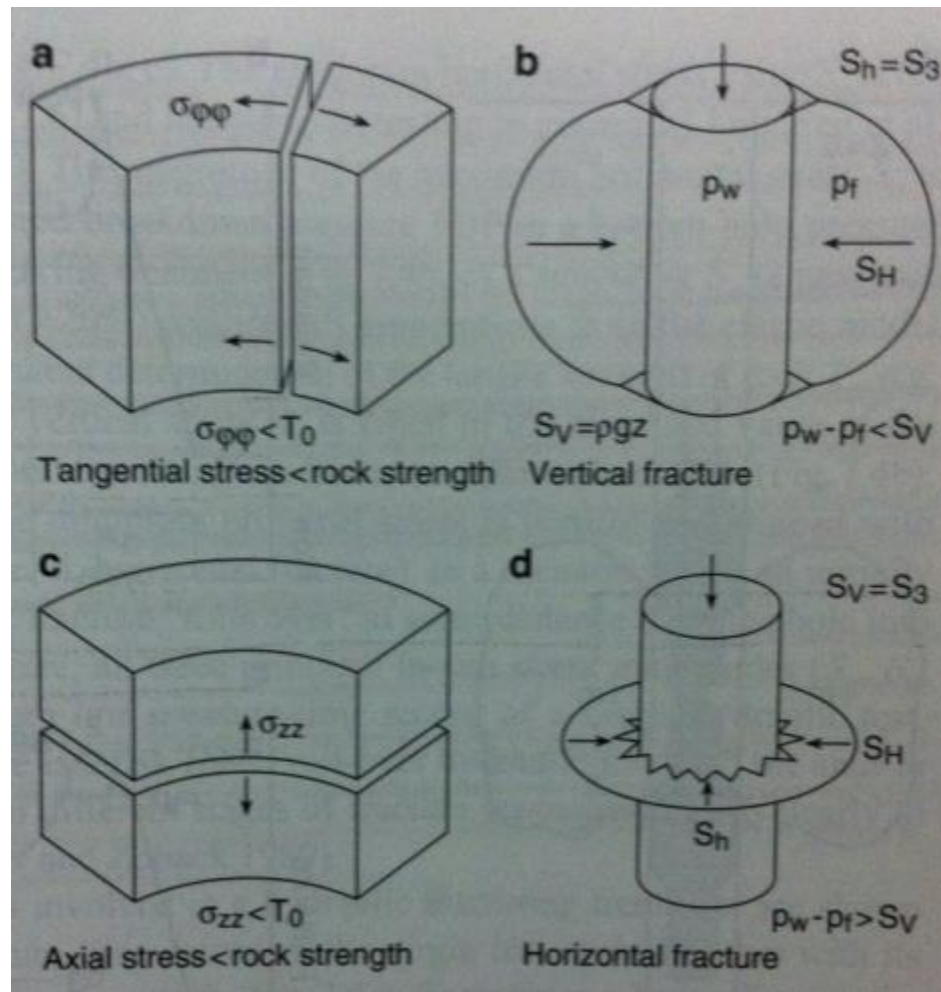
# In Situ Stress Measurement Method

## Direct – mini-frac



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- Vertical fracture vs. horizontal fracture (in vertical hole)



# In Situ Stress Measurement Method

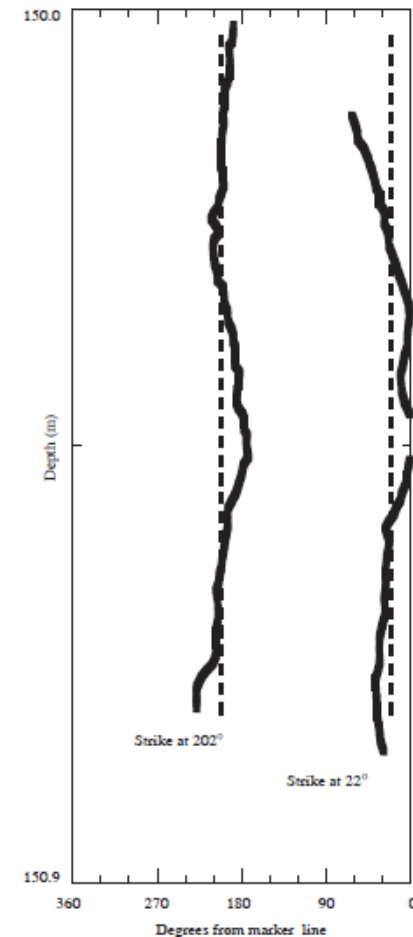
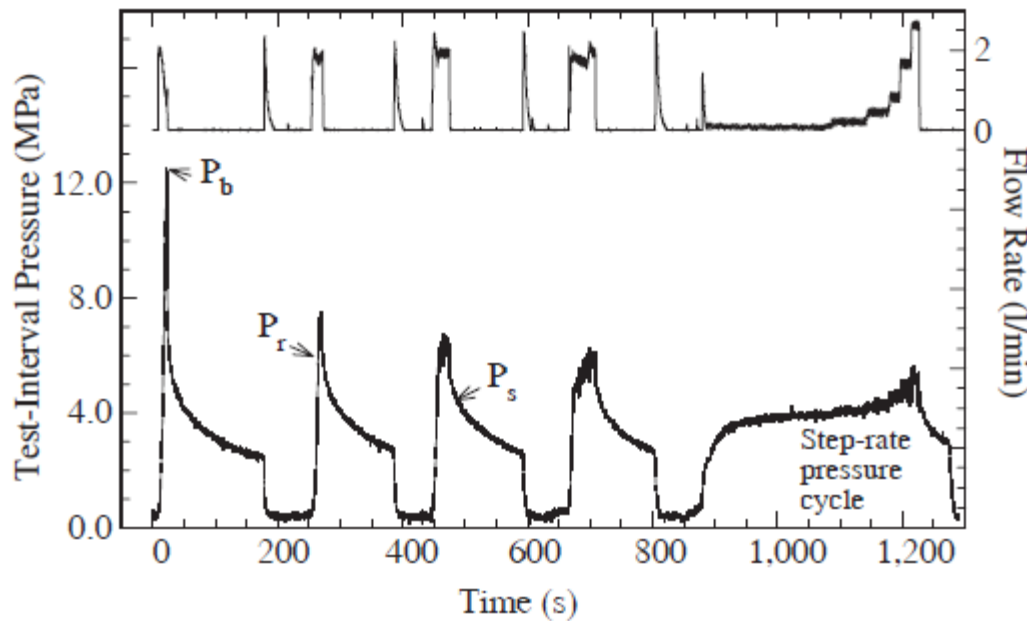
## Direct – mini-frac



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- Actual records

“When open hole is available”



Actual impression packer record  
(Haimson & Cornet, 2003)

# In Situ Stress Measurement Method

## Indirect method



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- Borehole breakout
- Anelastic Strain Recovery (ASR)
- Kaiser effect
- Core diking
- Focal mechanism of earthquake

