

Rock Mechanics & Experiment

암석역학 및 실험

Lecture 6. Fractures – geometrical and mechanical properties/Fractured rock

Lecture 6. 균열의 기하학적 및 역학적 성질/균열암반

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Seoul National University



- Fractures (Discontinuities)
- Geometrical properties of fractures
 - Orientation
 - Spacing (and frequency)
 - Rock Quality Designation (RQD)
 - Persistence, Roughness and Aperture
- Mechanical properties of fractures
 - Normal and shear stiffness
 - Strength

- Fractures (Discontinuities)

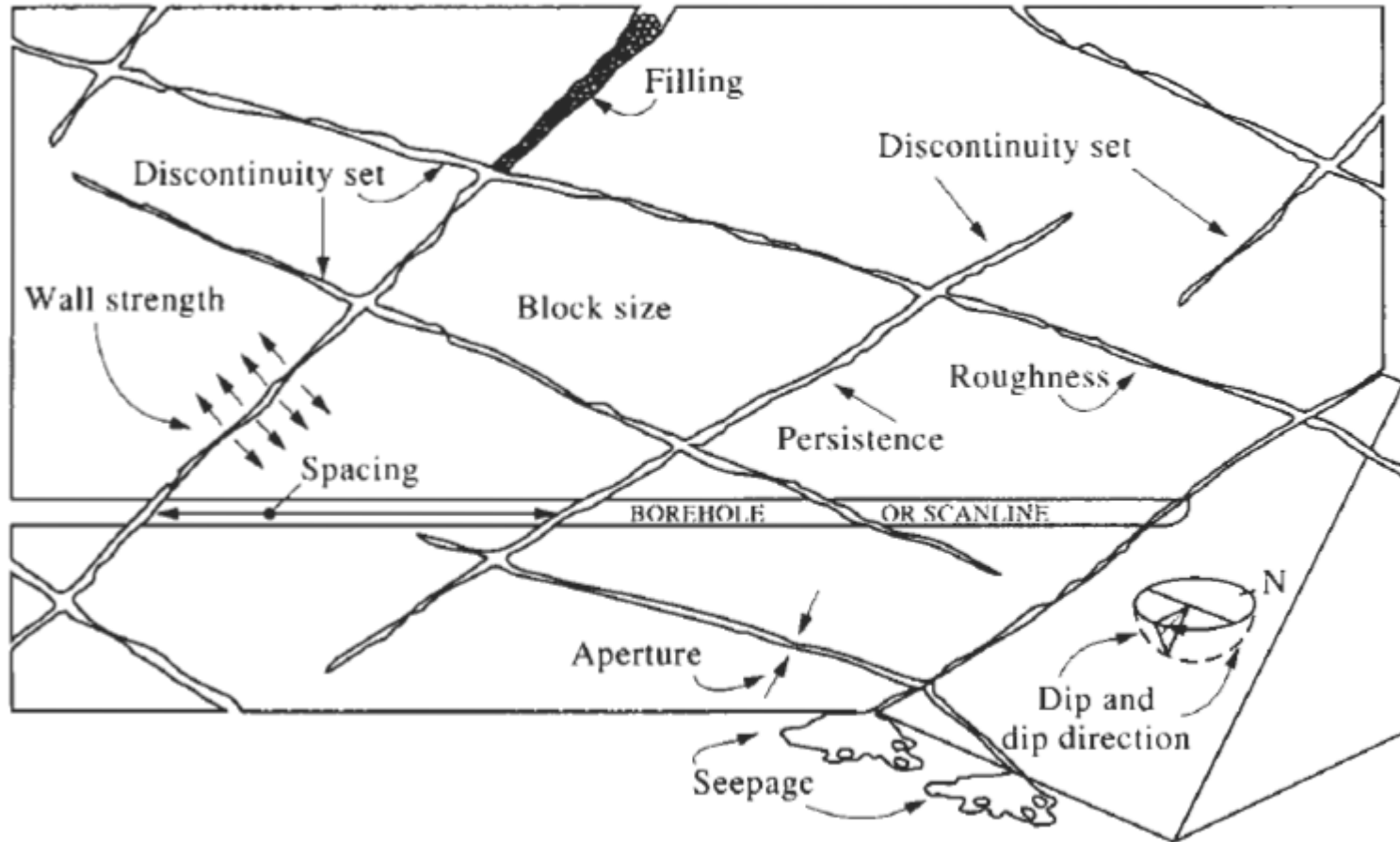


Figure 7.3 Schematic of the primary geometrical properties of discontinuities in rock (from Hudson, 1989).

Geometrical properties

Orientation - definition



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- Plane

- Dip (경사) & Dip direction (경사각)

- ↗ Dip angle: angle between the **steepest line** and horizontal plane

- ↗ Dip direction: bearing of this steepest line measured from North (clockwise)

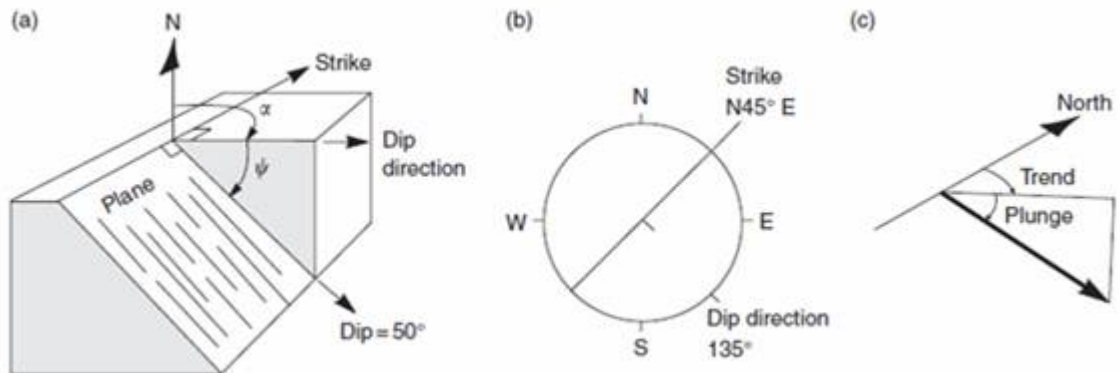
- ↗ Ex) 130/50 (dip direction/dip)

- Or Dip & Strike (주향)

- ↗ Ex) strike N40E, dip 50SE

- Line

- Trend & Plunge



Geometrical properties

Orientation – hemispherical projection

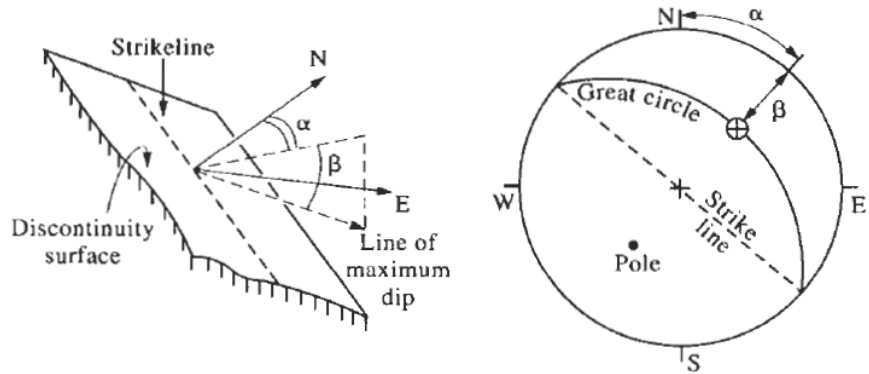
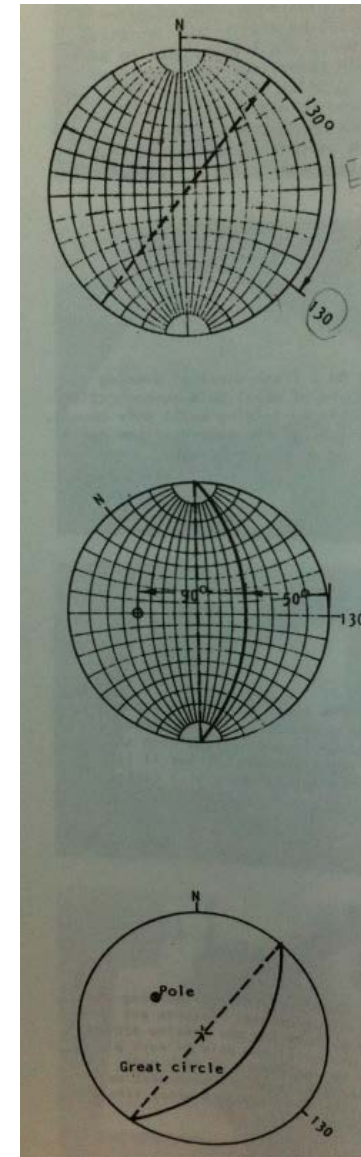
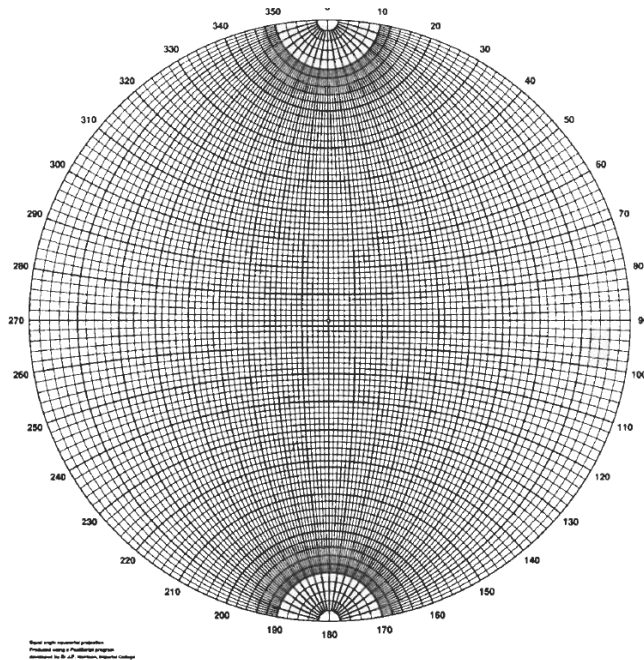


Figure 7.12 Discontinuity plane and the associated hemispherical projection.