# In Situ Stress Measurement Method

## Indirect method – borehole breakout

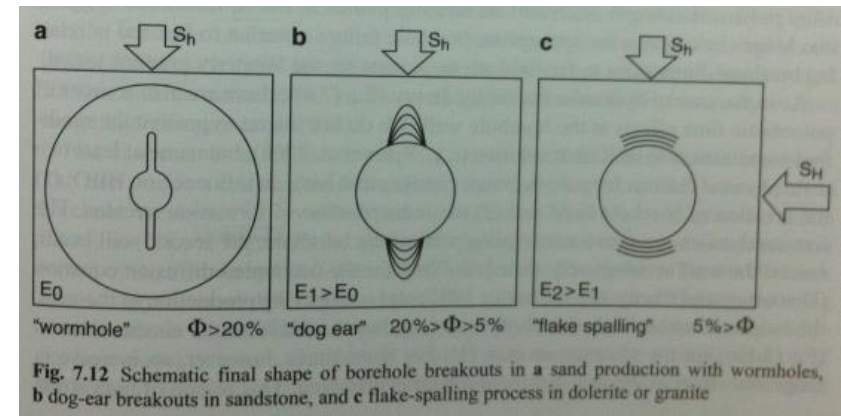
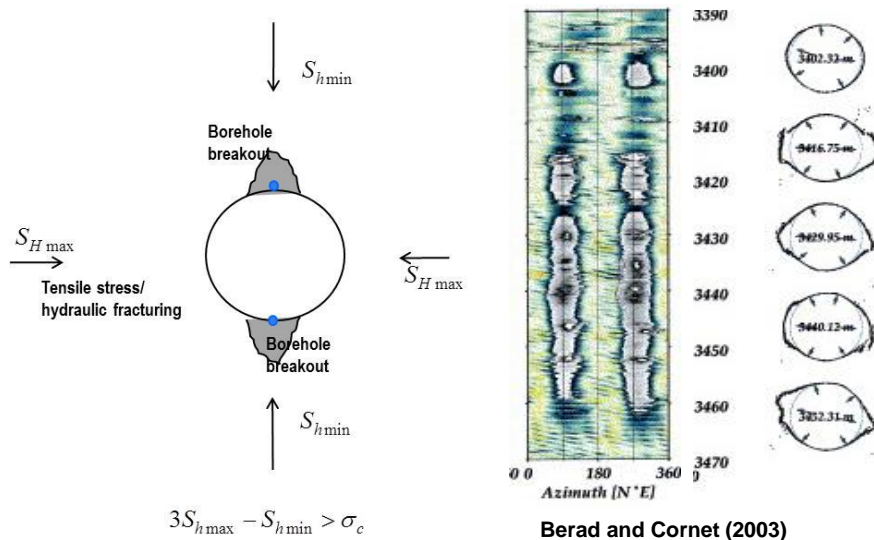


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### • Borehole breakout

- Enlargements of the borehole wall caused by stress-induced failure of wells occurring 180° apart.
- In vertical wells, the diametrically faced zones of broken material occur at direction of minimum horizontal stress.

#### Borehole breakout



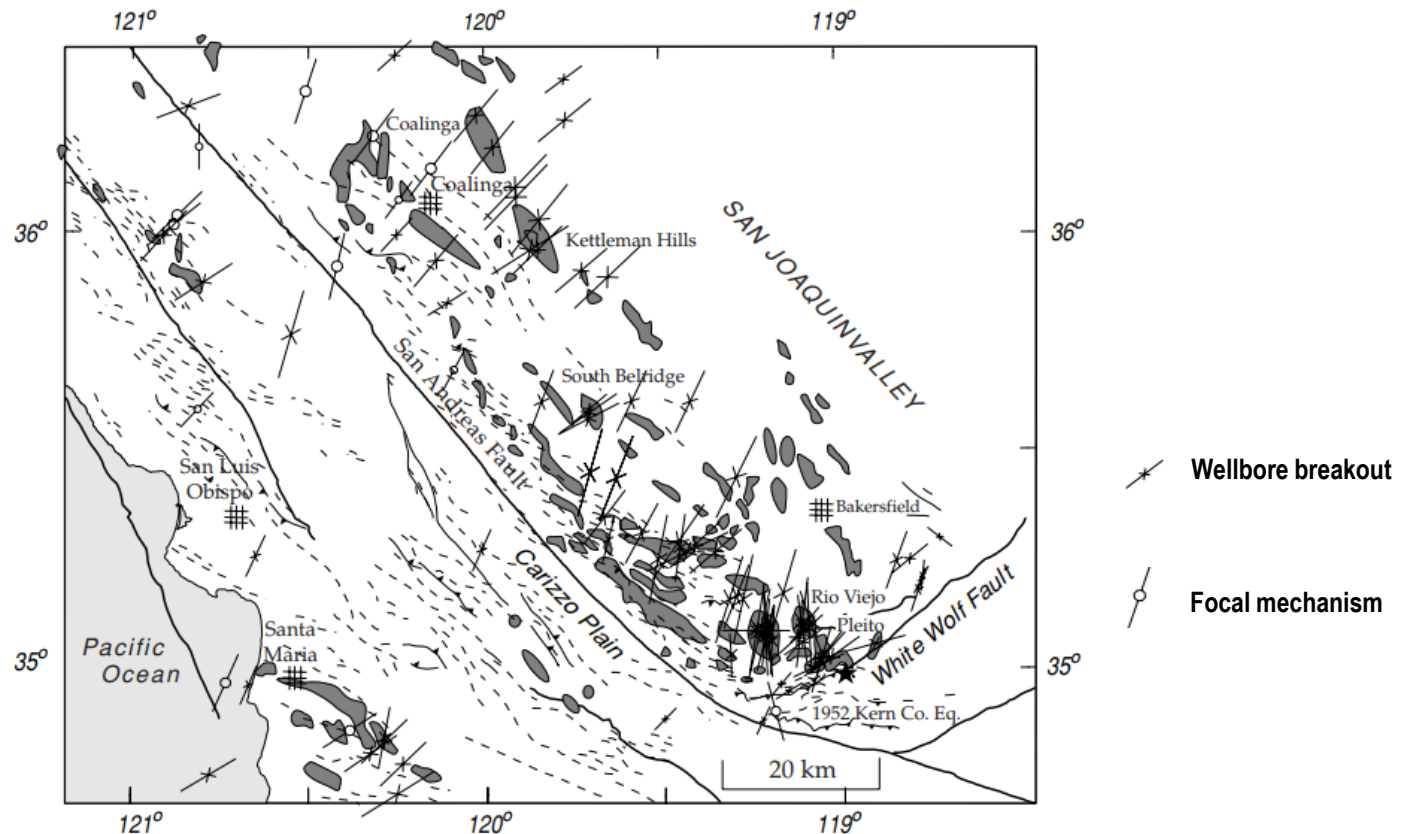
# In Situ Stress Measurement Method

## Indirect method – borehole breakout



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- Stress map of southern San Joaquin



**Figure 1.6.** Stress map of central California (after Castillo and Zoback 1994) showing  $S_{Hmax}$  directions obtained from wellbore breakouts (inward pointed arrows) and earthquake focal plane mechanisms (symbols with open circle). AAPG©1994 reprinted by permission of the AAPG whose permission is required for further use.



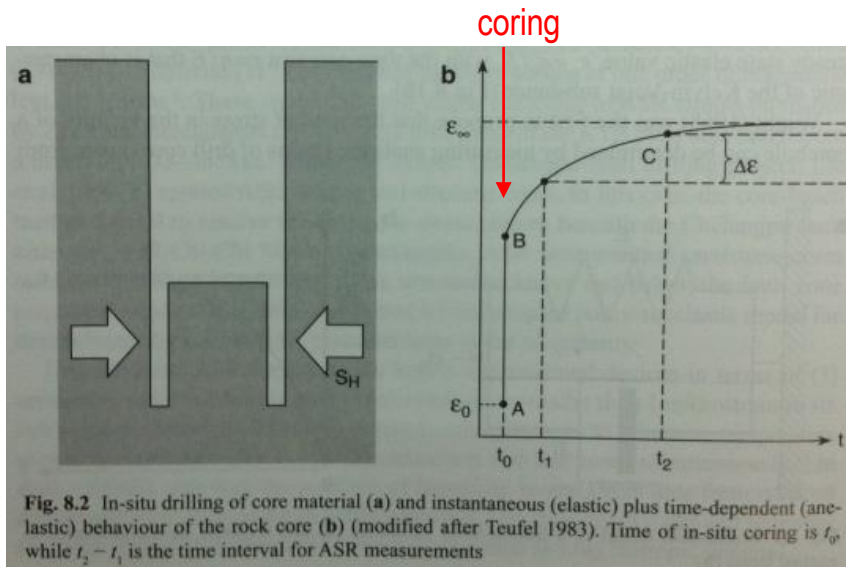
# In Situ Stress Measurement Method

## Indirect method – ASR

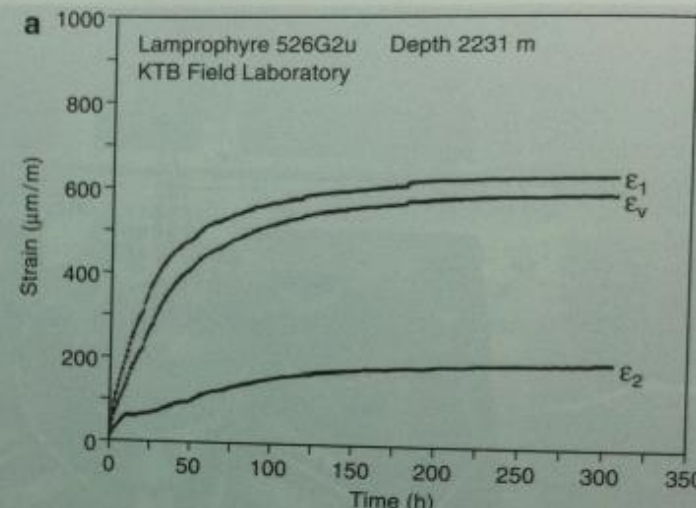


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- Anelastic Strain Recovery (ASR)
  - Core-based method to estimate in-situ stress magnitudes and orientations from instrumenting a freshly recovered drill core obtained from deep wells.
  - The direction of maximum strain recovery is parallel to the maximum horizontal stress in the borehole.