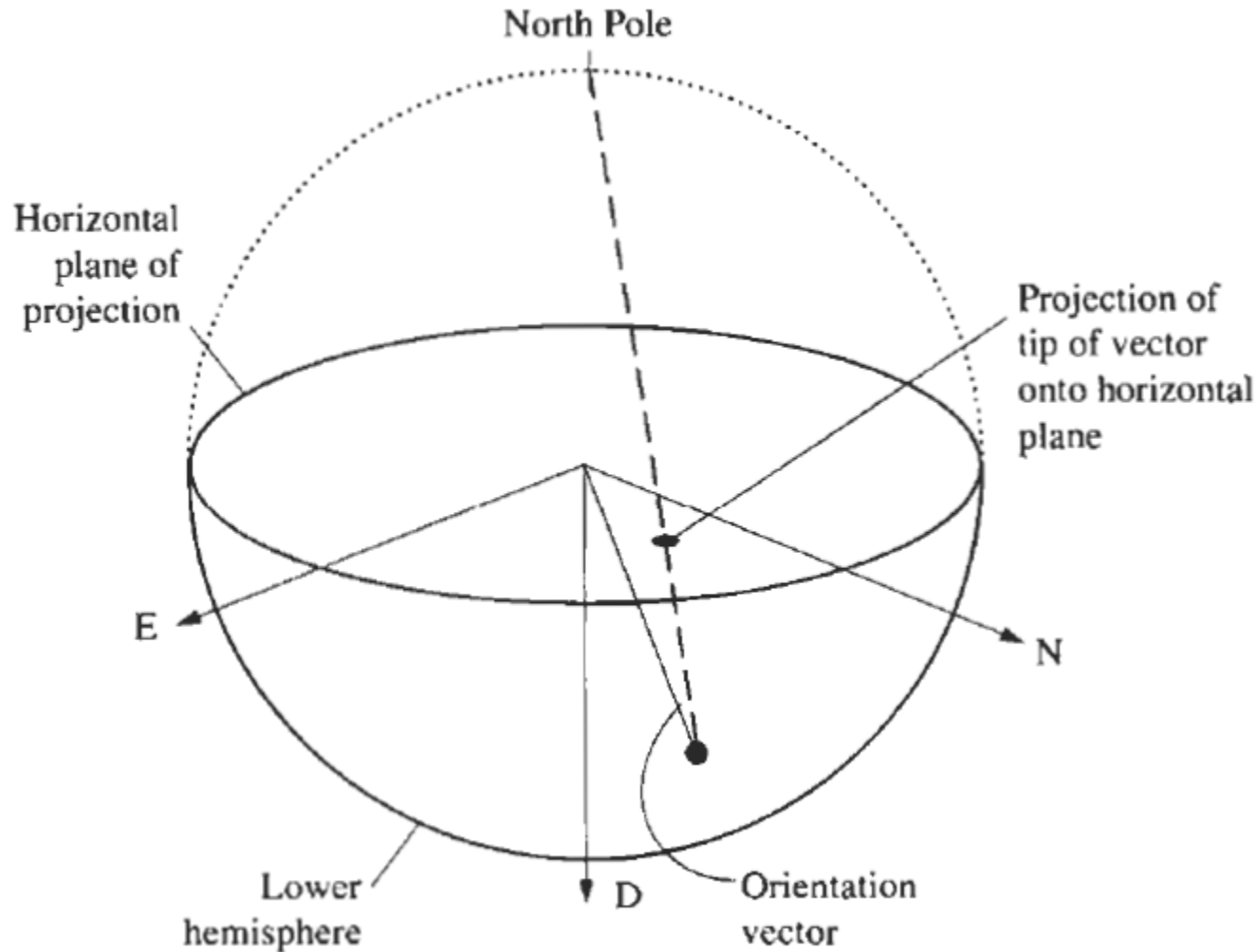
130/50의 예

Hemispherical projection

projection of a line onto two dimensions



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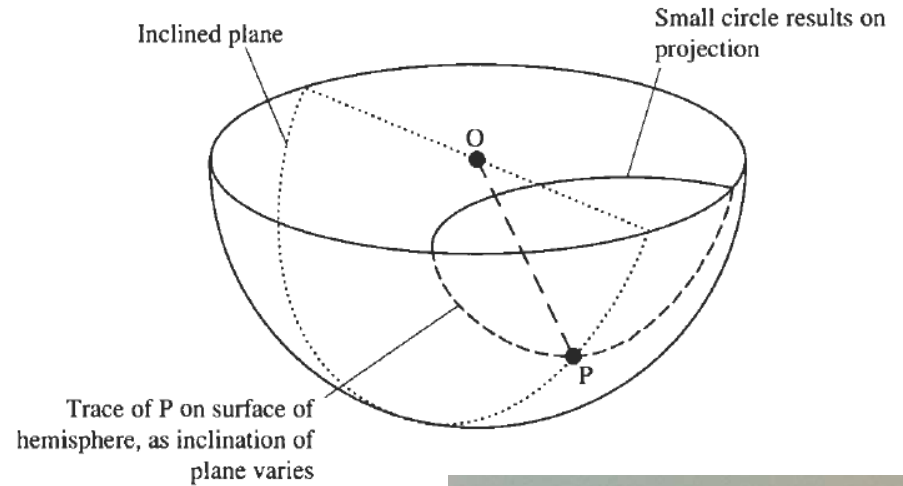
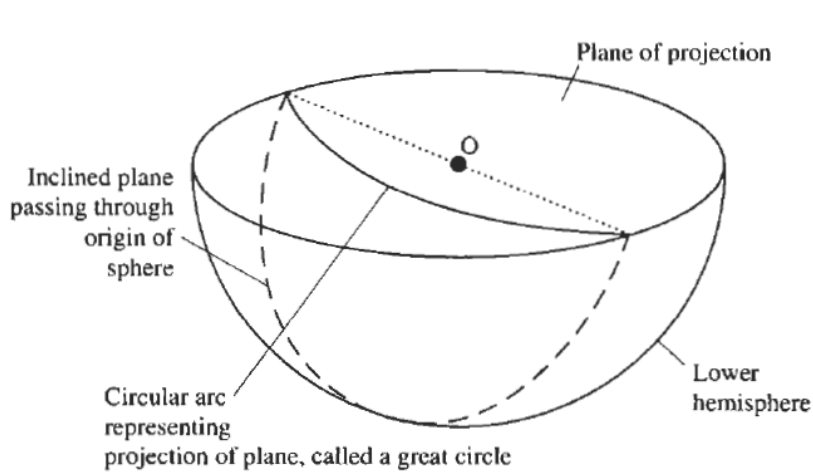


Hemispherical projection

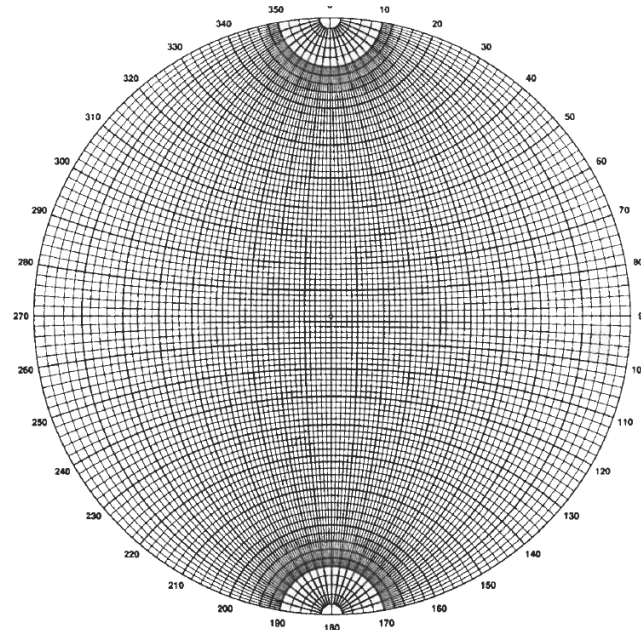
projection of a plane: great circles & small circles



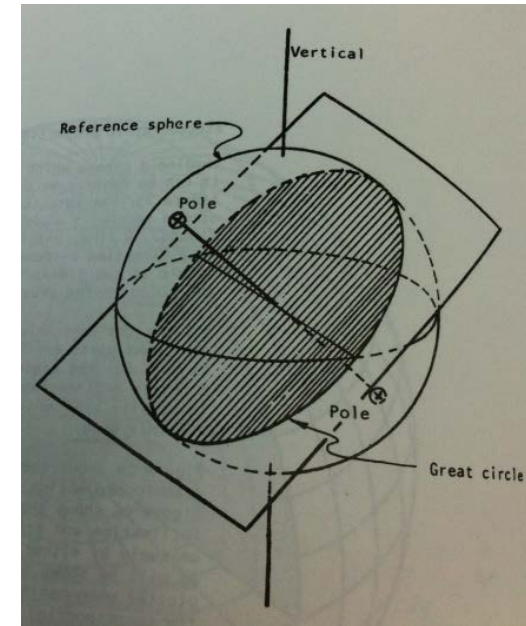
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Hemispherical projection net
(반구투영망)



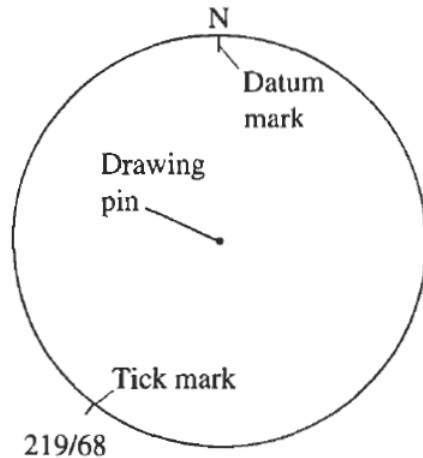
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Hemispherical projection plotting vectors

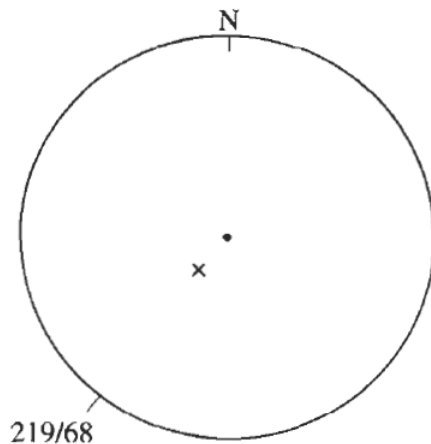
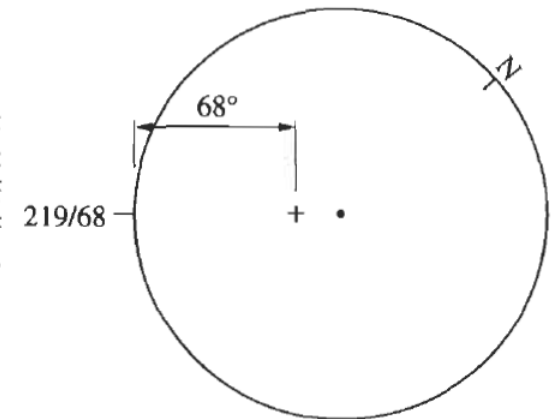


- Trend/plunge: 219/68



Mark, with a tick on the perimeter of the net at the correct azimuth, the vector to be plotted. Write on the projection the orientation. **Only write on the tracing paper, not the net.**

Rotate the tracing paper so that the tick is on the E-W line. Count in an amount equal to the dip of the vector. Mark the position of the vector. **Only write on the tracing paper, not the net.**



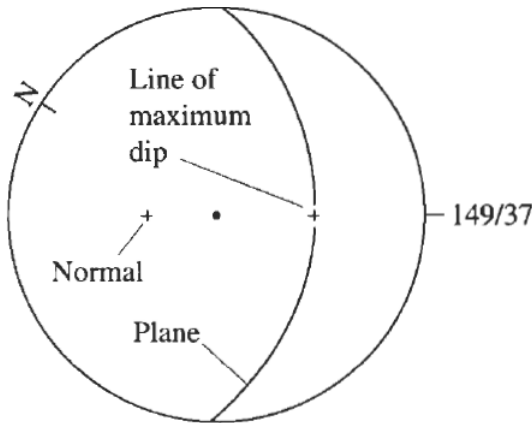
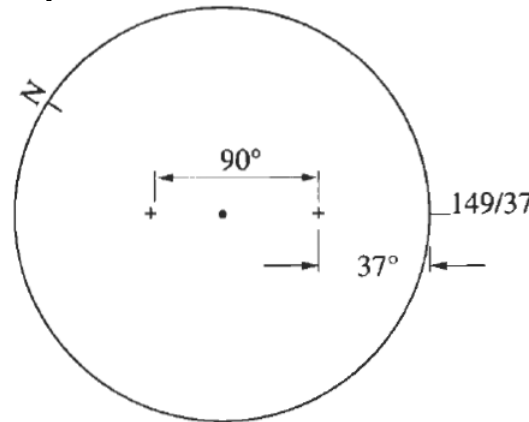
Rotate the tracing paper back to the datum: the position of the vector is now correct relative to north.

Hemispherical projection plotting planes



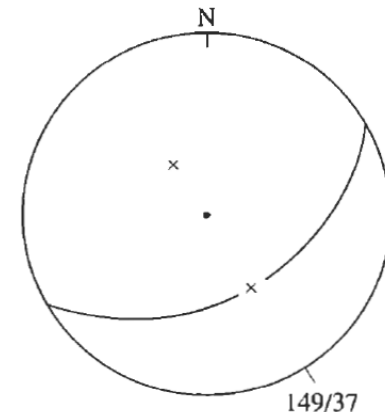
- Dip direction/dip 149/37

Rotate the tracing paper so that the tick lies on the E-W line, and count in an amount equal to the **dip amount** of the plane. Count a further 90° along the E-W line, and mark this new position.



Trace the **great circle** that passes over the first point: this represents **the plane**. The second point represents **the normal to the plane** (i.e. the vector that is perpendicular to the plane). The first point represents the **line of maximum dip**; the second point is termed the **pole**.

Rotate the tracing paper back to the datum: the positions of the plane and the normal are now correct relative to north.



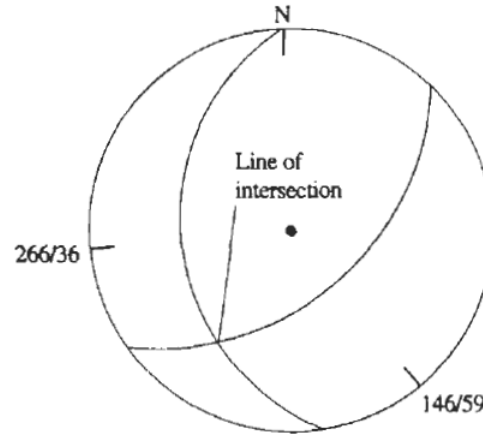
Hemispherical projection

Determining the line of two intersection of two planes

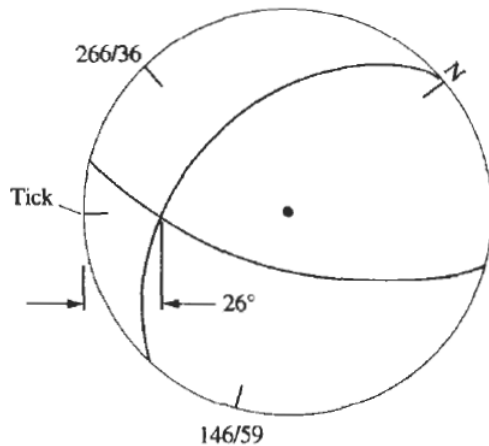


- 266/39 & 146/59

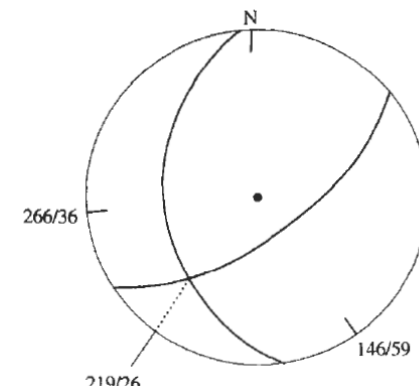
For any pair of planes, there is a line of intersection: it is where the two great circles cross. Finding its orientation is easy.



Rotate the tracing paper so that the intersection of the great circles lies on the E-W line. Mark the azimuth of this with a tick on the perimeter of the projection. Measure the plunge of the line by counting in from the perimeter, along the E-W line.



Rotate the tracing paper back to the datum, and measure the azimuth of the intersection. Thus we can see that the planes 266/36 and 146/59 have an intersection of 219/26. You should be able to measure graphically all such angles to the nearest degree.



Geometrical properties

Orientation



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- Fracture sets
 - Fractures with similar orientations are identified as a 'set'

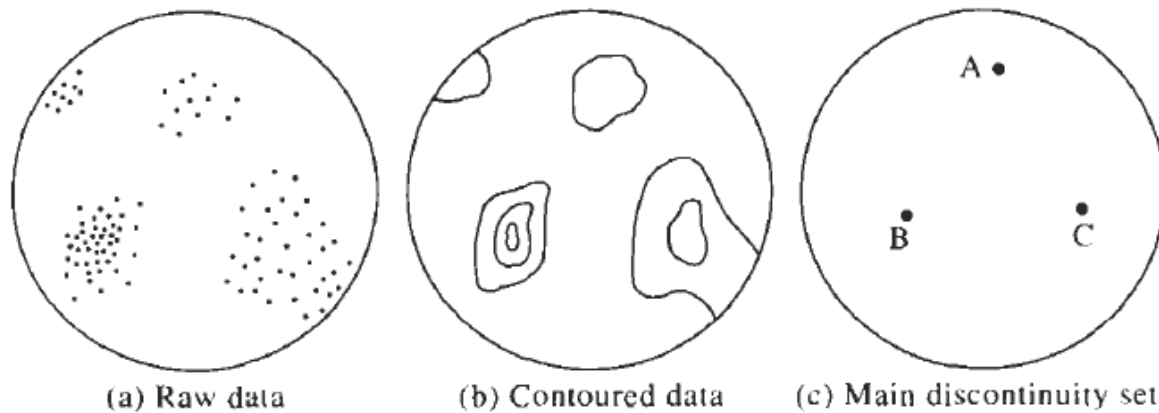


Figure 7.13 Discontinuity orientation data plotted on the lower-hemispherical projection.

Geometrical properties

Spacing (frequency)



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- Spacing
 - Distance between adjacent fracture intersections with the measuring scanline
- Frequency
 - Number of fractures per unit distance (reciprocal of spacing)
 - P_{10} : number of fracture per unit length of scan line
- Intensity
 - 2D, P_{21} : length of fracture traces per unit area of sampling surface
 - 3D, P_{32} : area of fractures per unit volume of rock mass

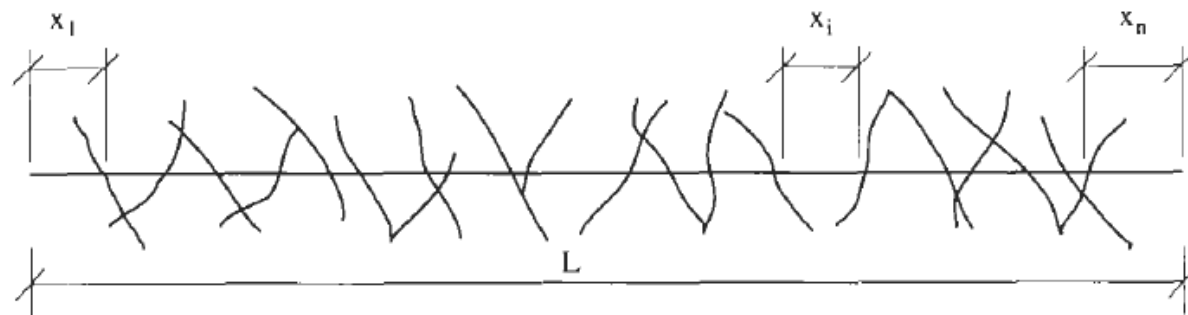


Figure 7.4 Quantifying discontinuity occurrence along a sampling line.

Geometrical properties

Spacing (frequency)



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- Geological Survey

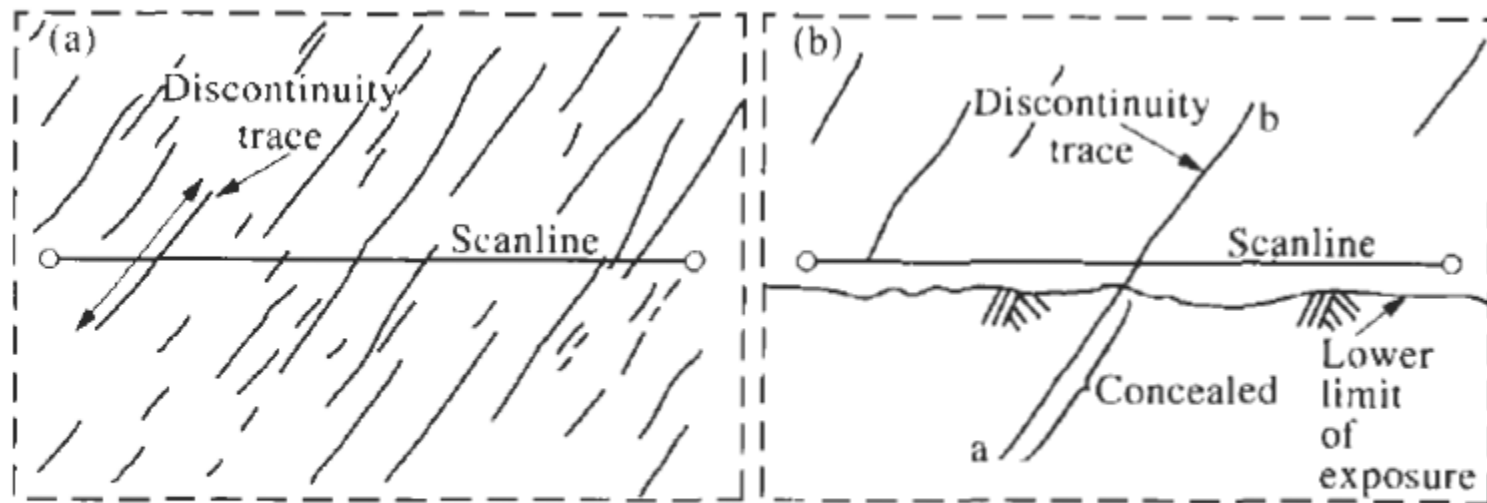


Figure 7.15 Diagrammatic representation of discontinuity traces intersecting a scanline set up on a rock face.