**Fig. 8.6** a Principal recovery strains and b semi-logarithmic, normalized recovery strains, both versus time for quasi-isotropic dyke rock (core #526G2u, depth of 2231 m) from the KTB pilot borehole (modified after Wolter and Berckhemer 1989)



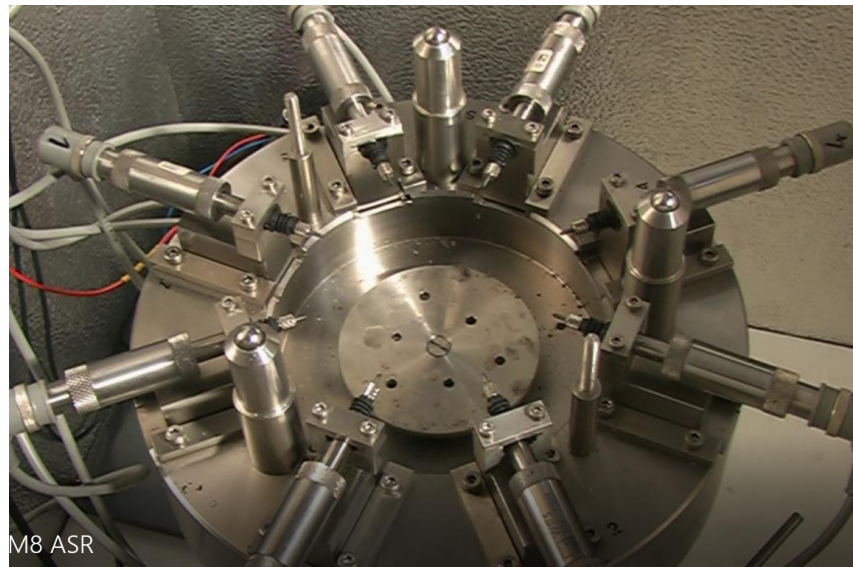
# In Situ Stress Measurement Method

## Indirect method – ASR



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- Anelastic Strain Recovery (ASR) - video





# In Situ Stress Measurement Method

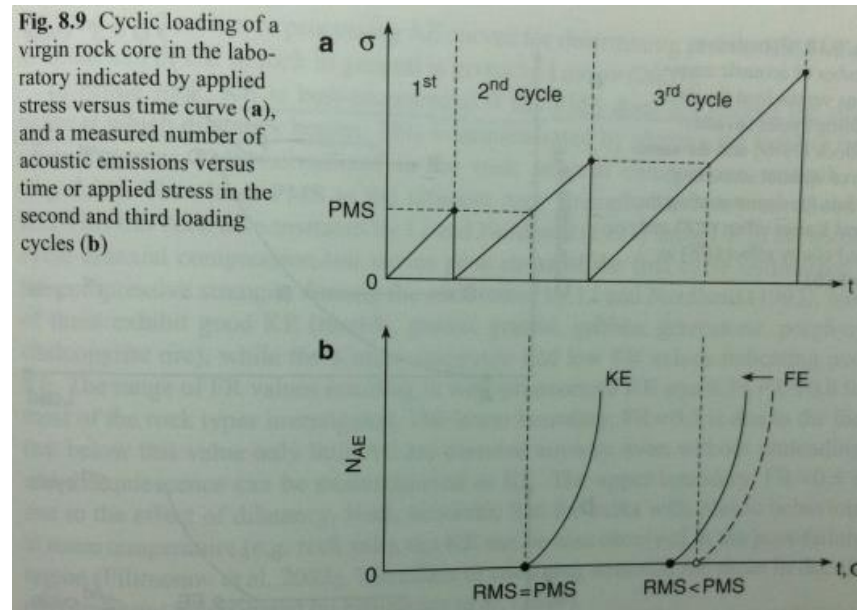
## Indirect method – Kaiser Effect



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- Kaiser effect
  - phenomenon that a material under stress emits acoustic emissions only after the previous maximum stress is reached.
  - Joseph Kaiser (1950, metal, rock and wood in tension)

PMS: previous maximum stress  
RMS: Recalled maximum stress



# In Situ Stress Measurement Method

## Indirect method – core dicing



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- Core dicing
  - Assemblage of cored disks in highly stressed rock
  - Often shaped like a horse saddle (axis ~ maximum horizontal stress)
  - The thinner thickness, the greater the horizontal stress

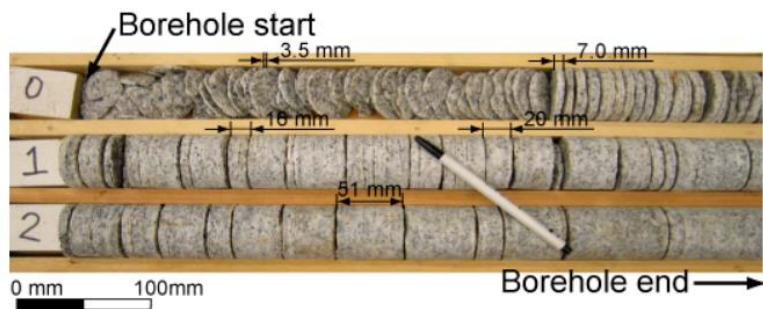
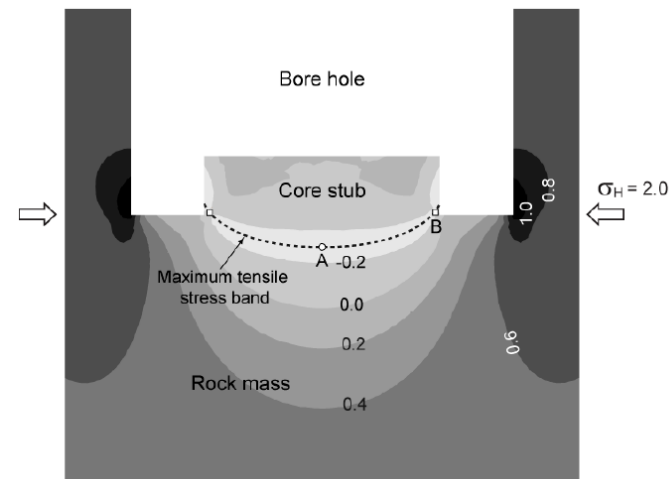


Fig. 1. Typical core dicing observed in 75-mm-diameter boreholes drilled from tunnels at the 420-m depth Level of AECL's Underground Research Laboratory.