- Rock Quality Designation (RQD)
 - Percentage of the sampling line (cores) consisting of spacing values greater than 10 cm.

$$RQD = 100 \sum_{i=1}^n \frac{x_i}{L} \%$$

where x_i = spacing values greater than 0.1 m, and n is the number of these intersected by a borehole core or scanline of length L .

Geometrical properties

Persistence, roughness and aperture



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- Persistence (연속성)
 - Lateral extent of a discontinuity plane
 - Dimension (length, area) of fracture + connectivity
 - Dimension
 - ↗ Trace length
 - ↗ Area
- Aperture
 - The distance between adjacent walls of a discontinuity. Openness of the discontinuity
 - Important for hydraulic and mechanical behavior
 - Has distribution

Geometrical properties

Discrete Fracture Network (DFN)



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- Measurement of Aperture
 - Direct measurement
 - ↗ Insertion of feeler gauge (steel with different thickness)
 - ↗ Impression packer
 - ↗ Borehole camera
 - Indirect measurement
 - ↗ Back calculation from the flow rates



$$Q = \frac{\rho_w g e^3}{12\mu} \frac{\partial h}{\partial x}$$

unknown

$$e = \left(\frac{12\mu Q}{\rho_w g} / \frac{\partial h}{\partial x} \right)^{1/3}$$

Geometrical properties

Discrete Fracture Network (DFN)



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- Geometrical properties of fractures
 - Orientation (dip and dip direction), size (trace length in 2D), intensity (spacing in parallel infinite fracture), location, aperture, roughness
- Characterization method
 - Exposed rock faces
 - ↗ scanline sampling: line-based sample, use measuring tape (줄자).
 - ↗ Window sampling: area-based sample, rectangle of measuring tapes
 - Borehole sampling
- We then need to construct a geometric model of fractured rock – deterministic or stochastic generation of fractures
 - Monte Carlo Simulation

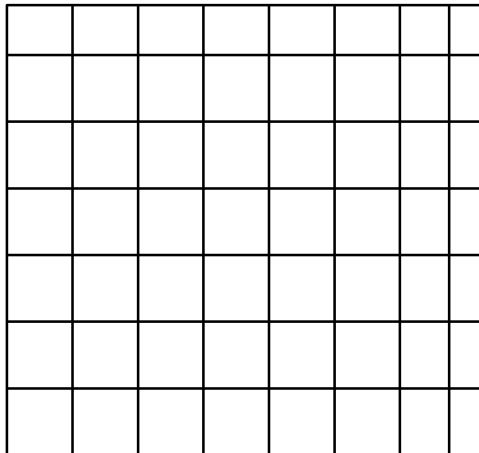
Geometrical properties

Discrete Fracture Network (DFN)

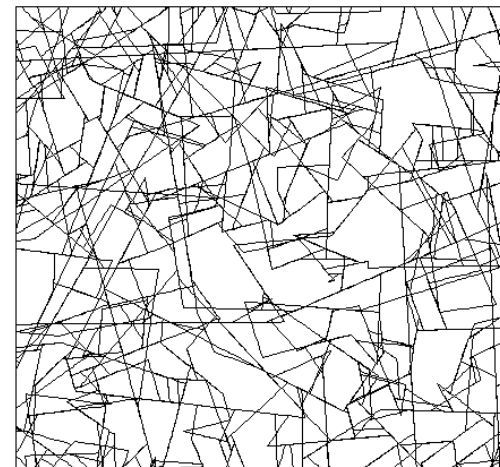


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- Geometrical models based on the characterisation



Idealized regular fracture model



Discrete Fracture Network
(암반균열망)

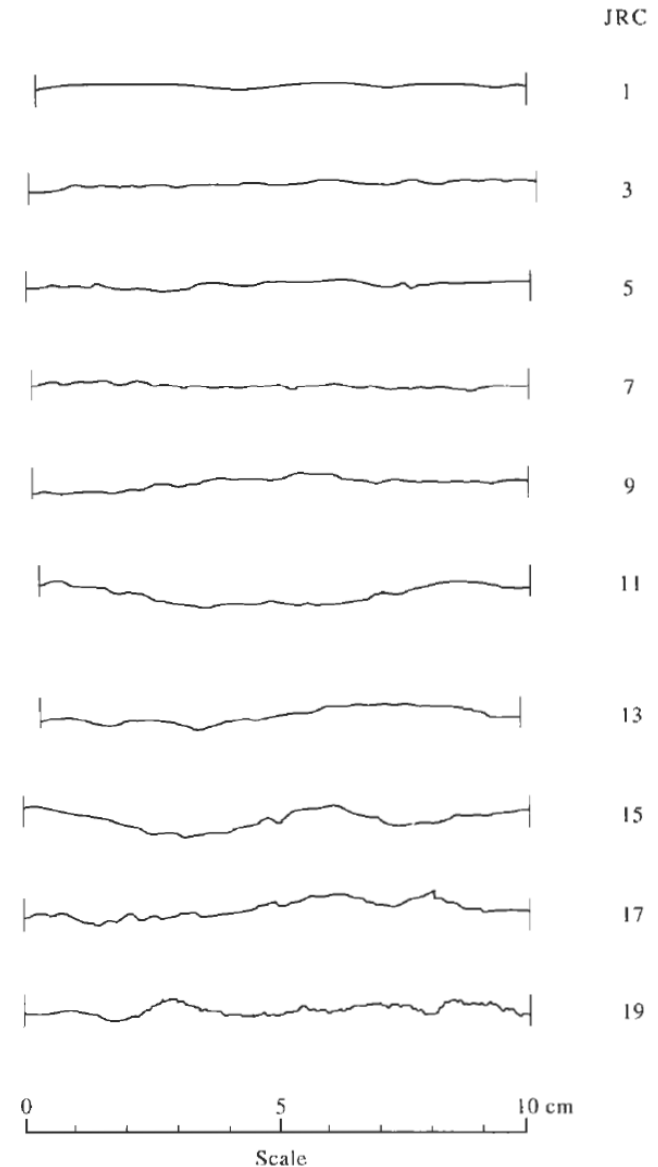
Geometrical properties

Persistence, roughness and aperture

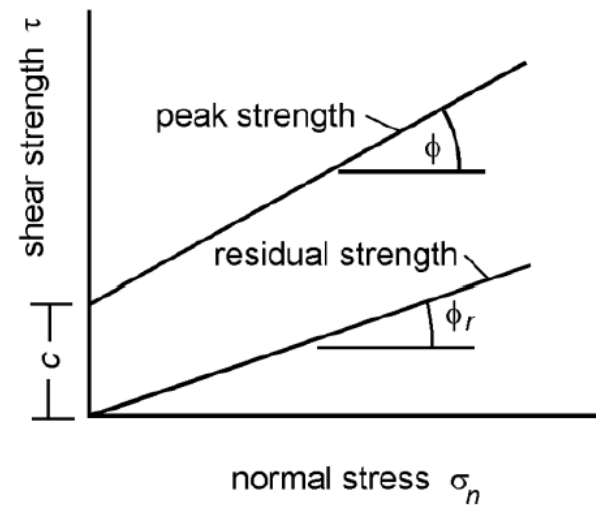
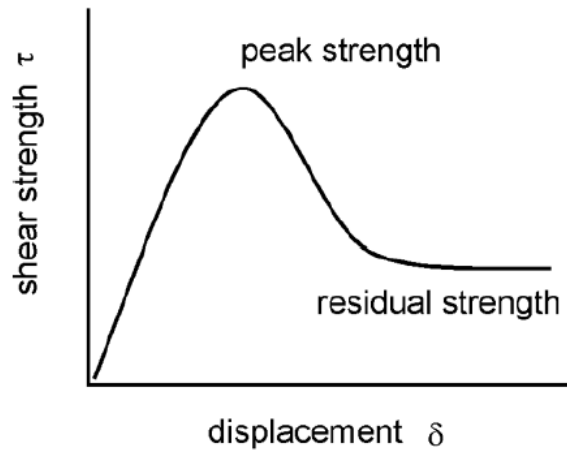
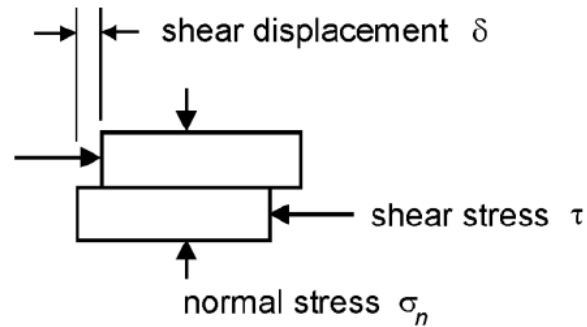


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- Roughness (거칠기)
 - Deviation of a discontinuity surface from perfect planarity
- Joint Roughness Coefficient (JRC)
 - Compare the profiles of discontinuity surface with standard roughness profile and assign numerical values.
 - Developed by Nick Barton (1977)
 - Not robust way but practically very important for mechanical and hydraulic properties of fractures



- Shear testing of fractures



Mechanical properties

Stiffness

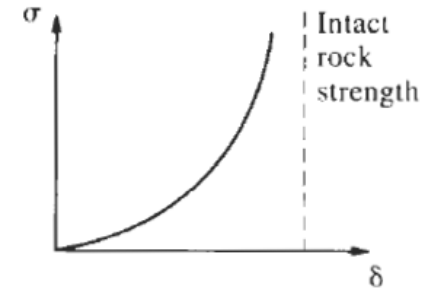
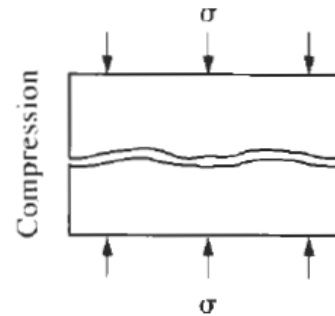


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- Normal stiffness

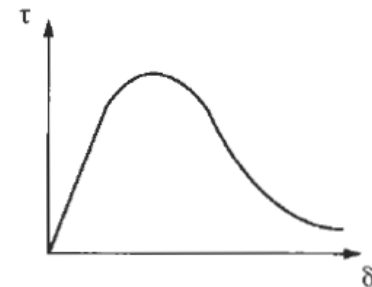
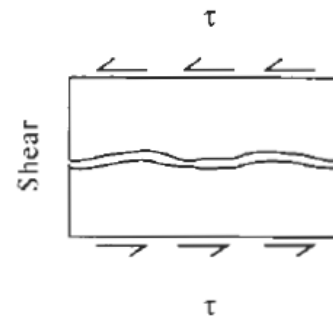
- Unit: Stress/length (MPa/m)
- Linear model $\sigma_n = K_n \delta_n$
- Non-linear model

$$\delta_n = \frac{\sigma_n}{c + d\sigma_n}$$



- Shear stiffness

- Unit: Stress/length (MPa/m)
- Linear model $\sigma_s = K_s \delta_s$



- Non-linear model: e.g., Barton's equation (strength + full path)

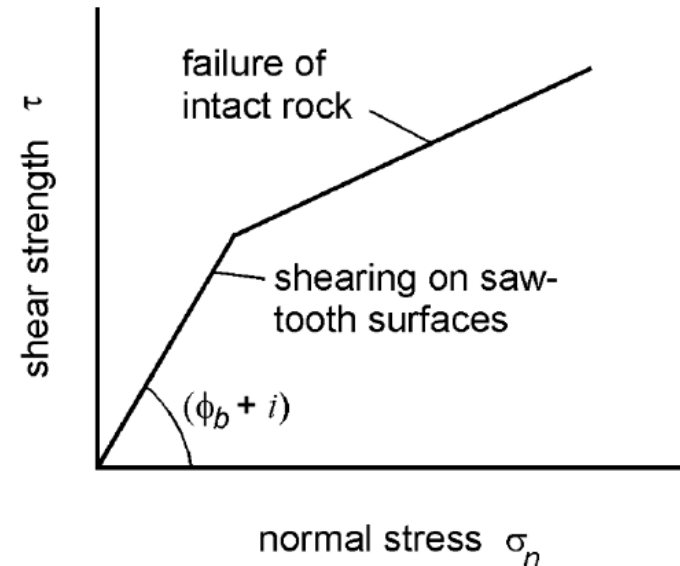
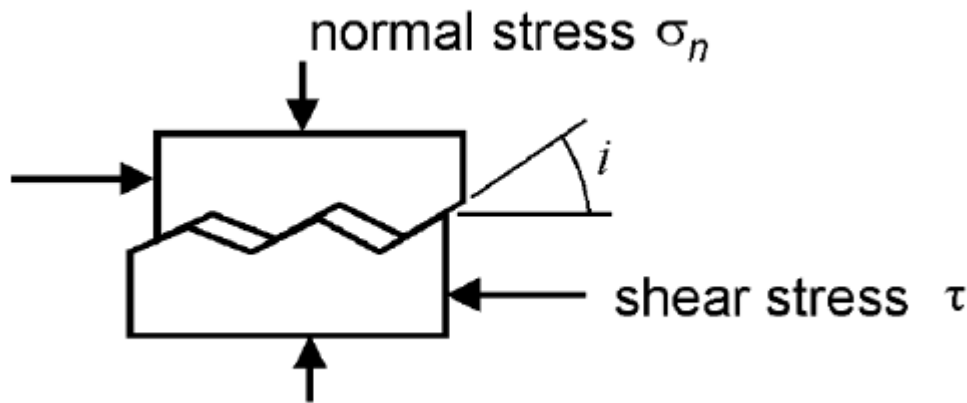
Mechanical properties

Strength



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- Patton's law (1966) for fracture friction angle



$$\tau = \sigma_n \tan(\phi_b + i)$$

where ϕ_b is the basic friction angle of the surface and i is the angle of the saw-tooth face.

Mechanical properties

Strength



- Barton's equation (1977)

$$\tau = \sigma_n \tan [JRC \log_{10} (JCS / \sigma_n) + \phi_r]$$

τ : Shear Strength of a fracture

JRC: Joint Roughness Coefficient

JCS :Joint Wall Compressive Strength

ϕ_r : residual friction angle

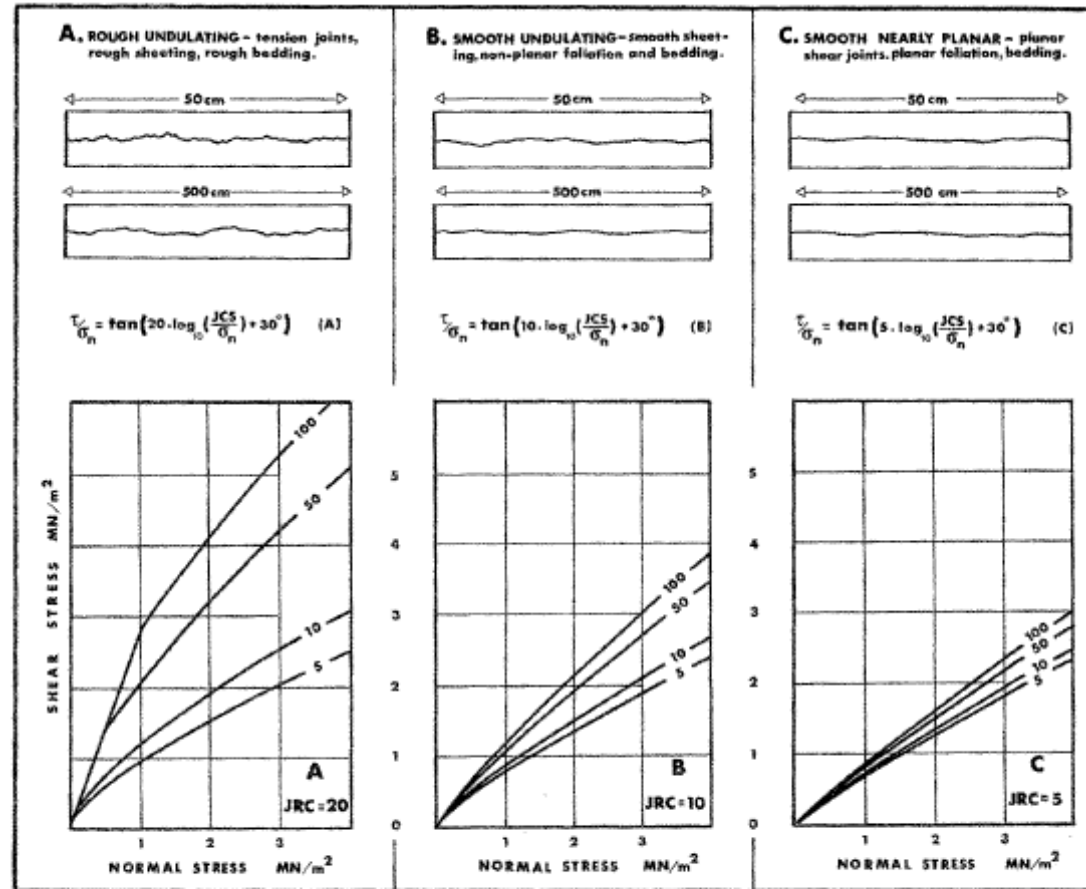


Fig. 1. Empirical law of friction in graphical form. Each curve is numbered with the appropriate JCS value (units of MN/m²). The roughness profiles are intended as an approximate guide to the appropriate JRC values 20, 10 and 5. Completely smooth planar joints have JRC = 0

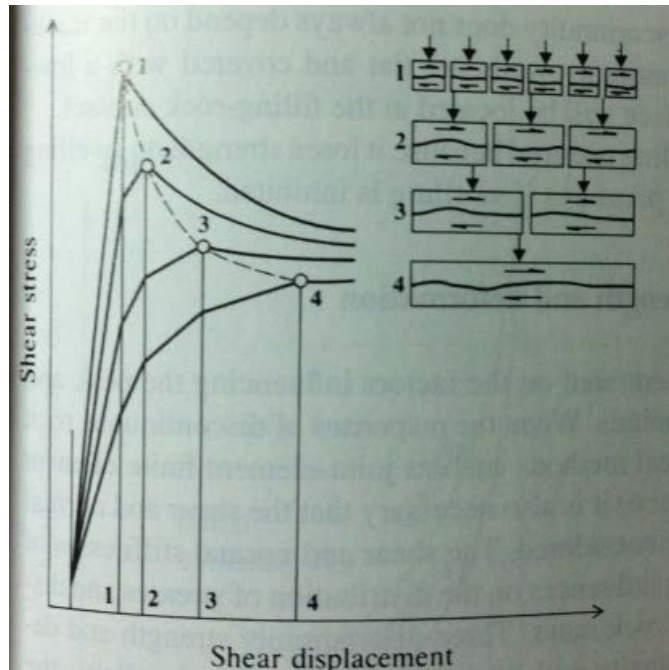
Mechanical properties

Strength



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- Size dependent behavior of a fracture



$$JRC_n = JRC_o \left(\frac{L_n}{L_o} \right)^{-0.02JRC_o} \quad (8)$$

where JRC_o , and L_o (length) refer to 100 mm laboratory scale samples and JRC_n , and L_n refer to in situ block sizes.

Mechanical properties Strength

Table 1: Shear strength of filled discontinuities and filling materials (After Barton 1974)

Rock	Description	Peak c' (MPa)	Peak ϕ°	Residual c' (MPa)	Residual ϕ°
Basalt	Clayey basaltic breccia, wide variation from clay to basalt content	0.24	42		
Bentonite	Bentonite seam in chalk Thin layers Triaxial tests	0.015 0.09-0.12 0.06-0.1	7.5 12-17 9-13		
Bentonitic shale	Triaxial tests Direct shear tests	0-0.27	8.5-29	0.03	8.5
Clays	Over-consolidated, slips, joints and minor shears	0-0.18	12-18.5	0-0.003	10.5-16
Clay shale	Triaxial tests Stratification surfaces	0.06	32	0	19-25
Coal measure rocks	Clay mylonite seams, 10 to 25 mm	0.012	16	0	11-11.5
Dolomite	Altered shale bed, \pm 150 mm thick	0.04	1(5)	0.02	17
Diorite, granodiorite and porphyry	Clay gouge (2% clay, PI = 17%)	0	26.5		
Granite	Clay filled faults Sandy loam fault filling Tectonic shear zone, schistose and broken granites, disintegrated rock and gouge	0-0.1 0.05 0.24	24-45 40 42		
Greywacke	1-2 mm clay in bedding planes			0	21
Limestone	6 mm clay layer 10-20 mm clay fillings <1 mm clay filling	0.1 0.05-0.2	13-14 17-21	0	13
Limestone, marl and lignites	Interbedded lignite layers Lignite/marl contact	0.08 0.1	38 10		
Limestone	Marlaceous joints, 20 mm thick	0	25	0	15-24
Lignite	Layer between lignite and clay	0.014-.03	15-17.5		
Montmorillonite Bentonite clay	80 mm seams of bentonite (montmorillonite) clay in chalk	0.36 0.016-.02	14 7.5-11.5	0.08	11
Schists, quartzites and siliceous schists	100-15- mm thick clay filling Stratification with thin clay Stratification with thick clay	0.03-0.08 0.61-0.74 0.38	32 41 31		
Slates	Finely laminated and altered	0.05	33		
Quartz / kaolin / pyrolusite	Remoulded triaxial tests	0.042-.09	36-38		