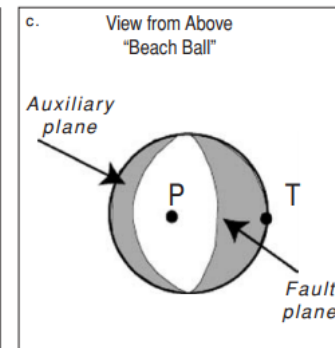
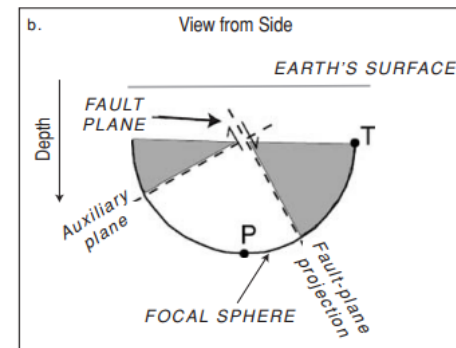
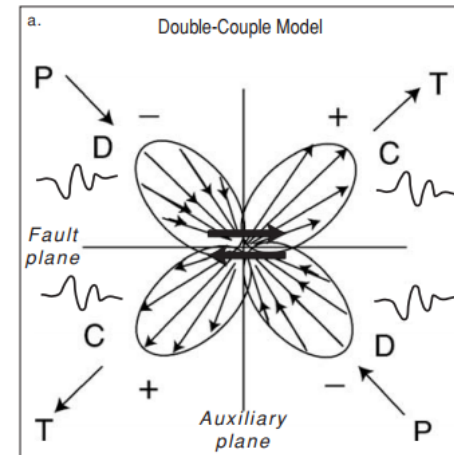
# In Situ Stress Measurement Method

## Indirect method – Focal Mechanism (also available at ch.5)



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- Earthquake Focal mechanism (= fault plane solution)
  - Two orthogonal nodal planes = fault plane + auxiliary plane
  - 90 degrees ambiguity
- Definition:
  - P-axis: bisects dilatational quadrant
  - T-axis: bisects compressional quadrant
  - B-axis: orthogonal to P and T
- Usefulness:
  - 1) Style of faulting,
  - 2) approx. direction of the principal stress



**Figure 5.11.** (a) Schematic illustration of the radiation pattern and force-couple associated with earthquakes as the basis earthquake focal plane mechanisms. An east–west striking, vertical right-lateral strike slip fault intersecting a half space is shown. The polarity of the P-waves defines the compressional and dilatational quadrants. (b) Cross-sectional view of the nodal planes, radiation pattern and P- and T-axes associated with an east-dipping normal fault. The radiation pattern does not uniquely distinguish the fault plane from the auxiliary plane. (c) Lower hemisphere stereonet representation of the normal faulting focal mechanism.

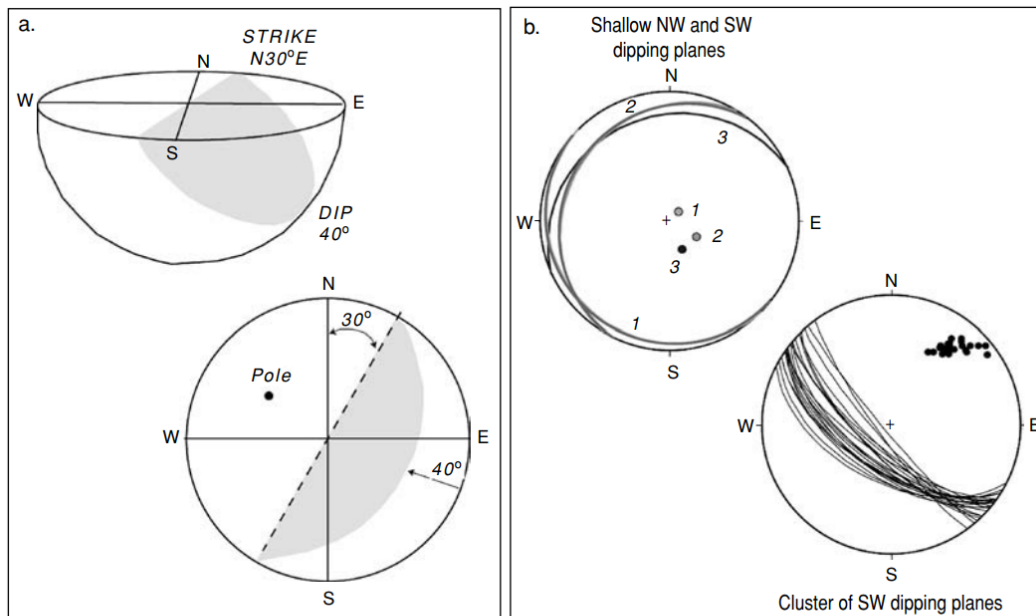
# Supplementary material

## Stereographic lower hemisphere projection



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- Stereographic lower hemisphere projection
  - show the trace of a fracture plane (where it intersects the lower half of the hemisphere) or
  - the intersection of fracture poles (normals to the fracture planes) and the hemisphere
- The circular diagrams (Figure 5.6b) used to represent such projections: stereonet



**Figure 5.6.** Illustration of the display of fracture and fault data using lower hemisphere stereographic projections. Either the intersection of the plane with the hemisphere can be shown or the pole to the plane. Planes which are sub-horizontal have poles that plot near the center of the stereonet whereas steeply dipping planes have poles which plot near the edge.



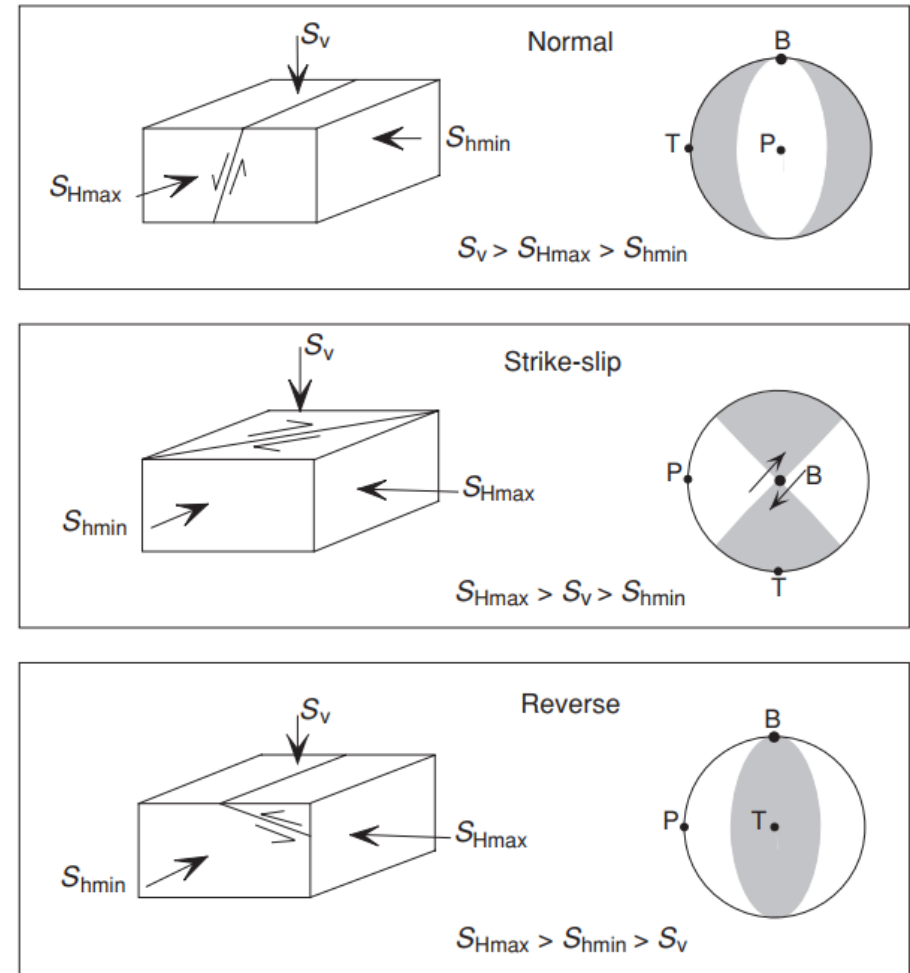
# In Situ Stress Measurement Method

## Indirect method – Focal Mechanism



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- Focal mechanisms of earthquake
  - Provides the orientation of principal stresses using Coulomb failure criterion
  - Relative magnitude of the three principal stress
  - 77% of World Stress Map data
  - Based on the analysis of observed seismic waveform (first motion of P-wave).  
Upward: compression,  
downward: dilational



**Figure 1.2.** E. M. Anderson's classification scheme for relative stress magnitudes in normal, strike-slip and reverse faulting regions. Earthquake focal mechanisms, the *beach balls* on the right, are explained in Chapter 5.