Mechanical properties

Strength - dilation



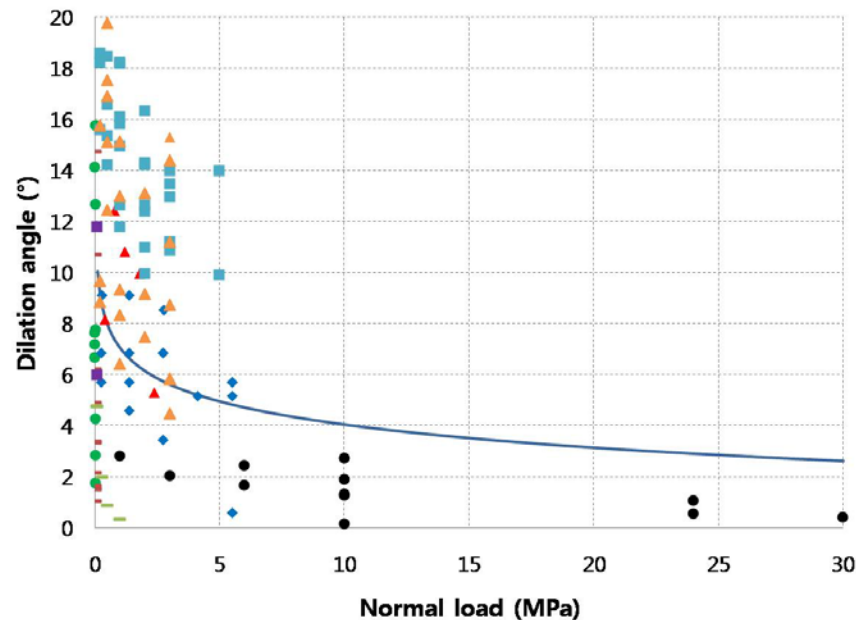
- Fracture dilation during fracture shearing
 - Great implication for geo-environmental engineering

$$\phi_{dilation} = \tan^{-1} \left(\frac{\text{normal dilation}}{\text{shear displacement}} \right)$$

$$\phi_{dilation} = \frac{1}{2} JRC \log \left(\frac{JCS}{\sigma_n} \right)$$



Shear
dilation
----->



- Forsmark
- Barton (1982) - granite
- Barton et al. (1985) - granite
- ▲ Homand et al. (2001) - replica
- Bandis et al. (1981) - sandstone
- Huang et al. (2002) - replica
- ◆ Wibowo et al. (1994) - replica
- Lee (1999) - Granite
- ▲ Lee (1999) - Marble
- Yeo et al. (1998) - replica

Mechanical properties

Strength-dilation

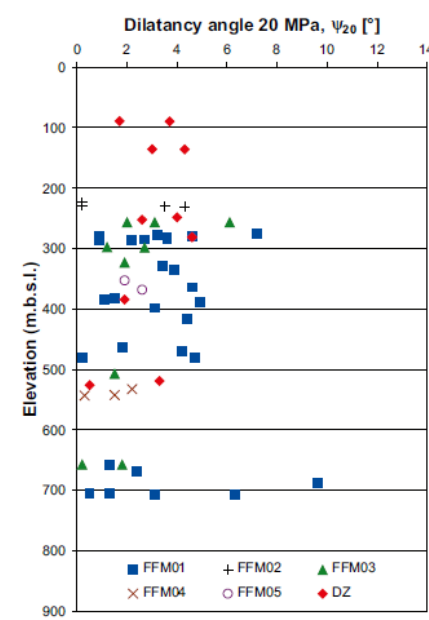
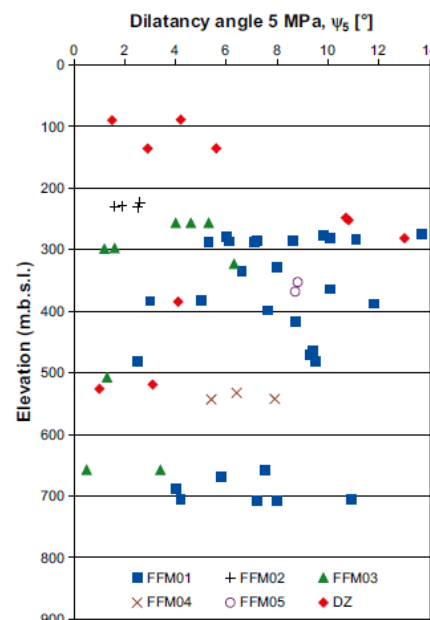
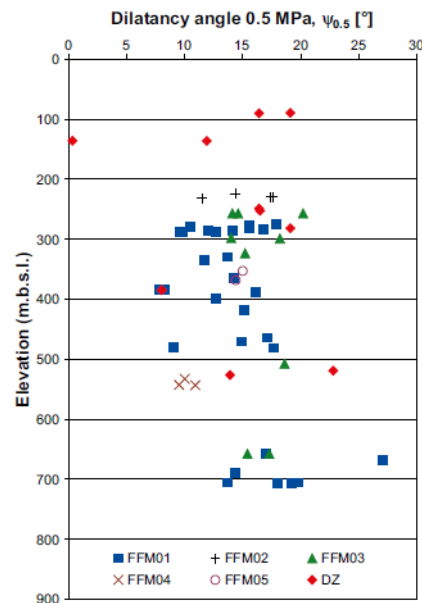


– Aperture change due to shear dilation (SKB, 2007)

 ≈ 3.2° dilation angle under 20 MPa

 ≈ Dilation angles from direct shear tests for the FFM01 fracture domain

Normal load (MPa)	Mean (°)	Std. dev. (°)	Minimum (°)	Maximum (°)	Uncertainty of mean (%)
0.5	14.6	4.1	7.8	27.1	±10.2
5	7.7	2.7	2.5	13.7	±12.8
20	3.2	2.1	0.2	9.6	±23.9

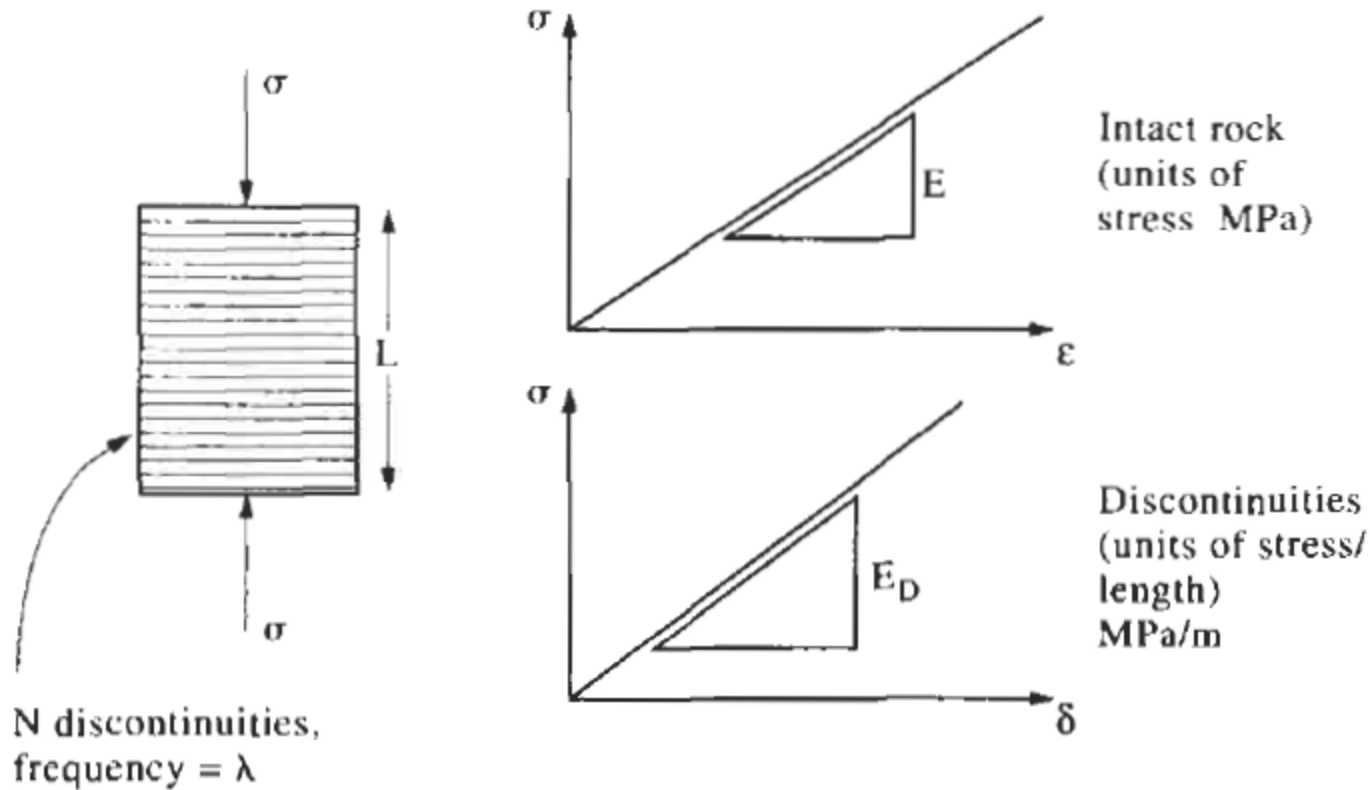


Fractured Rock Masses

Elastic modulus (deformation modulus)



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$$\delta = 1 / \lambda = \text{spacing of fractures}$$

Fractured Rock Masses

Elastic modulus (deformation modulus)



- Elastic (deformation) modulus of fractured rock mass

$$\frac{1}{E_m} = \left(\frac{1}{E_i} + \frac{1}{k_n \cdot S} \right) \quad \frac{1}{E_m} = \left(\frac{1}{E_i} + \frac{\lambda}{E_D} \right)$$

E_m : Elastic modulus of rock mass

k_n : normal stiffness of a fracture

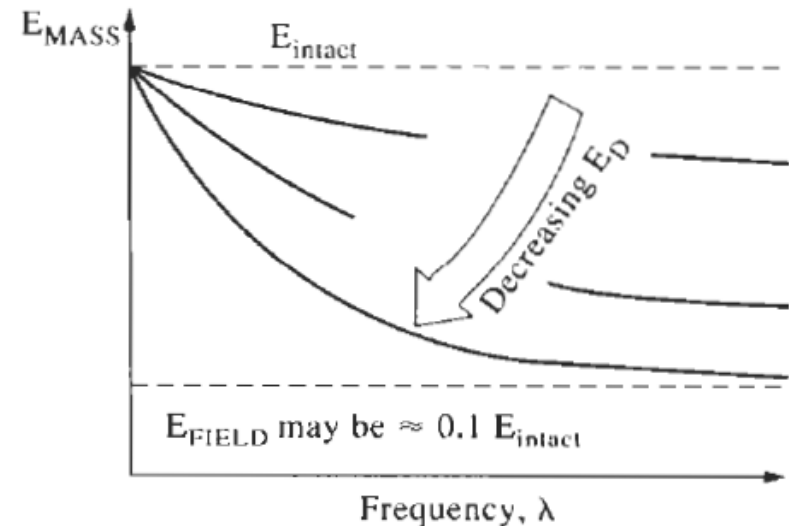
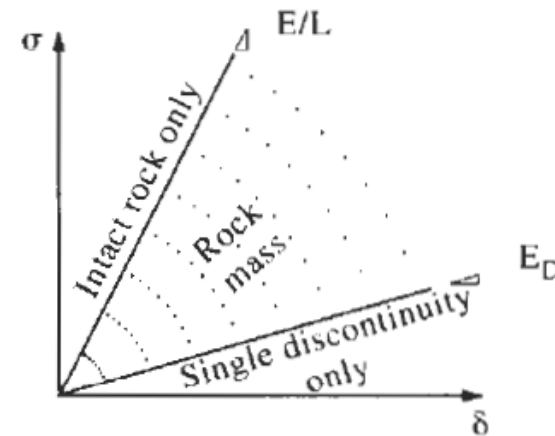
S : spacing of fractures

$$\frac{1}{G_m} = \left(\frac{1}{G_i} + \frac{1}{k_s \cdot S} \right)$$

G_m : Shear modulus of rock mass

k_s : shear stiffness of a fracture

S : spacing of fractures



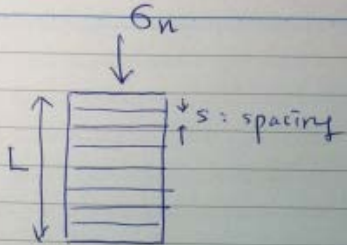
Fractured Rock Masses

Elastic modulus (deformation modulus)



- Derivation

DATE



σ_n
 L
 s : spacing
 σ_n
 number of fracture = $\frac{L}{s}$

E_c : elastic modulus of intact rock
 E_m : " " fractured rock mass

$\sigma_n = k_n \cdot \delta$

normal displacement of a single fracture = $\frac{\sigma_n}{k_n}$
 " multiple fracture = $\frac{L}{s} \times \frac{\sigma_n}{k_n}$

total displacement of fractured rock mass (δ_t) = $L \cdot \frac{\sigma_n}{E_c} + \frac{L}{s} \cdot \frac{\sigma_n}{k_n}$
 (displ of intact rock) (displ of fractures)

total strain (ϵ_t) = $\epsilon_c + \epsilon_f$
 $= \frac{\sigma_n}{E_c} + \frac{1}{s} \frac{\sigma_n}{k_n}$

$E_m = \frac{\sigma_n}{\epsilon_t} = \frac{\sigma_n}{\frac{\sigma_n}{E_c} + \frac{1}{s} \frac{\sigma_n}{k_n}}$

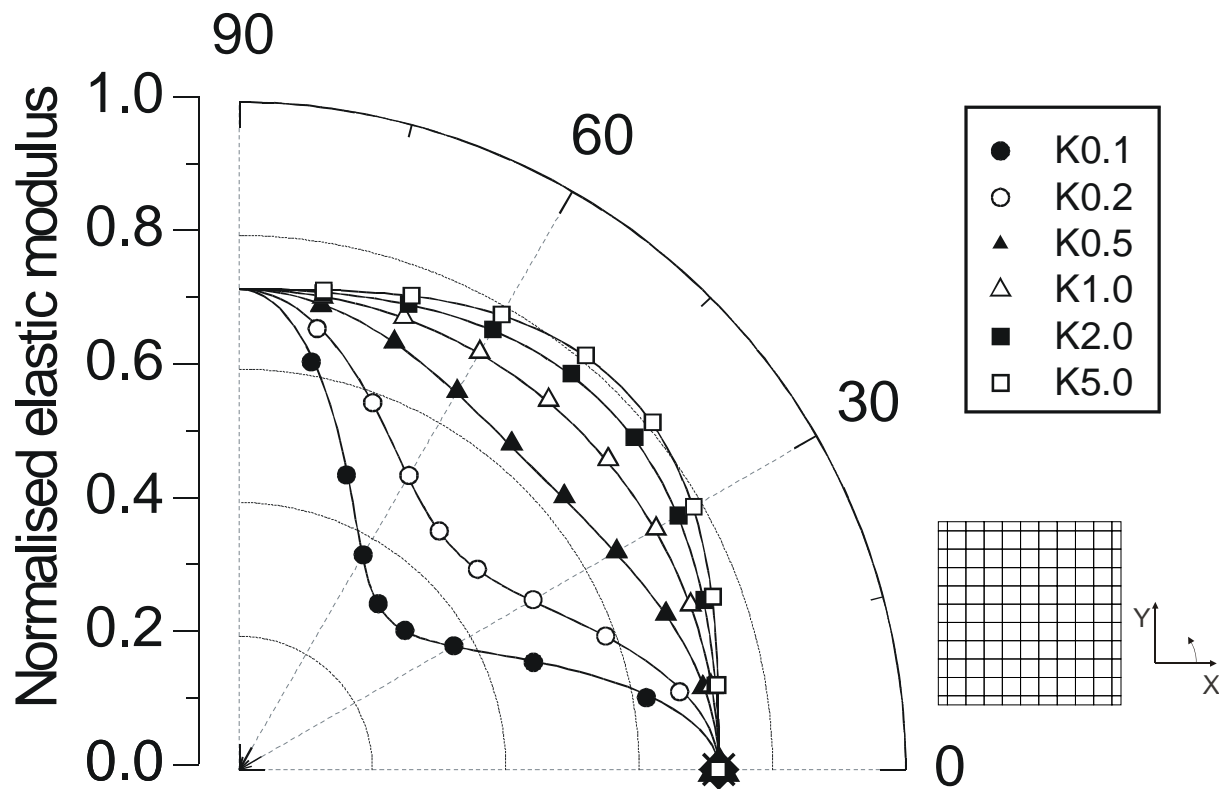
$\frac{1}{E_m} = \frac{1}{E_c} + \frac{1}{s k_n}$ or $\frac{1}{E_m} = \frac{1}{E_c} + \frac{\lambda}{k_n}$

Fractured Rock Masses

Elastic modulus (deformation modulus) - anisotropy



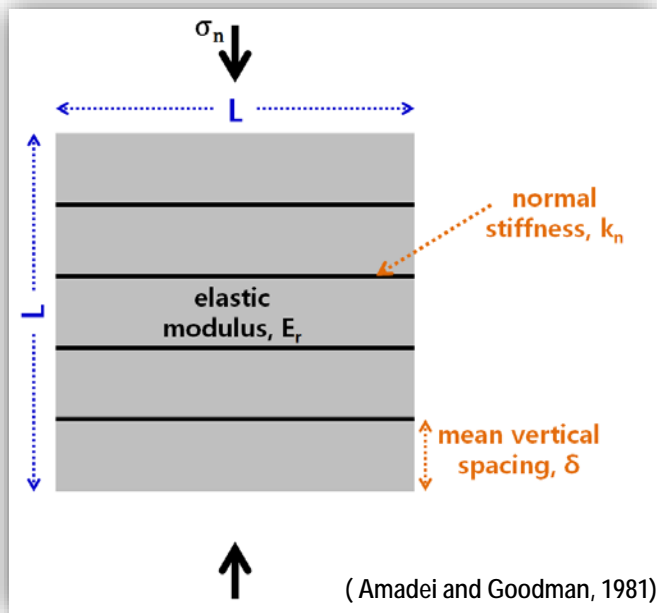
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Fractured Rock Masses Data from Forsmark, Sweden



Equivalent Continuum Model



$$\frac{1}{E_e} = \left(\frac{1}{E_r} + \frac{1}{k_n \cdot S} \right), \quad \frac{1}{G_e} = \left(\frac{1}{G_r} + \frac{1}{k_s \cdot S} \right)$$

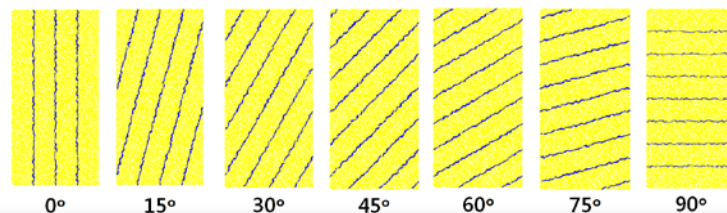
E_e, G_e : equivalent elastic & shear modulus,
 E_r, G_r : intact rock elastic & shear modulus,
 k_n, k_s : normal & shear stiffness on weak planes,
 S : mean vertical spacing

K: shear to normal stiffness ratio

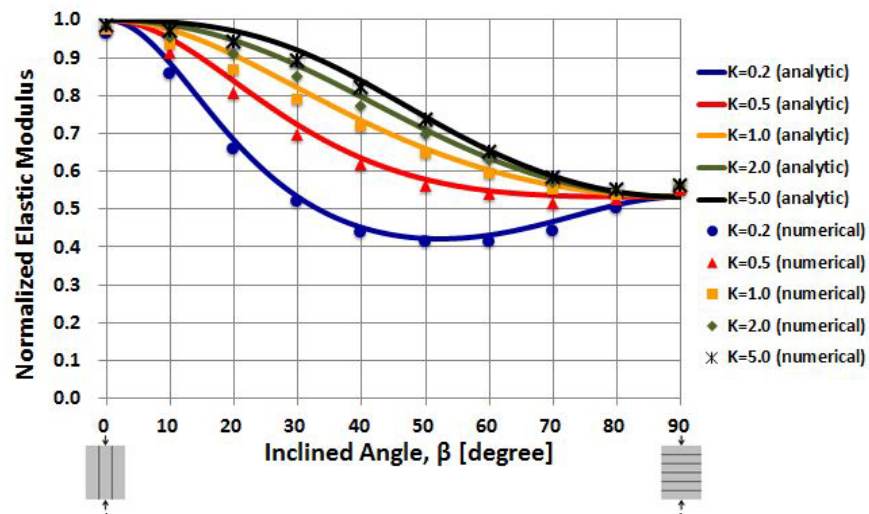
Tensor Transformation

$$\frac{1}{E_\beta} = \frac{1}{E_r} + \cos^2 \beta \cdot \left(\frac{\cos^2 \beta}{k_n \cdot \delta} + \frac{\sin^2 \beta}{k_s \cdot \delta} \right)$$