# In Situ Stress Measurement Method

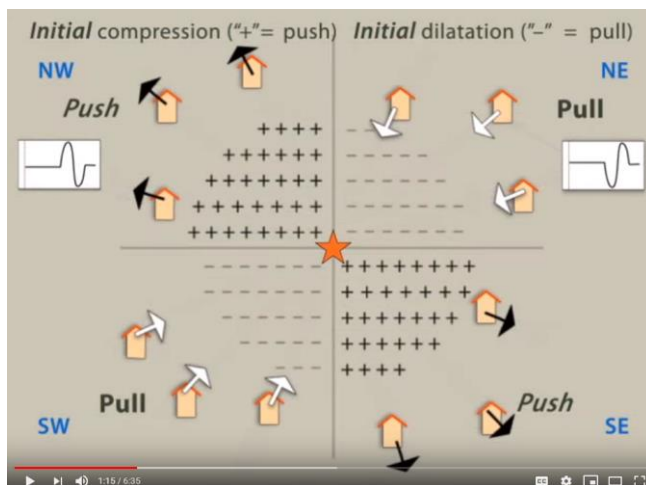
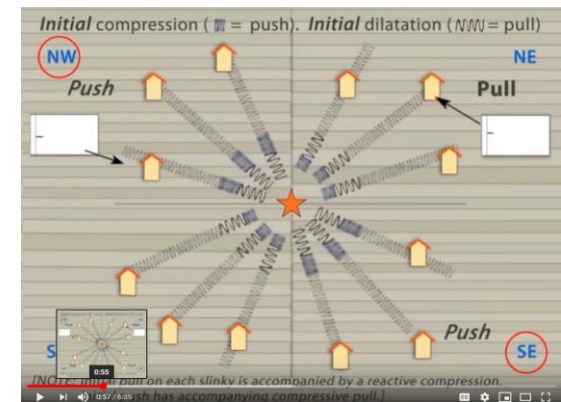
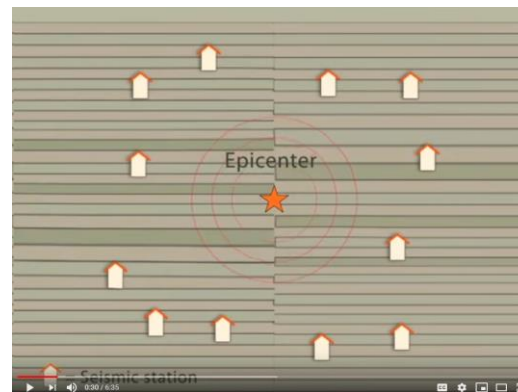
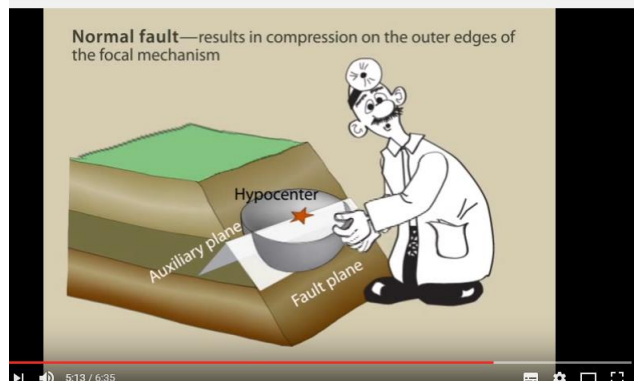
## Indirect method – Focal Mechanism



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- You tube video

- <https://www.youtube.com/watch?v=MomVOkyDdLo>



ambiguity

# In Situ Stress Measurement Method

## Indirect method – Focal Mechanism



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- Ambiguity can be clarified with additional knowledge, e.g., geology



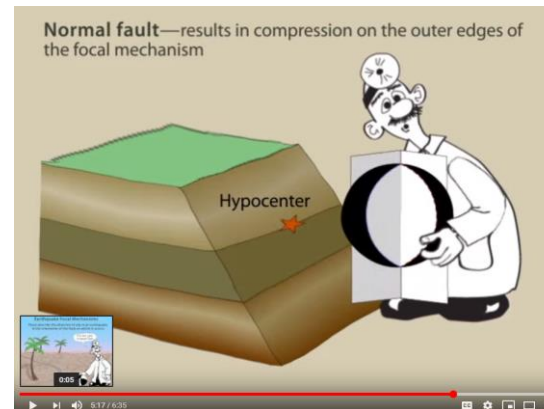
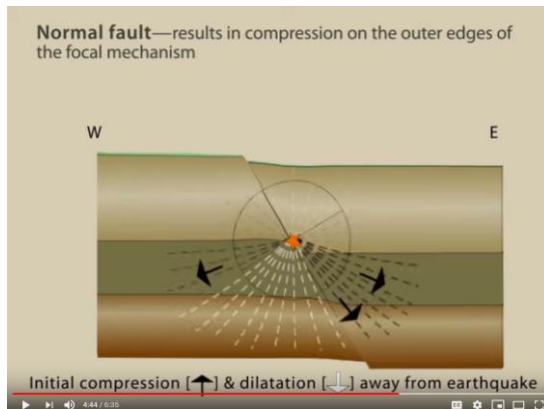
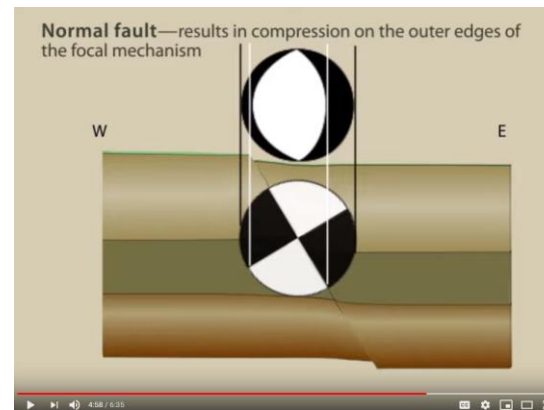
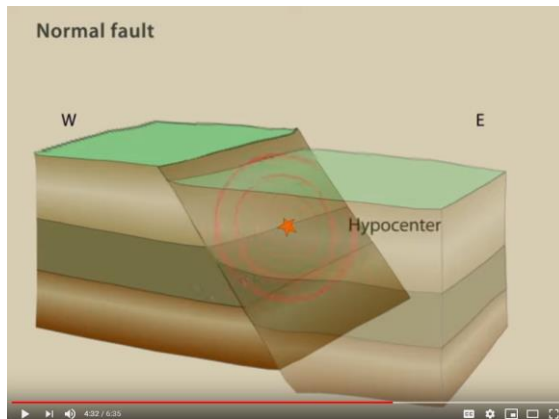
# In Situ Stress Measurement Method

## Indirect method – Focal Mechanism



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- Normal fault





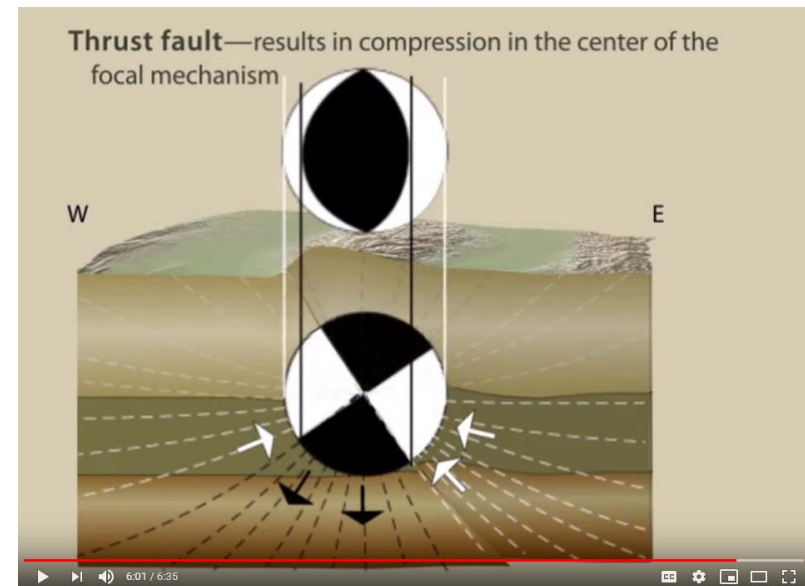
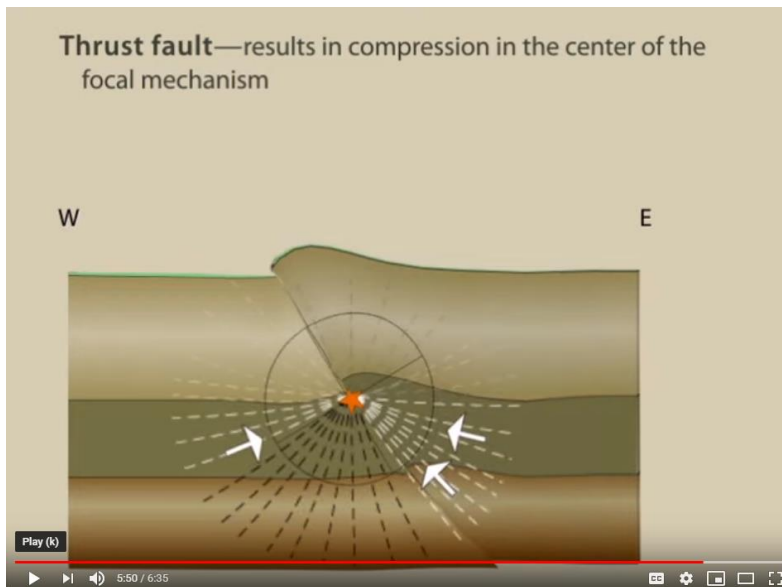
# In Situ Stress Measurement Method

## Indirect method – Focal Mechanism



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- Reverse Fault (Thrust Fault)



**Table 1.2.** *Summary of horizontal principal stress measurement methods*

## Stress orientation

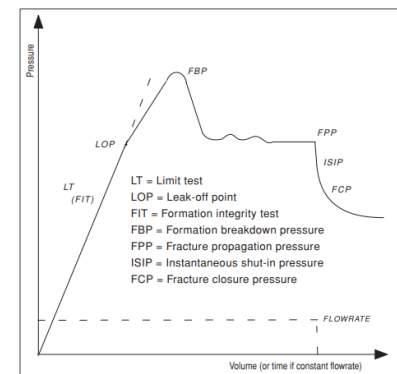
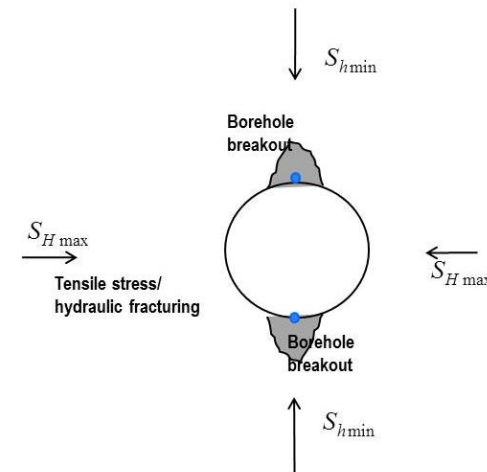
Stress-induced wellbore breakouts (Chapter 6)  
 Stress-induced tensile wall fractures (Chapter 6)  
 Hydraulic fracture orientations (Chapter 6)  
 Earthquake focal plane mechanisms (Chapter 5)  
 Shear velocity anisotropy (Chapter 8)

## Relative stress magnitude

Earthquake focal plane mechanisms (Chapter 5)

## Absolute stress magnitude

Hydraulic fracturing/leak-off tests (Chapter 7)  
 Modeling stress-induced wellbore breakouts (Chapter 7, 8)  
 Modeling stress-induced tensile wall fractures (Chapter 7, 8)  
 Modeling breakout rotations due to slip on faults (Chapter 7)



**Figure 7.2.** A schematic mini-frac or extended leak-off test showing pressure as a function of volume, or equivalently time (if the flow rate is constant). Modified after Gaarenstroom, Tromp *et al.* (1993). The significance of the various points indicated on the pressure record is discussed in the text.

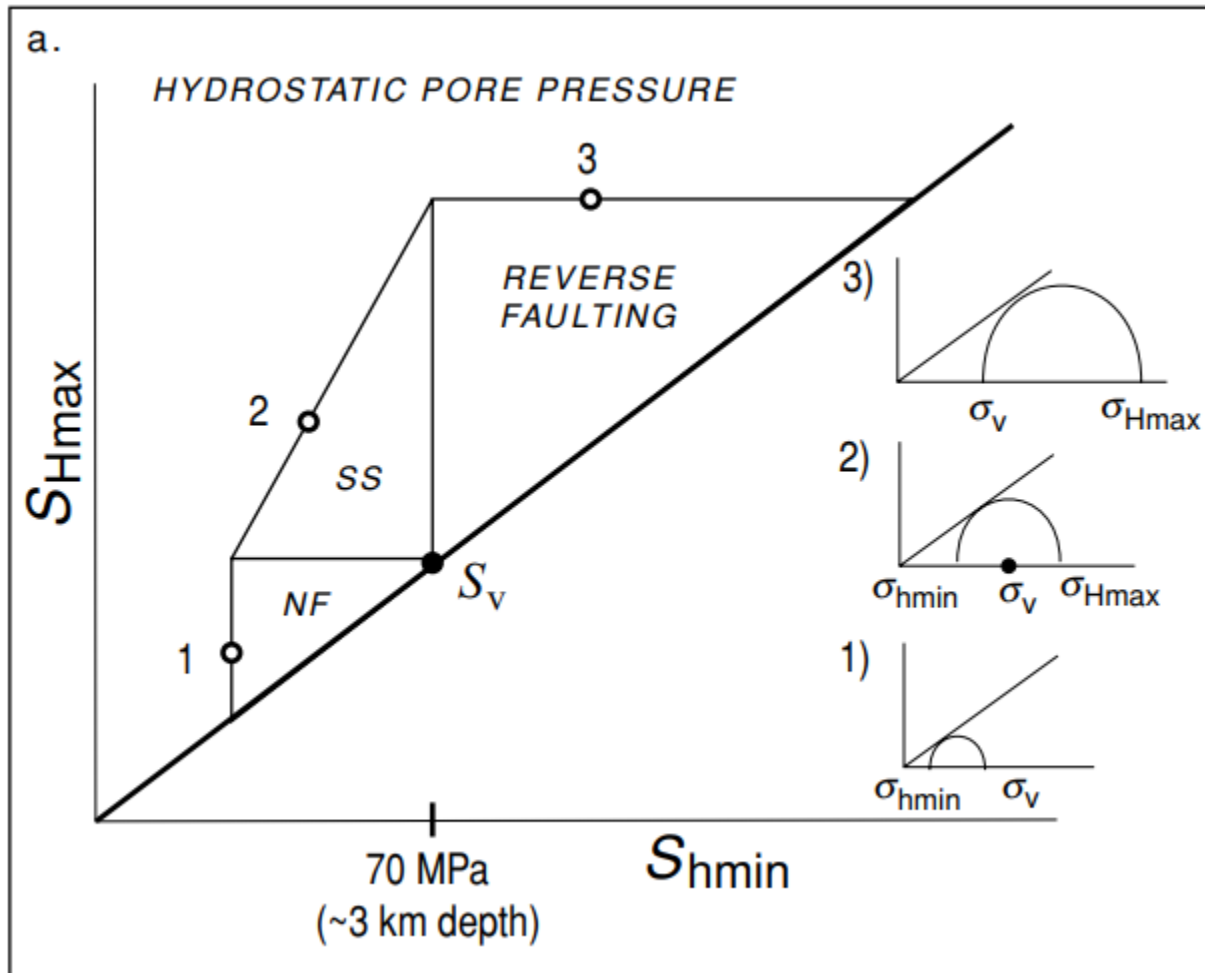
# Constraints of in situ stress

## Stress Polygon



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- Stress polygon – range of allowable stress magnitude