(E_β varies with respect to inclination, β)



Elastic Modulus of Smooth Joint Model

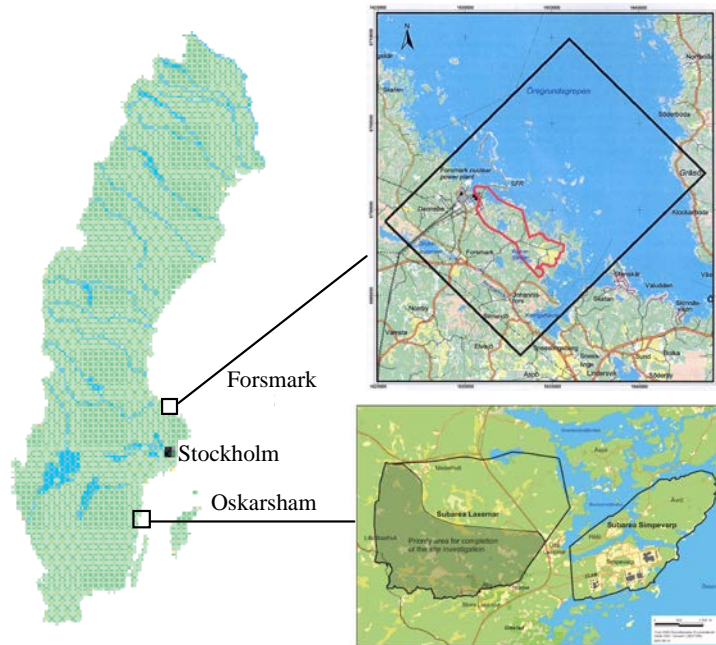


Fractured Rock Masses

Elastic modulus (example from Forsmark, Sweden)



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Map of Sweden



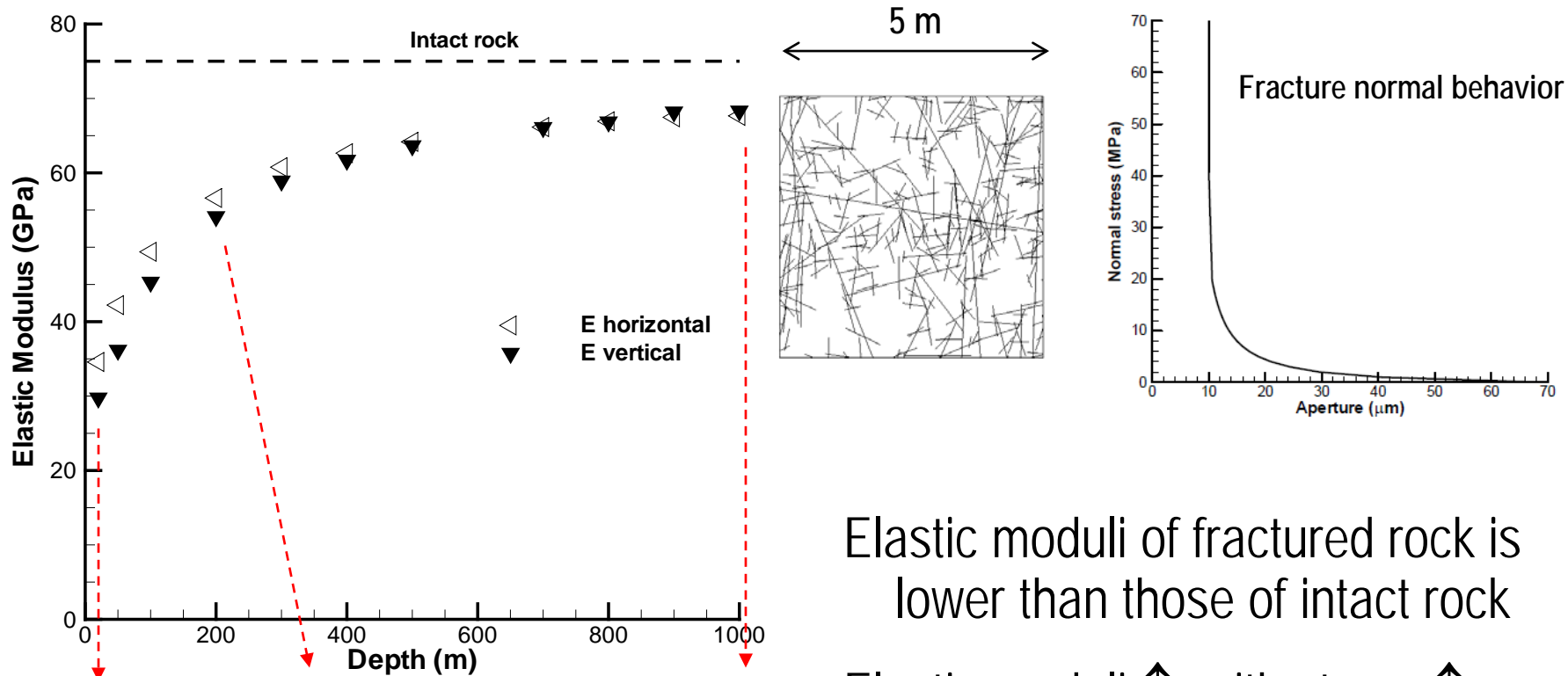
- Swedish geological repository of nuclear waste
- Forsmark and Oskarshamn, two candidate sites for Swedish Program.
- 2002-2009: site investigation.
- 2009: Forsmark, as the final site
- 2011: License application
- 2015: decision (?)
- 2025: Operation

Fractured Rock Masses

Elastic modulus (example from Forsmark, Sweden)



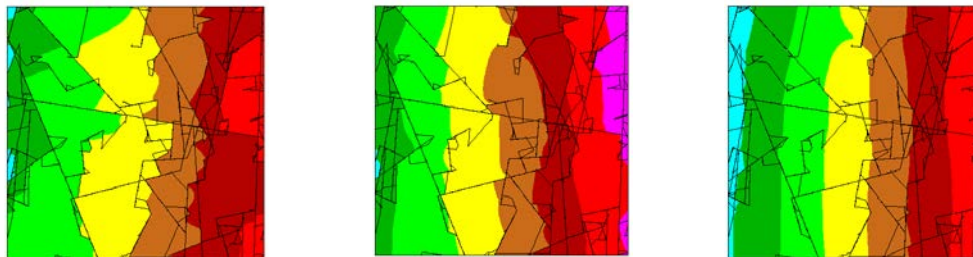
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Elastic moduli of fractured rock is lower than those of intact rock

Elastic moduli \uparrow with stress \uparrow
- highly stress-dependent

Effect of fracture/stress is more evident in low stress condition.



Displacement distribution

Fractured Rock Masses Strength



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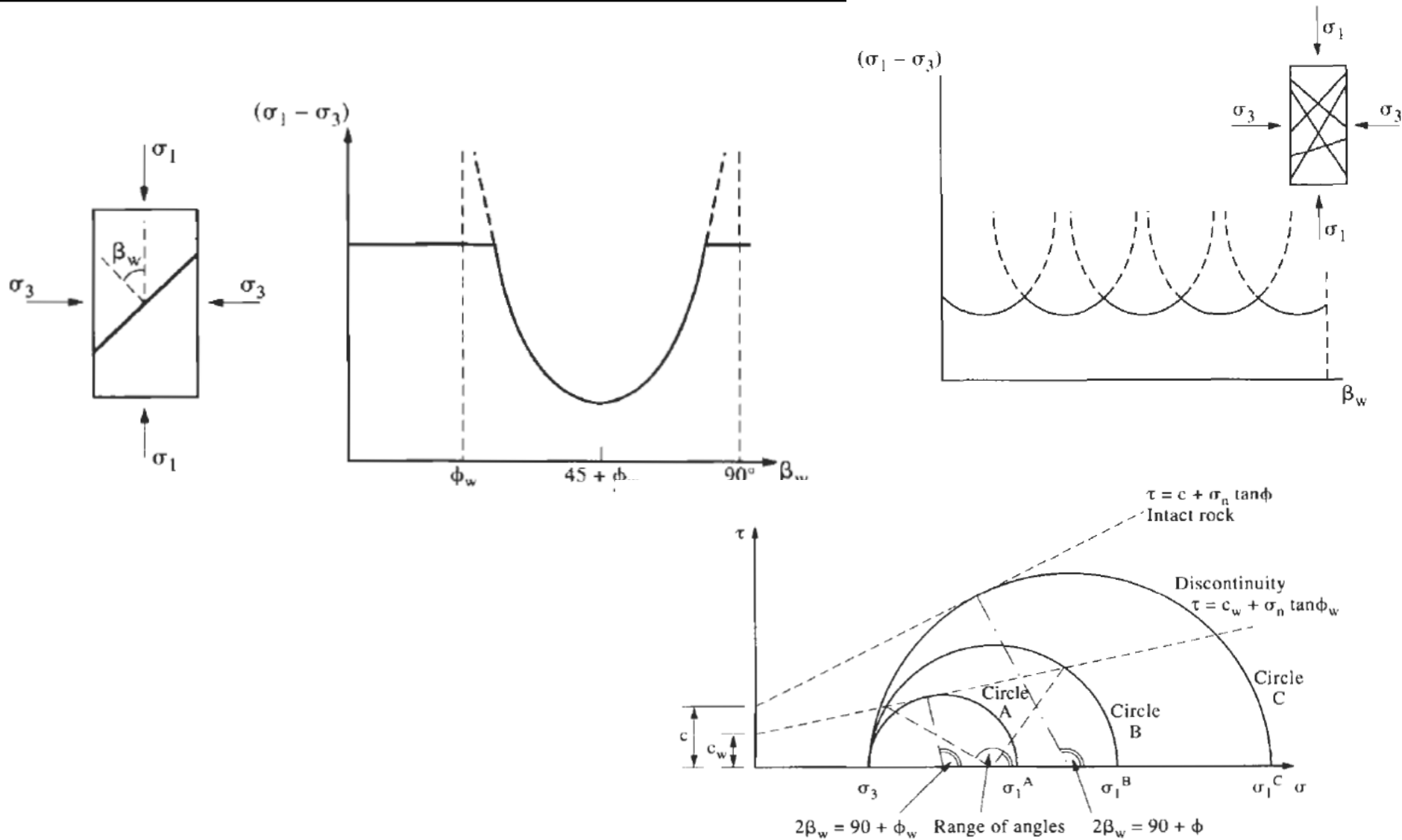


Figure 8.5 Mohr's circle representation of the possible modes of failure for rock containing a single plane of weakness.