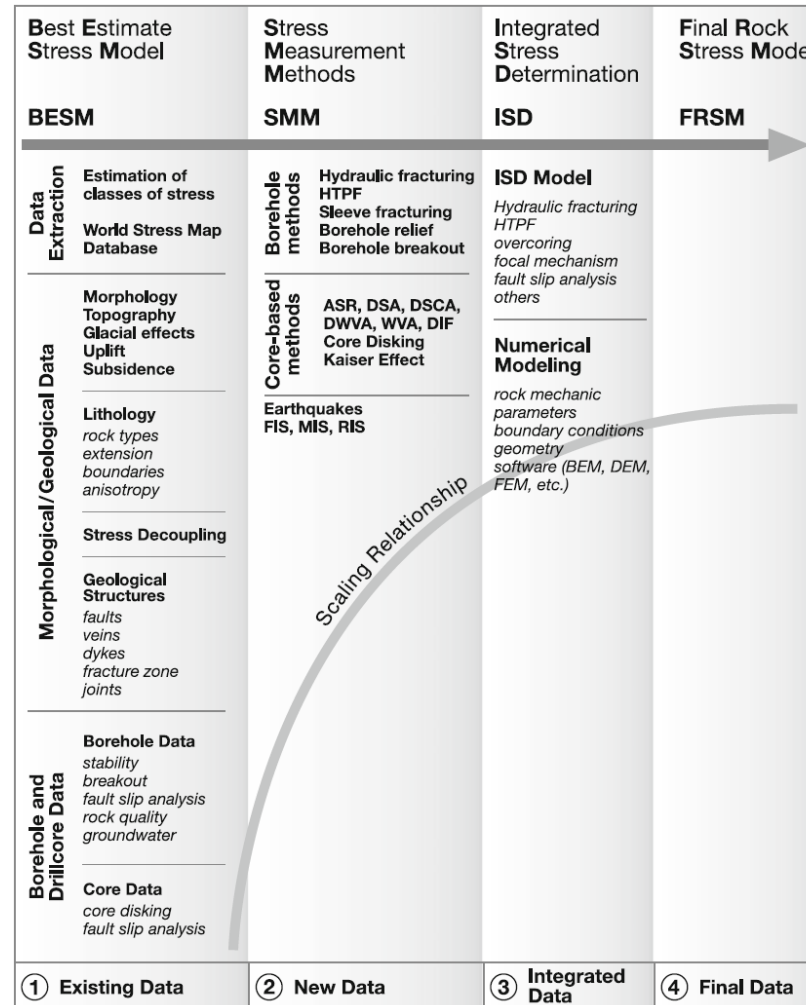
# In Situ Stress Measurement Method

## Integrated stress measurement



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- Multiple methods are often needed



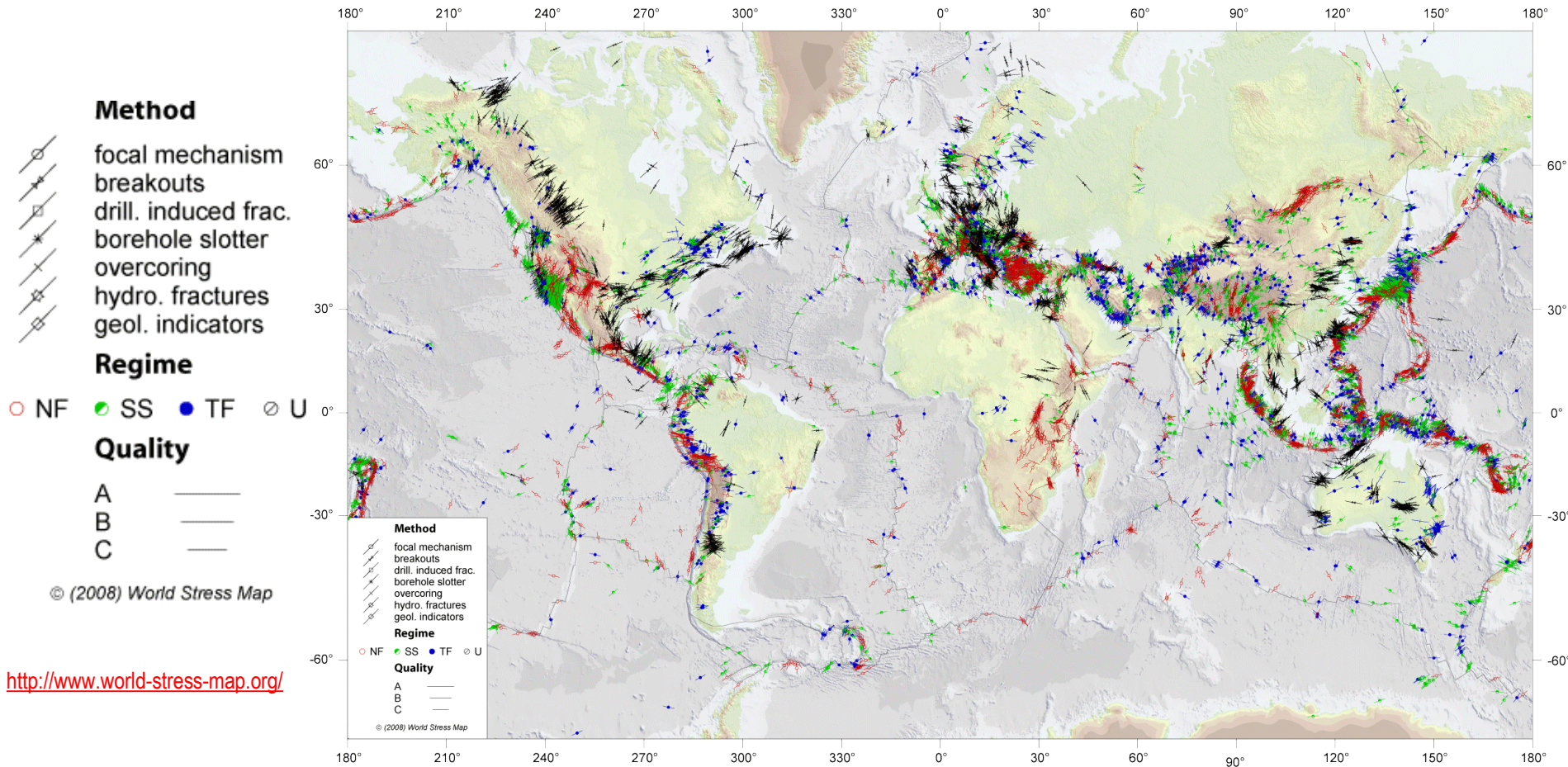


# World wide in situ stress data

## World stress map



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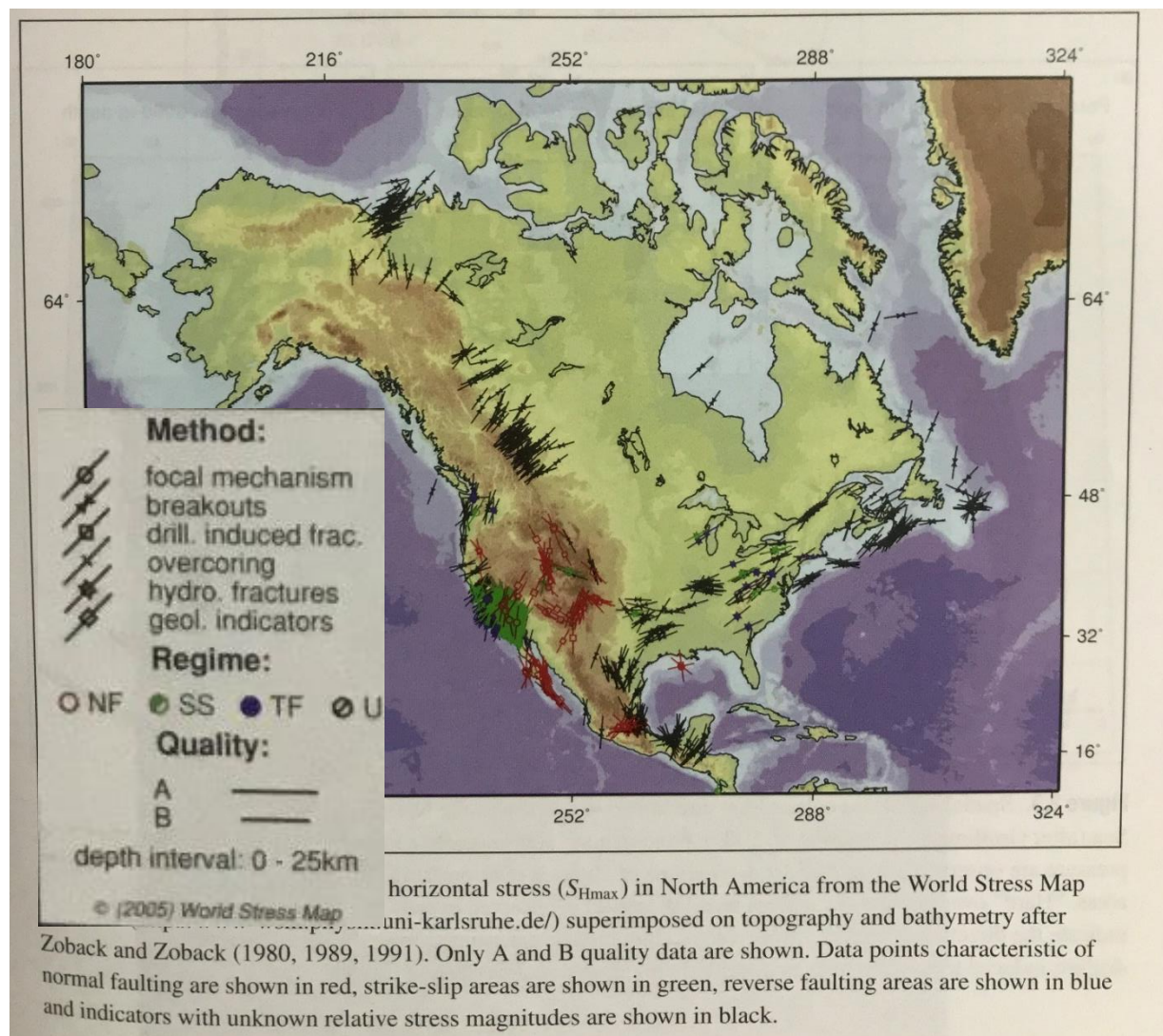


# World wide in situ stress data

## World stress map



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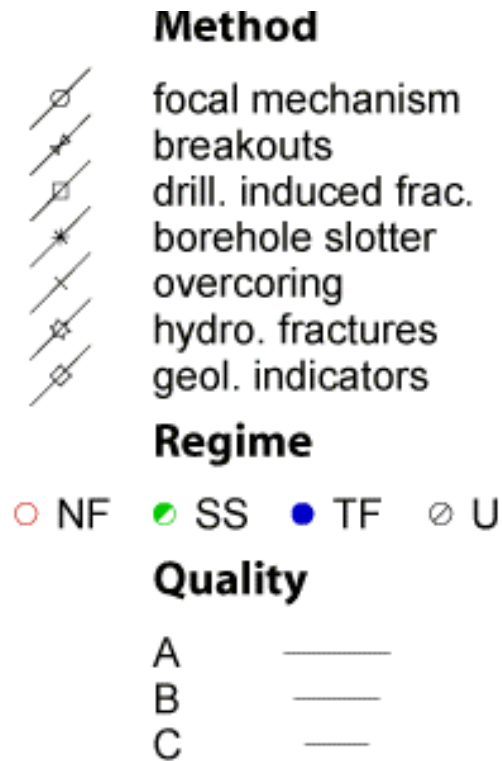
# World wide in situ stress data

## World stress map

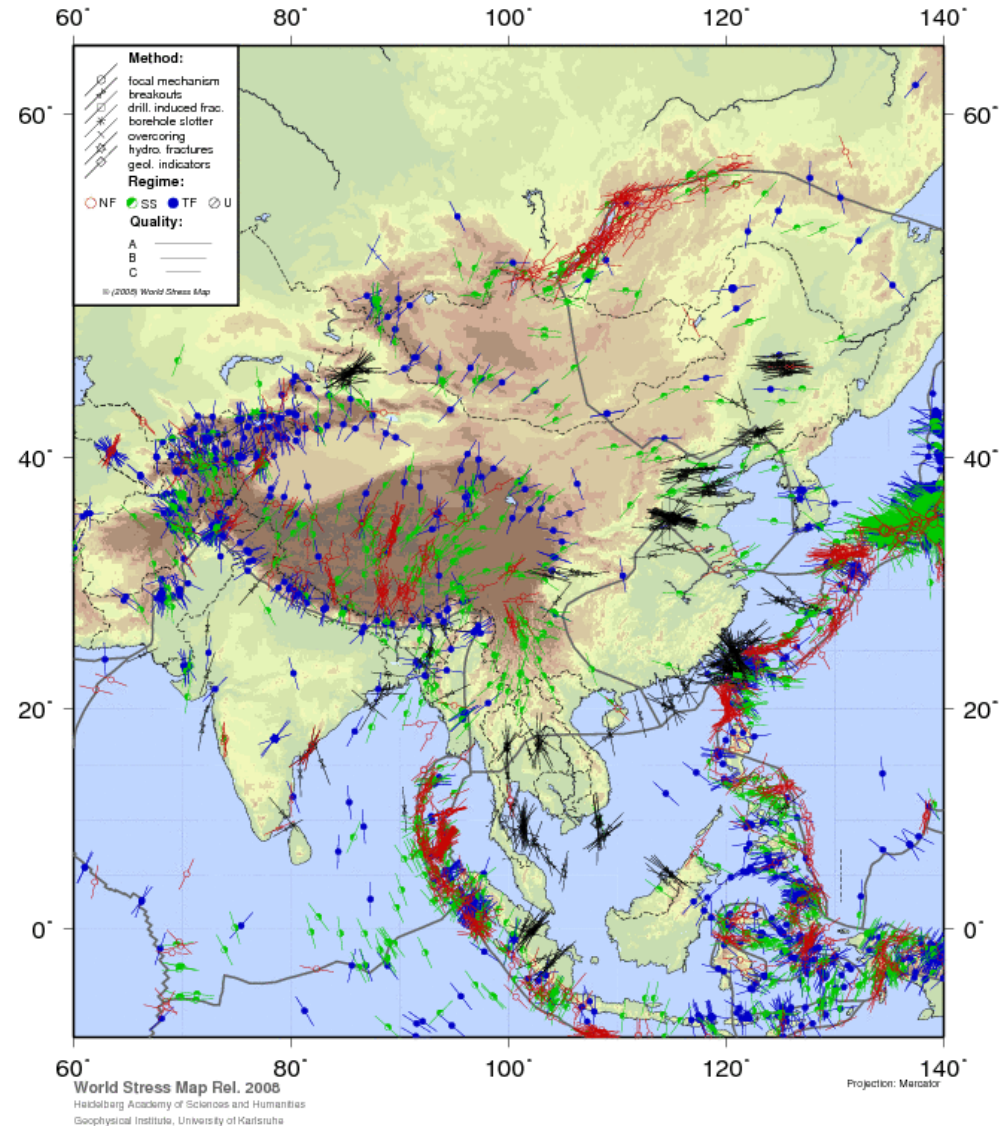


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- <http://www.world-stress-map.org/>



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# Summary

## Content of the course



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- Understand the importance of in situ stress for geomechanics
- Nature of in situ stress measurement;
- Direct method
  - ㄹ Flatjack method
  - ㄹ USBM overcoring method
  - ㄹ CSIRO (type) overcoring method
  - ㄹ Hydraulic fracturing method (mini-frac or micro-frac)
- Indicator method (indirect method)
  - ㄹ Borehole breakout
  - ㄹ Anelastic Strain Recovery (ASR)
  - ㄹ Kaiser effect
  - ㄹ Core diking
  - ㄹ Focal Mechanism of earthquake
- Estimation
  - ㄹ World Stress
  - ㄹ Integrated method